# Using near infrared photography to link spatial patterns in stratigraphy with stability

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### Limited research on non-persistent weak layers

A majority of avalanche research has focused on persistent layers. In comparison, we know very little about non-persistent weak layers despite their prevlance in ski areas and other areas with predominantly artificial triggers. One study (Brown and Jamieson 2006) finds that failure layers in new snow are less dense and composed of larger crystals than adjacent strata, while another (Stethem and Perla 1980) finds failure layers in new snow are denser than the slab, but that crystal types are often the same. Micropenetrometer studies of new snow (Pietzch 2009; Peilmier 2003) and observations of atmospheric forms (Casson et al. 2008) show that new snow is highly stratified, especially in wind exposed areas like avalanche starting zones. Fist hardness new snow, classified as a uniform layer by traditional methods, can actually vary in hardness by an order of magnitude over 10cm (Pietzch 2009).

The competing effects of stress and compaction from overburden cause substantial changes within hours (Birkeland et al. 2006). The dyanimic nature of the position and strength of new snow requires a new approach to stability modeling. Stability models that assume the weak

### nIR measurements at two crown faces that failed in the storm layer

layer exists in a static location are not correct. For new snow avalanches, preliminary results suggest failure layers are often not substantially different optically than the slab. Because of the high proportion of avalanches that fail at the interface or in storm snow, we propose a spatially distributed stability index model that examines stability at all depths to predict small to large avalanches for ski areas.



## 4/23/2010 @ 7:30 Skateboard Park



	Snow Pit Profile Observer: Ned Bair	Stability on similar slopes: Good Stability Test Notes:	Layer notes:
	MMSA – Skateboard ParkFri Apr 23 14:38:00 PDT 2010	Air Temperature: <b>4.6 C</b>	64-106: 1A,B,C@95cm
	Elevation (m) 3035 Slope: 38	Sky Cover: Crear Precipitation: None	65-64: Problematic Lay
2.5	Aspect: 66 Wind loading: previous	Wind: Calm	40-63: 2A,B,C@63cm
	Specifics: Ski Area Pit; We skied slope. Ski tracks on slope.		0-40: 3A,B,C@20cm
	Notes: dug 2 meters uphill from crown line fx for Skateboard	Park, R1D1, 0.5m max crown, 3m wide slab, traveled 50m.	
	-10 -9 -8 -7 -6 -5 -4 -3 -2	-1 0 Crystąl ρ Stabilit	y Tests
		Temp C Form Size (mm) kg/m <sup>3</sup>	
	-	- 100	
2	-		
	-	- 0.3 - 0.5 M	
	-	- 80	
4 5			
1.5			
	-	- 70	

Occurrence record estimate:
Type: Soft slab
Trigger: 1 kg handcharge by
author
Class: R1D1
Bed surface: in new snow
3m wide
ran 50m

#### 3/13/2010 @ 6:25 Comeback Cliffs 5 photographed on 3/14/2010 @ 13:47

Occurrence record estimate:
Type: Soft slab
Trigger: ski cut
Class : R2D1
Bed surface: old/new snow inter-
face
7m wide

# effective grain diameter (6/SSA), mm (1m left): ECT 1/NP @ 18cm,

notes: the failure did not occur at the obvious meltfreeze crust or on the graupel layer on top of it

50 cm max crown

30 cm max crown estimate

ran 35m

new wind blown snow made it difficult to find the bed surface during examination, but I concluded it was in the storm layer because of the lack of significant change in grain size.



The radiative transfer problem from a snow pit wall can be solved by considering that snow acts as a collection of equivalent sized spheres (Warren and Wiscombe 1980). My nIR profiles suggest that these spheres are the same size in failure layers and adjacent strata for non-persistent weak layers. Since density does not affect albedo (Bohren 1975), it's possible that these grains are the same size, but packed differently. Packing differences could be caused by different atmospheric forms and the influence of wind. I propose that packing differences are largely responsible for formation of nonpersistent weak layers.

To investigate and model non-persistent weak layers, I will:

1) make measurements of failure layers at crown faces and at protected study plots using an nIR camera, a snow micro-penetrometer, and traditional snow pit observations such as detailed temperatures and densities.

2) investigate a snow camera which photographs flakes as they fall to determine how atmospheric form affects packing.

### Research plan:

### Measurements:

Stability tests to identify weak layer depth: •Extended column •Tilt-board

Instruments to measure strength and stratigraphy: Snow micropenetrometer •nIR camera disdrometer (snowflake camera) •traditional methods (i.e. hand lens, 250cc density cutter)

## Models:

•Geostatistical model to examine spatial varation in weak layer location •Continuum stability index model to examine fractures within storm layer •Spatially distributed stability index driven by a wind model

3) develop a continuum stability model which models the dynamic vertical evolution of nonpersistent weak layers.

4) investigate development of a spatially distributed stability index that uses a wind model, SnowTran-3D (Liston et al. 2007), driven by a high-resolution (1m<sup>2</sup>) DEM, 7-8 anemometers, and 2 precipitation stations over Mammoth Mountain (1600 hectares).

### Software:

•Develop open source (MATLAB) package to process and calibrate near infrared digital images, possibly eliminating frame and flat-field correction requirements

Research questions: 1) Do non-persistent weak layers form at different depths throughout the course of a storm?

2) For microstructure, how does variability within a storm slab compare to the differences between a failure layer and a slab?

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