

REFERENCES

- Bayne, C. K. 1958. Geology, Mineral Resources and Ground-water Resources of Elk County, Kansas. Volume 14, State Geological Survey of Kansas, Lawrence.
- Bradford, J. M., and R. F. Piest. 1980. Erosional Development of Valley-bottom Gullies in the Upper Midwestern United States. In: Thresholds - in Geomorphology, George Allen and Unwin, Boston. pp. 75-101.
- Daniels, R. B., and R. H. Jordan. 1966. Physiographic History of the Soils, Entrenched Stream Systems, and Gullies, Harrison County, Iowa. Tech. Bull. 1348, Soil Conservation Service, U.S. Dept. of Agri. 103 p.
- Flint, R. F. 1957. Glacial and Pleistocene Geology. John Wiley and Sons, New York. 553 pp.
- Frye, J. C., and A. B. Leonard. 1952. Pleistocene Geology of Kansas. Bull. 99, State Geological Survey of Kansas, Lawrence. 223 p.
- Grissinger, E. H. Unpublished data.
- Jamkhindikar, S. 1967. Sedimentary Characteristics of Pleistocene Deposits, Neosho River Valley, Southeastern Kansas. Bull. 187, Part 5, State Geological Survey of Kansas, Lawrence. 13 p.
- Knox, J. C. 1975. Concept of the Graded Stream. In: Theories of Landform Development, State University of New York, Binghamton, pp. 169-198.
- Ruhe, R. V. 1969. Quaternary Landscapes in Iowa. Iowa State University Press, Ames, Iowa. 256 p.
- Soil Conservation Service, and U.S. Dept. of Agriculture 1975. Soil Survey of Butler County, Kansas. U.S. Gov't. Printing Office, Washington, D.C. 142 p.
- Wendland, W. M., and R. A. Bryson. 1974. Dating Climatic Episodes of the Holocene. Quaternary Res. 4: 9-24.

PRELIMINARY ASSESSMENT OF CHANNEL CHANGE, ELK RIVER, KANSAS

Peter F. Lagasse
Assistant Dean for Research, United State Military Academy,
West Point, New York, U.S.A.

INTRODUCTION

This paper considers the geomorphic response of the Longton reach of the Elk River to anticipated modifications in stream flow and sediment load as a result of construction of a number of small flood retention reservoirs in the upstream watershed. Within the constraints of data and time for problem analysis, only a preliminary qualitative assessment of response is attempted, but a methodology is outlined that will provide both quantitative results and a reasonable prediction of anticipated response. The approach has been applied to river systems as large as the Upper and Lower Mississippi River and as small as ephemeral arroyos in New Mexico, with excellent results. The methodology is currently being applied to an analysis of the response of the Cochiti to Isleta reach of the Rio Grande in New Mexico to the construction of the main-stem dam at Cochiti in 1973. As results of this latter study are pertinent to the Workshop theme of downstream river channel changes from diversions and reservoir construction, they will be available during the general discussion session of the Workshop to illustrate the methodology outlined in this paper.

METHODOLOGY

The approach recommended for analysis of Elk River response has been described in detail and applied to the Upper Mississippi River in a reference document, "The River Environment," prepared for the Fish and Wildlife Service, Twin Cities, Minnesota by Simons, Lagasse, Chen, and Schumm in 1975. It was refined further relative to an "Assessment of Geomorphic Response of River Systems to Hydraulic Structures" in a paper I prepared for an International Symposium on "Environmental Effects of Hydraulic Engineering Works," held at Knoxville, Tennessee, 12-14 September 1978. The brief outline of the methodology which follows is extracted from these references.

Depending on the data and resources available for analysis the problem of response of the Elk River to the construction of a number of small hydraulic structures should be approached in three phases.

1. A qualitative analysis based on general geomorphic parameters.
2. A quantitative analysis based on specific geomorphic and hydraulic data.

3. A mathematical model of watershed and channel processes in the reach or system of concern.

As listed, each phase requires an increasing commitment of resources, but individually each phase yields meaningful results. These range from a purely qualitative assessment of trends to the numerical results and predictive capability of physical process computer modeling. When applied sequentially this multi-phase approach constitutes a powerful methodology for the evaluation of short and long range response of river systems to development.

To the extent that data permits it is desirable to establish, first, the morphologic and hydraulic conditions of the river or reach under consideration before man's intervention. Conditions on the "natural" river form a baseline against which the impacts of man's activity can be assessed. Unfortunately, the historic record and data are usually not sufficient to establish a complete picture of the natural river, except on major systems such as the Mississippi or Rio Grande. As a minimum, it is usually possible to reconstruct the history of engineering activity on a river or reach of concern. Although data on the Elk River provided for the Workshop do not provide a complete picture of engineering activity in the watershed, indications are that such information could be developed with a minimal investment of research effort.

A qualitative analysis of geomorphic response should include an examination of the river in planform, longitudinal profile, and cross section. Where data are available for different time periods (for example, before and after construction of a dam), this analysis produces a time-sequenced picture of morphologic change in three dimensions which can be correlated with the history of engineering activity in the study reach. This correlation provides a qualitative assessment, in terms of trends, of the impacts of man's activity in the reach. In systems that have experienced multiple development techniques (dredging, dikes, navigation dams, levees) an attempt can then be made to isolate the system response to a particular activity of concern, or to predict response to hydraulic structures and development measures being considered.

Township plats normally provide the earliest accurate planform data on a river system. Comparison of these with later topographic surveys, aerial photographs, or the current USGS quadrangle sheets establishes the degree of bankline stability in a system. Even within relatively stable banklines most alluvial rivers exhibit changes in bankline and the number, location and configuration of bars and islands. As these features influence resistance to flow and act as controls in a river reach, an understanding of their evolution is imperative.

Time-sequenced comparison of selected reaches and measuring and comparison of river widths, island area, and river bed area, all provide useful data for establishing and evaluating morphologic characteristics of an alluvial river. Graphical, tabular, and plan view representation of this data provide insight into the evolution of alluvial processes of a particular region and their relation to man's development in the

region. The impact of such development techniques as bankline stabilization, contraction dikes, and jetty fields is usually quite apparent in the planform comparisons.

While longitudinal profiles are normally not directly available, they can be constructed from cross section data derived from hydrographic surveys. Comparisons of profiles can be made using either the average depth of the cross section or the thalweg depth, that is, the deepest point in the cross section. Evidence from analysis of the Upper Mississippi indicates that thalweg bed elevations, in general, vary in the same manner as average river bed elevations, and so provide a good, readily obtainable indicator of trends in bed elevation. Both tabular compilations and time-sequenced longitudinal profiles are useful in establishing trends in aggradation or degradation.

The tabulation and plotting of cross section data reveal changes in such morphologic parameters as surface width, bed elevation, average depth, and thalweg position. Cross sections can be established at equal intervals throughout a reach or can be concentrated in those portions of the reach that planform analysis has shown to be morphologically active. Comparison of cross sections adds a third dimension to the qualitative analysis and normally provides the clearest indicator of the impact of such stabilization measures as revetment, contraction dikes, and jetty fields.

In correlating the planform, longitudinal, and cross-sectional data accumulated by these and other techniques, it is not uncommon to encounter apparently conflicting indicators. Here, an understanding of the natural morphologic conditions of the system provides an essential baseline for interpreting apparently anomalous behavior. Efforts directed at synthesizing the geologic history of the region and identifying existing geologic controls, as well as the historical research required to provide clues as to the natural conditions of the river system, often pay unexpected dividends in this regard.

While a qualitative analysis, alone, will yield meaningful results, refinement of conclusions resulting from such an analysis and a more precise assessment of trends as well as a predictive capability can be derived from both quantitative analysis and mathematical modeling. It should be noted, that calibration of a mathematical model involves evaluation and modification of supplementary relations to the basic process equations using field data or theory so that the model will reproduce the historical response of the modeled river system. Establishing the trends in geomorphic components which constitute this historical response by a qualitative analysis, then, provides an essential base for model calibration.

The conventional or traditional quantitative techniques of river engineering are well known and do not require discussion here, however, a few general comments on these techniques in relation to geomorphic analysis and mathematical modeling may be appropriate. Quantitative techniques include using unit hydrographs for water routing and yield from watersheds, the Universal Soil Loss Equation for estimating erosion from watersheds, time-lag methods for flood routing in the channel,

sediment rating curve techniques for estimating deposition in reservoirs, and developing relationships for hydraulic geometry of the reach under study. These techniques can be applied to only a relatively small number of conditions and alternatives because of cost limitations for most studies. It is often difficult to predict the response of a system to development alternatives using these methods as they are based on the assumption of homogeneity in time and space. While this approach may be feasible for selected reaches, application on a system-wide basis quickly leads to unmanageable computational requirements. Such dynamic features as degradation and aggradation are difficult to account for, and integrating the subsets of a traditional quantitative analysis into a basin-wide system analysis presents complex problems. The traditional quantitative techniques are, however, extremely valuable in either refining or substantiating the conclusions of a qualitative analysis, or in providing a basis for mathematical model calibration.

To conduct an analysis with a physical process computer model, several major tasks must be accomplished. The first is to develop a mathematical model for routing water and sediment from the watershed and through the existing and modified channel and reservoir systems. This model could be used to determine flow lines and identify areas where excessive sediment aggradation and degradation may occur, as well as to indicate major sources of sediments that flow from tributary systems to the main channel. With such a comprehensive model developed and verified, the second task is to evaluate various operational alternatives or development scenarios. The final task involves selecting an optimum plan for the development or operation of the river basin considering flood control, sediment control, minimizing environmental impact, maximizing water salvage, and other factors of concern for a particular region.

The detailed development of such a mathematical model involves the following steps: data assembly and inventory; data evaluation; development of a data storage and retrieval system; collection of required data that cannot be synthesized; identification of data gaps and synthesis of additional data required for analysis; overall system design including spatial and temporal design, subsystem model development such as the main-stem model, tributary models and watershed models; validation and linking of the subsystem models; development of application data files; model calibration and validation; application of the models to evaluate system response for different design alternatives; and finally, conducting a detailed analysis of selected alternatives.

ANALYSIS

If the analysis of the Elk River were to concentrate on only the Longton reach, that is the 8 river miles between Sections 11-15 and 3-3, it is doubtful that much more than a qualitative geomorphic assessment of response to the proposed flood retention reservoirs could be justified. Selected aspects of the trends revealed by this geomorphic analysis could be verified by quantitative calculations. However, if a complete analysis of the response of the Elk River system to the construction of the 48 floodwater retarding structures is desired, the application of physical process computer modeling would be required.

With the data and time presently available, and considering just the eight miles of the Longton reach, only a preliminary qualitative assessment of geomorphic response will be attempted here. This should provide a basis for more detailed analysis and discussion during the Workshop.

Control of floodwater run off from 59% of the Elk River drainage area and retention of the expected 100 year accumulation of sediment from this area will induce geomorphic change along the Elk River, including the Longton reach. On each of the many tributaries with retention dams one would expect some tendency toward the classic response of a river to dam construction. As a result of clear water release local effects below each dam could include local scour at the dam, channel degradation below the dam, and possible bank instability if significant degradation takes place. The downstream effects could include degradation and bank instabilities along the entire reach between the retention structure and the junction of the tributary with the Elk River. Over the short term this could produce an increase in sediment supply from these tributaries to the main-stem, even though the dam might trap virtually all the sediment produced from the upstream watershed. The magnitude of any short term increase in sediment load depends on resistance of each tributary reach to degradation. This resistance would come from bed and banks of cohesive material, bed rock or geologic control, or armoring of alluvial reaches of the tributary.

As each tributary adjusts to regulated flows of water and sediment over the long-term, and cohesive materials, or bed rock, or a well developed armor layer become sufficient to resist scour by regulated flows, a net reduction in sediment supply to the Elk River should be anticipated. (Note pre- and post-project estimates of Elk River sediment yield in the data package.) The magnitude and duration of any short term increase in sediment supply from the tributaries following construction of a retention dam and the development of a stable channel in the reach below the structure will depend on individual channel and watershed characteristics. Based on the characteristics of bed, bed material, and armor layer of the Elk River cross-sections described in detail in the data package, and information in the soil survey of Chautauqua County, any initial increase in sediment load from the tributaries should be of short duration, and degradation in tributary reaches below retention structures should be limited.

For the Elk River main-stem, the cumulative effect of this sequence of geomorphic change on multiple tributaries could be significant. There is potential for both short- and long-term change in meander pattern and planform configuration, composition of the bed, and the riffle/pool sequence along the Elk River. With specific reference to the Longton reach the magnitude of this change will depend, to some extent, on the ability of the channel to absorb and redistribute any short term increase in tributary sediment load, but primarily on the resistance of the reach to a long term tendency toward degradation because of reduced sediment inflow from the tributaries.

The characteristics of the bed, bed load, armor layer and banks described for Sections 11-15 to 3-3 of the Longton reach indicate that

this reach should be quite resistant to both degradation and planform change through bank erosion. Of the sections described in detail most show a geologic control (limestone, rock, shale) either exposed in the section or a few feet below an armored alluvial bed. For sections such as 3-1, 3-3, and 11-9, where the bed material is apparently alluvium, the resistance to degradation will depend on development and stability of an armor layer. With indications of an armor layer at 7 of the 13 sections described in the data package, there is sufficient evidence to assume that the channel bed contains the necessary gradation and quantity of coarse material to produce armoring as degradation progresses. However, the response of the Elk River to changes in discharge of sediment and water on multiple tributaries will be so complex that a quantitative determination of the actual amount of degradation to be anticipated in alluvial reaches would require physical process computer modeling of the watershed, tributary, main-stem system. A quantitative approach, short of modeling, would probably not produce an estimate any more reliable than that derived from a purely qualitative assessment. From the morphologic characteristics of the Longton reach one would expect no more than 2 or 3 feet of degradation before either bed rock control or armoring produces stability against the reduced flood peaks of controlled flows.

There are several techniques in the literature for computation of river bed degradation as a result of altered sediment regime, but the assumptions required in their application to a field situation are quite limiting. For example, Komura and Simons developed a technique for calculating "River-Bed Degradation Below Dams" in 1967 but the assumptions required in the numerical example include (with an indication of applicability to the Elk River problem):

1. Sediment transport is completely arrested by the dam (OK),
2. River banks are not erodible (No),
3. Seasonal variations in discharge and temperature of water do not occur (No),
4. Sediment injections by tributaries do not occur (No), and
5. Meandering and growth of vegetation do not occur (No).

Similarly, the USBR "Design of Small Dams" contains an approach for estimating degradation and armoring which places heavy reliance on "engineering judgement" and limiting assumptions.

In the case of the Longton reach of the Elk River the key factor in the analysis must be the influence on the main-stem of altered tributary flow conditions. As with the Komura and Simons approach, most quantitative techniques for estimating degradation and stability through armoring cannot handle this complexity. Important tributaries to the Longton reach include: Wildcat Creek - 2 miles above Section 11-15, Clear Creek - .5 miles above Section 11-15, Hitchen Creek at Section 11-1, Painterhood Creek below Section 3-1; and several smaller unnamed tributaries in the vicinity of Sections 11-11 and 11-15. Under

"natural" flow conditions (prior to the construction of flood detention dams) one would expect that slope and sediment inflow of these tributaries would be adjusted to the existing base level of the main-stem. While delta deposits from flood flows on the tributaries might temporarily divert or control the base level of the main-stem, these deposits would normally be redistributed during flood flows on the Elk River proper.

Under post-project conditions on the Elk River (as illustrated by the unit discharge hydrograph in the data package) the tributaries could assume a dominant role in controlling base level. This would be particularly true during any initial period of degradation below structures on the tributaries, when regulated flows on the main-stem may be incapable of moving either the size or volume of material deposited in tributary deltas.

As a case in point, on-going analysis of the response of the Rio Grande to the construction of the main-stem dam at Cochiti reveals that degradational processes have been far more complex than would be predicted by available quantitative techniques. The "classic" degradational wedge, deepest at the dam and tailing out at some downstream geologic control, has not developed. Instead, the initial 8 miles below the dam have shown remarkable stability, apparently because of the inability of regulated flows to move the size or volume of material in numerous arroyo deltas in the reach. The availability of significant quantities of gravel in these deltas has resulted in development of a stable armor layer in the reach. Downstream, beyond the influence of these arroyos, 6 to 8 feet of degradation has occurred.

Over the short-term, then, planform change could be expected as a result of diversion or blocking of the Elk River at tributary junctions. With time, redistribution of this material could alter the existing riffle and pool sequence as "slugs" or waves of deltaic sedimentary material are moved through the system. High flows will be comparatively rare, and extended periods of low flow will scour the crossings and fill the pools of the existing meander sequence. As tributaries adjust over the long term to altered flow conditions a period of degradation can be expected in reaches such as the Longton reach. A fairly recent (natural?) cut off of a meander bend is evident below section 11-11. Degradation could induce bank instability with a potential for additional meander loop cut offs at several locations in the reach. This would produce a radical change in slope, velocity, and transport capacity and, again, alter the riffle/pool sequence. However, degradation should be limited in most sections of the Longton reach to no more than 2 to 3 feet by either geologic controls such as the limestone ledge at Section 13-2 (photo 25-2) or development of an armor layer. Once developed, an armor layer should be sufficient to provide stability under conditions of reduced main stem flows. There are several locations in the Longton reach (Sections 11-15, 2180, 11-1, 3-5, 3-4) where degradation could expose rock, shale, or limestone, thus altering the existing substrate material as well as modifying the riffle and pool sequence.

SUMMARY

In summary, there is potential for geomorphic change in the Longton reach of the Elk River as a result of construction of flood retention structures on tributaries. While a qualitative assessment can indicate possible trends (aggradation/degradation, stability/instability) with reasonable assurance, a complete quantitative assessment of response would require application of the multiphase approach outlined at the outset of this paper. Response to altered conditions of water and sediment flow on 48 tributary reaches and 59% of the drainage area will be complex, and will require computer modeling to refine initial conclusions derived from qualitative analysis.

Although the data package provided for this analysis was reasonably complete, additional information concerning bank material and bankline stability would have been useful. With the Elk River system "fixed" in the vertical dimension by numerous geologic controls, bankline stability becomes the crucial indicator of geomorphic stability in terms of the meander pattern and the riffle and pool sequence. Along these lines, a site visit must be considered an absolutely essential element of analysis for any river related problem. Results of any analysis, even at the qualitative level, must be considered tentative until the analysis is supported by at least one day in the field on-site. For example, the preceding qualitative assessment assumed that bed rock control was dominant in the Longton reach, that the tributaries were similar to the main-stem Elk River in terms of geologic controls, alluvium, and bankline vegetation, and that active measures to insure the integrity of bankline vegetation were in effect or planned. If a site visit demonstrated that any of these assumptions was in error, the conclusions of this preliminary qualitative assessment would require revision.

An additional technique that I consider essential for analysis of response in any river system is a literature search for well-documented case studies of river response to similar engineering activity in related physiographic settings. This is particularly important where conclusions or projections of response are required using only a limited data base, as was the case with the Elk River. For example, Haug (1977) analyzed the response to large impoundment structures of seven major streams in Kansas. Qualitative methods of fluvial geomorphology as outlined in this paper and hydraulic engineering techniques (quantitative methods) were applied to the problem of river response, to include analysis of the Fall River just to the north of the Elk River watershed. While response to a single large impoundment structure on each river system was analyzed, several conclusions derived from Haug's study are of interest in anticipating response of the Elk River system to multiple small impoundment structures. Degradation below the impoundment structure was experienced on all streams considered by Huang, and in all cases the stream below the structure tended to form a relatively narrower and deeper channel, that is, width to depth ratio decreased.

CONCLUSIONS

The following specific conclusions on the Elk River problem are listed by categories for comparison with conclusions from other Elk River papers.

Aggradation/Degradation

Over the short-term some aggradation, particularly in the vicinity of tributary junctions, should be anticipated as tributaries respond to the construction of detention dams. After tributaries have adjusted to altered flow and sediment regimes, degradation along the main-stem Elk River including the Longton reach can be anticipated. Because of geologic controls and the probability of armoring in alluvial reaches, this degradation should not exceed 2 to 3 feet. In fact, geologic controls will limit degradation to alluvial reaches where lowering of bed elevations might be more accurately characterized as "local scour."

Bed Material Size

Over the long-term bed material size will increase along the Elk River and in the Longton reach as fines are removed and armoring develops through hydraulic sorting. In terms of altered substrate conditions reaches where a geologic control is covered by a thin veneer of alluvium could be swept clean.

Bank Stability and Width

Two fundamental assumptions of this analysis have been that the banks are composed of predominately cohesive materials and are (and will remain) stabilized by vegetation. Unless a site visit were to contradict these assumptions, channel banks, and width can be considered stable in most reaches. Degradation in alluvial reaches might undercut bankline vegetation and produce local instances of channel widening.

Width to Depth Ratio

With stable banklines and limited degradation, width to depth ratios should change only slightly. With reference to Huang's case studies of response of Kansas streams to impoundment a slight increase in width to depth ratio should be anticipated in some alluvial reaches.

Pool and Riffle Spacing

Redistribution of material produced by short-term tributary adjustment to detention structures could alter the composition of bed materials in the riffle and pool sequence as "slugs" of deltaic sediments move along the main-stem. Over the long-term under regulated conditions, high flows will be comparatively rare and extended periods of low flow will tend to scour the crossings and fill the pools of the existing meander sequence; however the spacing of the riffle and pool sequence would not be altered by this process.

Sinuosity Change

There is evidence of at least one recent (?) natural (?) cut off in the Longton reach. A site visit would be required to determine age and origin, but cut off of a meander loop could produce a radical change in slope, velocity, and transport capacity of the reach. Cut off a meander loop is one geomorphic event that could alter the riffle and pool spacing in the reaches above and below the cut off. Documentation by Stevens (1980) of a flood of record of 200,000 cfs on the Elk River in July 1976 (six times larger than the mean annual flood of 32,800 cfs) indicates a potential mechanism for developing cut offs of the existing meander pattern.

As a result of the analysis of the Elk River problem and discussion during the Workshop several general conclusions are offered.

Site Visit

A site visit is an absolutely essential element of analysis for any river related problem. Results of any analysis, even at the qualitative level, must be considered tentative until supported by reconnaissance in the field. The time required for a site visit depends on the areal extent and complexity of the watershed. For the Elk River, one day on-site to include examination of the main-stem and several representative tributaries might suffice. Aerial reconnaissance is, in most cases, an essential adjunct to a site visit.

Case Studies

Well documented case studies of river response to similar engineering activity in related physiographic settings constitute an important supplement to analysis of river response. This is particularly true where conclusions or projections of response are required using only a limited data base, as was the case with the Elk River. Hydraulic engineering projects can induce major changes in the hydrologic, hydraulic, and sediment regimes of river systems, and at present, theory alone is not capable of predicting this complex response. Research effort committed to producing documented case studies can provide a valuable resource for evaluating river response.

REFERENCES

- Huang, T. 1977. Changes in Channel Geometry and Channel Capacity of Alluvial Streams Below Large Impoundment Structures. Thesis in partial fulfillment of the requirements for the degree of Master of Science, University of Kansas, Lawrence, Kansas.
- Komura, S., and D. B. Simons. 1967. River-Bed Degradation Below Dams. ASCE, Journal of the Hydraulics Division, Vol. 39, HY4, July, pp. 1-14.
- Lagasse, P. F. 1978. Assessment of Geomorphic Response of River Systems to Hydraulic Structures. Proceedings of the International Symposium on Environmental Effects of Hydraulic Engineering Works, Knoxville, Tennessee, IAHR, 12-14 September.
- Simons, D. B., P.F. Lagasse, Y. H. Chen, and S. A. Schumm. 1975. The River Environment. Reference Document for the U.S. Department of the Interior, Fish and Wildlife Service, Twin Cities, Minnesota, December.
- Simons, D. B., S. A. Schumm, M. A. Stevens, Y. H. Chen, and P. F. Lagasse. 1975. A Geomorphic Study of Pools 24, 25, and 26 in the Upper Mississippi and Lower Illinois Rivers. Colorado State University, for the Waterways Experiment Station, Vicksburg, Mississippi, NTIS Report AD-A012 845.
- Stevens, M. A. 1980. Assessment of Anticipated Channel Changes in the Elk River near Longton, Kansas Due to Upstream Developments. U.S. Fish and Wildlife Service Workshop on Downstream River Channel Changes from Diversions or Reservoir Construction, 27-29 August 1980, Fort Collins, Colorado.
- U.S. Bureau of Reclamation. 1977. Design of Small Dams. Water Resources Technical Publication, U.S. Government Printing Office, Washington, D.C.

PREDICTED RESPONSE OF THE ELK RIVER AT LONGTON, KANSAS

David C. Ralston
National Design Engineer, Engineering Staff
Soil Conservation Service, Washington, D.C., U.S.A.

GENERAL INFORMATION

The Elk River is a tributary of the Verdigris River in the Arkansas River basin in southeastern Kansas. The Elk River reach selected to evaluate is near Longton, Kansas and has a drainage area ranging from 285.7 square miles at the upper end (Section 11-15) to 390.5 square miles at the lower end (Section 3-3). This discussion is an analysis and evaluation of the expected response of this reach of river to the installation of 45 floodwater retarding dams in the upstream reaches of the watershed under the USDA Soil Conservation Service small watershed program authorized under P.L. 83-566.

The watershed consists of gently sloping flood plains and steep bluffs of the Flint Hills escarpment in the upper reaches. In the channel study reach, the area consists of thick beds of sandy shale with interbedded limestone ranging from thin beds to thick ledges. Soils of the watershed are thin over the thick limestone ledges. Floodplain soils are deep and friable and are mainly silty clay loams. Land use over the watershed is about 13% cropland, 82% grassland, 3% woodland, and 2% a mixture of other uses. Of the approximately 7,320 acres of woodlands, about 5,350 acres are on the flood plain, primarily in narrow belts adjacent to the Elk River and its tributaries.

The average annual precipitation for Howard, Kansas, located about 12 miles northwest of Longton near the middle of the watershed, is 35.07 inches. The largest total annual precipitation recorded at Howard is 56.07 inches and the smallest is 18.47 inches. Normally, about 75 percent of the precipitation falls during the growing season, April to October. The Elk River floodplain is flooded frequently--two to three flows a year exceed bank-full capacity. Flooding duration in the study reach is usually 24 to 36 hours. Sediment deposition on the flood plain during flooding causes problems in localized areas.

The average growing season is 185 days. An average year would be frost free from 15 April to 17 October. Daily temperatures average 35°F during January and 80°F during July. Extreme temperatures have been above 115°F and below -20°F.

The proposed project is a system of 45 floodwater retarding dams above the study reach to be installed on the major tributary drainages to the Elk River. The dams are to be earthen with vegetated or rock emergency spillways to provide safe passage of the runoff that exceeds the reservoir detention storage capacity. These spillways are planned so that their chance of operation in any 1 year ranges from 4 to < 1 percent. The principal spillways for the dams are of reinforced concrete with a crest at the elevation of the 100-year accumulation of

sediment and have an uncontrolled release rate of 20 cubic feet per second per square mile of contributing drainage area (CSM).

Reservoir detention storage capacity is planned to handle from 3.0 to 5.25 inches of runoff from the contributing drainage area. The storage allocated for sediment accumulation ranges from 0.51 to 1.79 inches from the contributing drainage area. An ungated orifice through the principal spillway is to be provided at the elevation of the 50-year accumulation of sediment. After depletion of this storage, the orifice can be plugged and submerged storage is then available below the principal spillway crest elevation.

The State Geological Survey of Kansas issued a report in July 1958 on the rock formations and mineral and ground-water resources for Elk County where the study reach is located. The report states that the flood plain alluvium consists of two strata. The lower stratum is coarse material, predominately chert, limestone, and sandstone gravel and ranges from a fraction of an inch to 10 feet in thickness but generally is about 5 feet thick. Sand is intermingled with the pebbles, some of which are 2 to 3 inches in diameter. These deposits are the better aquifers in the area, even though their yield is not large. Household wells in the area usually yield less than 50 gallons per minute. The lower stratum yields water freely but is not continuous over the entire valley. The upper stratum of alluvium is a deposit of mostly clay and silt which grades downward to more sandy materials.

ANALYSIS

This section describes the concepts and methods used to estimate the future form and substrate of the selected reach (between Sections 3-3 and 11-15). The floodplain along the reach is 6.8 miles long and the stream channel length is 10.2 miles long.

General Geomorphology

To make a prediction of performance, it first is necessary to establish the general state of equilibrium of the stream. This is best done by evaluation of the geomorphic setting and as many quantitative factors as possible that are useful in supporting the identified state of equilibrium.

The general trend is erosional; the channel is gradually cutting into the underlying bedrock. The local irregularities of rock elevation in the bed and of depth of alluvium are due to different degrees of erosion resistance and bedding thickness. Within this reach, the stream flows across the Lawrence Shale Formation, which has a thick limestone stratum and a sandstone member of variable thickness. The sandstone member (Ireland) is a significant aquifer for domestic water wells, even though its yield is relatively low (1 to 10 gpm).

Field inspection, probing, and sampling of the channel bed indicate that, with a few exceptions, bedrock is within 48 inches below the stream bed. This indicates, without doubt, that the stream is in a very

active state of degradation, but degradation is restricted by the resistant bed material. With the exception of limestone encountered at the upper end of the reach and sandstone at the lower end, all bedrock in the bed is shale.

Field inspection indicated the streambanks are in alluvium, with two exceptions. One is where the stream encroaches on the valley flank, exposing shale bedrock in the right bank (Section 3-5). The other exception is located in Section 11-12 where the channel is not encroaching on the valley flank. The streambank is in shale that extends 12 inches above the waterline and is capped with alluvium.

The limestone at the upper end of the reach (Sections 13-2 to 11-14) has significantly restricted the valley width compared to the shale upstream and downstream from this location. This restriction is verified by the existence of a limestone ledge outcrop across the channel. Another valley restriction occurs just below Longton near Section 11-1. There is no rock ledge outcrop in this area but limestone talus is identified on the right streambank. In addition, very coarse angular limestone fragments were described in an alternate bar about 1 mile downstream near Section 3-4. The fragments range in size to 36 inches in the longest dimension with the average being 6 to 8 inches. Downstream from the study reach, about 2 miles below Section 3-3 near Oak Valley, there is a third valley restriction, probably the result of sandstone (believed to be the Ireland member). This restriction, along with a significant addition of drainage area between Sections 11-1 and 3-7, needs to be recognized in any prediction about the study reach.

The stream valley form is dominated by the resistant geologic formations and the pattern cannot be predicted in terms of a consistently recurring sinuosity. There are two major patterns of sinuosity established by the channel--one with a wave length of about 2 miles, and the other with a wave length varying from 1,000 to 3,000 feet.

Hydraulics

The data submitted indicate that the stream transports 90 percent of the sediment as wash load. The project is expected to reduce the sediment discharge by about 60%, but the proportion of wash load to bed load is not expected to change much. The total sediment yield at the lower end of the study reach is estimated to be 31,000 tons per year before project installation and 13,000 tons per year after installation.

Storm runoff values are tabulated below for the bank-full channel condition. This is generally the maximum tractive stress condition and can indicate the time of maximum bed load movement. The existing bank-full discharge in the study reach is about 24 csm.

Table 1. Channel bankfull conditions.

Project Status	Runoff (in.)	Frequency Return Pd. (yr.)
w/o project	0.8	0.6
w/ project	1.6	1.8

Based upon this examination, it can be seen that the "maximum stress" on the bed will be less frequent. Instead of at least annually, it will be only about every other year.

Installation within the watershed of floodwater retarding dams with fixed releases will result in above-average flows for a prolonged period. Table 2 shows the approximate amount and duration of prolonged flow.

Table 2. Approximate amount and duration of prolonged flow.

Section	Drainage Area		Spillway Controlled		Flow Depth (ft)	Bank Full Discharge (cfs)
	Size (mi ²)	Controlled (%)	Release Rate (cfs)	Outflow Duration (hours)		
11-15	284.5	59	3,370	110	13	5-6,000
11-1	297.1	57	3,370	110	9	5-6,000
3-7	341.0	56	3,791	110	10	7-9,000
3-3	390.5	60	4,651	110	10	7-9,000

The prolonged flows will start 18 to 24 hours after the flood crest passes (based upon the synthetic 6-hour storm provided).

Twenty-three of the 45 project dams to be installed are designed to control the 25-year frequency flood event. These dams have overflow emergency spillways which will result in outflow greatly beyond the 20 csm release rate for storms with a magnitude greater than those with a 25-year return period. The runoff for a 6-hour storm of a 25-year return period is 4.5 inches. About 3.5 inches of this amount will be temporarily detained by each retarding structure. Nearly 60 percent of the drainage area above the study reach is controlled by floodwater retarding structures.

After the project is installed, out-of-bank flow will still occur as a result of the runoff from a 6-hour storm having a 25-year return period. This is based on the data supplied. The unit hydrograph

discharge is 14.74 csm per inch of runoff. There is about 4.5 inches of runoff for this size storm. Therefore, the stream discharge will be 14.74×4.5 , which is about 66 csm and exceeds the bankfull capacity of about 24 csm.

After the project is installed, the frequency of flooding due to out-of-bank flow can be determined by first dividing the 24 csm by 14.74 csm per inch to obtain a storm runoff value of 1.62. By use of the frequency-runoff curve provided, a storm of 6-hour duration and a 1.8-year return period will result in 1.6 inches of runoff.

The frequency of flooding due to out-of-bank flow will be reduced but still will be a common occurrence. However, the duration of flooding will be less. The long duration flows from the principal spillways are about one-half the channel capacity. The added flows from that portion of the watershed not controlled by dams will contribute the remainder of the runoff necessary to cause flooding. This portion of the runoff will be flashy or of short duration.

ADDITIONAL DATA REQUIREMENTS

The review and analysis of information provided indicates that the existing physical characteristics of the channel study reach are dominated by the geologic conditions and not the hydraulic or fluvial forces. The streambed is controlled by bedrock. However, information about the streambanks is very limited and general. To fully verify the streambank stability, it is desirable to collect additional data on the streambank material. These data would allow evaluation of both the structural stability of the streambanks against slides and their resistance to erosion from flow in the channel.

A review of old aerial photos and maps to provide information on the rate of stream alignment changes is advised. Photos taken before and after major storms would be especially valuable.

Examination of other watersheds in the area which have had similar project development would be very valuable in making evaluations. A detailed profile of the channel thalweg would be of value in contrast to the low flow water surface for hydraulic purposes.

It is necessary to examine upstream channel conditions for availability of bed materials to be transported into the reach under study. The channel performs and responds as a system in conjunction with the watershed, and a single reach cannot properly be evaluated without information about the other parts. In this case, the long-term performance of the channel bed materials will be greatly influenced by the bed materials upstream as well as the runoff.

This project is located in a subhumid region and, as a result, vegetation within the channel section will seasonally affect its flow characteristics. Knowing the woody and herbaceous climax species can provide some insight into the potential for an accelerated rate of channel choking, protection of the banks against erosion due to flow

impingement, and also the need for maintenance through the removal of snags and windfalls. Some tree species are significantly more susceptible to damage at an early age than others. Some have deeper or more dispersed root structure than others. Each of these items is valuable in anticipating the effects of vegetation.

EVALUATION

On the basis of available data, the channel study reach has been evaluated to anticipate the project's short- and long-term alterations to

1. Meander pattern,
2. Configuration of the channel,
3. Substrate material, and
4. Pool-riffle sequence.

The meander pattern is very stable. The streambanks are covered with mature vegetation, the channel is very narrow and uniform, and there are no extensive point bars. The hydraulic stress of bankfull flows has been very frequent. On the basis of past demonstrated stability, a reduced frequency of stress will likely result in an even more stable meander pattern in the future, or essentially no change.

The channel cross section shape or configuration also appears to be very stable. In many areas it is made up of a compound slope on each bank. There is a flattened slope in the bottom 2 to 4 feet, and above it is the steepest slope, about 1 to 1, for a height of 6 to 15 feet. Above this steep section there is usually a flattened bench-like slope, topped off with another steep section. The base of the streambank is a gravel layer resting on bedrock of shale, limestone, or sandstone. The generally benched slope and solid foundation provide a generally stable section, as demonstrated by the presence of mature vegetation on the bank.

The Elk River location in a subhumid area is adequate for bank vegetation. The average velocity of streamflow at bankfull capacity ranges from 2.5 to 4 feet per second throughout the reach. The maximum velocity at the point of extreme nonuniformity on curves is greater, but probably no more than 8 feet per second. These relatively low velocities will not inhibit vegetative growth from maturing in the future. The vegetation will provide increased bank protection against hydraulic erosion. The vegetated banks will tend to reduce the velocity at the boundary layer causing some sediment deposition. Channel cross sectional areas are not very likely to diminish as the stream normally carries little sediment and the project should reduce the sediment load in the future by one half.

The substrate material consists of either exposed bedrock or coarse gravel. The change in gradation is not anticipated to be significant.

Because the streambanks in the study reach will be more stable, the dominant source of bed materials will increasingly be the transient bed load moving out of upstream reaches. There should be no long-term change in bed material characteristics due to the existing volume of material in the "pipeline." The rate of delivery of bed material will be reduced to about one-half the present rate and, therefore, the volume of bed material should be ample for a very long time since the maximum hydraulic stress at bankfull flow still will move the bed materials but at a less frequent rate.

The bed materials are expected to become generally coarser through removal of some of the sand sizes, because the reservoir release will increase the duration of one-half bankfull flows. This should result in a veneer type of armoring on the bed after each flood event. The veneer size can be evaluated for stability using the Shields diagram. The gravelly bed material is not expected to become choked with silts and clays, since there will be a cleansing at least biannually (on the average) when bankfull flow disturbs the bed surface during its transport.

There will be no significant deepening of the channel since it rests on bedrock of moderate resistance. Bed materials form the pool riffle sequence, but the underlying bedrock restricts the pool depth. In at least one location the riffle is created by a bedrock ledge outcrop crossing the stream.

CONCLUSIONS

The channel study reach of the Elk River demonstrates itself as being very stable, and its characteristics are controlled by geologic factors rather than hydraulic forces. The potential for degradation of the channel bed is limited because of bedrock that is exposed or at shallow depth.

The installation of floodwater retarding dams for controlling runoff from 60% of the watershed will generally result in an even more stable channel. A slight coarsening of the bed material will occur.

The potential for channel choking or obstruction of flow will increase because of increased aging of trees and other vegetation. In addition, growth will provide increased protection against erosion at locations that are lower on the banks and more vulnerable to flow impingement. Both of these effects will result in increased stability of channel alignment.

Most of the answers on anticipated changes would be provided by an examination of the study reach data along with field verification by a person knowledgeable of stream channel behavior and experience with similar streams in the humid or subhumid area. This, along with a simple analysis using the Shields diagram, should provide a good estimate of the substrate size likely to result.

ASSESSMENT OF ANTICIPATED CHANNEL CHANGES IN THE ELK RIVER NEAR LONGTON, KANSAS DUE TO UPSTREAM DEVELOPMENTS

Michael A. Stevens
Consultant, Boulder, Colorado, U.S.A.

INTRODUCTION

The Elk River is a stream draining the east flank of the Flint Hills Escarpment primarily in Elk County in southeastern Kansas (Figure 1). The reach of interest here is at Longton, Kansas which is approximately 70 miles east-southeast of Wichita. At the downstream end of the reach, the drainage area is 405.3 mi².

The floodplain of the Elk River is the prime agricultural land in the area. Beef production is the main agricultural activity. The native pasture, which covers over 76% of the watershed, is utilized as grazing land. Feed grain and alfalfa are produced on the floodplain. Almost 80% of the floodplain is in crops. The 15,375 ac of floodplain cropland represent approximately 45% of the cropland in the watershed so farmers try to keep these lands in production despite frequent damaging floods. On the average, two to three flows a year exceed bankfull capacities. Flooding usually occurs during the growing season.

The U.S. Department of Agriculture, Soil Conservation Service has designed a system of 48 floodwater retarding structures to be installed in the Elk River catchment. The function of these structures is to store floods in the headwaters so as to mitigate flood damage to the cropland on the floodplain downstream.

The system of earth dams will provide 50,253 ac ft of floodwater detention storage and 12,942 ac ft of sediment storage. The system of structures will control the runoff from 239.8 mi² which is 59% of the watershed area.

Floodwater storage will be provided for from 3.0 to 5.25 in. of runoff from the upstream drainage area. Also, storage will be provided for the 10-year accumulation of sediment, this volume being equivalent to 0.51 to 1.79 in. of sediment yield from the upstream drainage area. An ungated orifice will be placed at the elevation of the 50-year accumulation of sediment.

The principal spillway of each structure will be reinforced concrete or comparable quality material with a single stage inlet. The uncontrolled design release rate is 20 ft³/s per mi² of drainage per in. of runoff.



Figure 1. Elk River Watershed upstream from Longton, Kansas

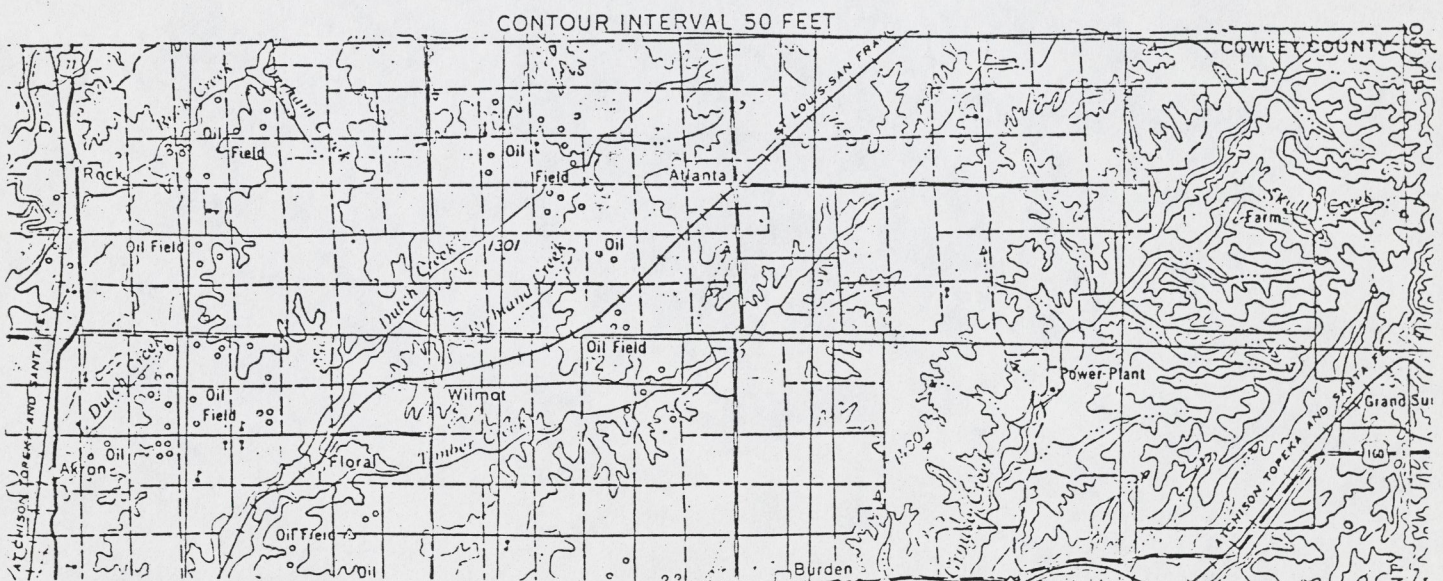
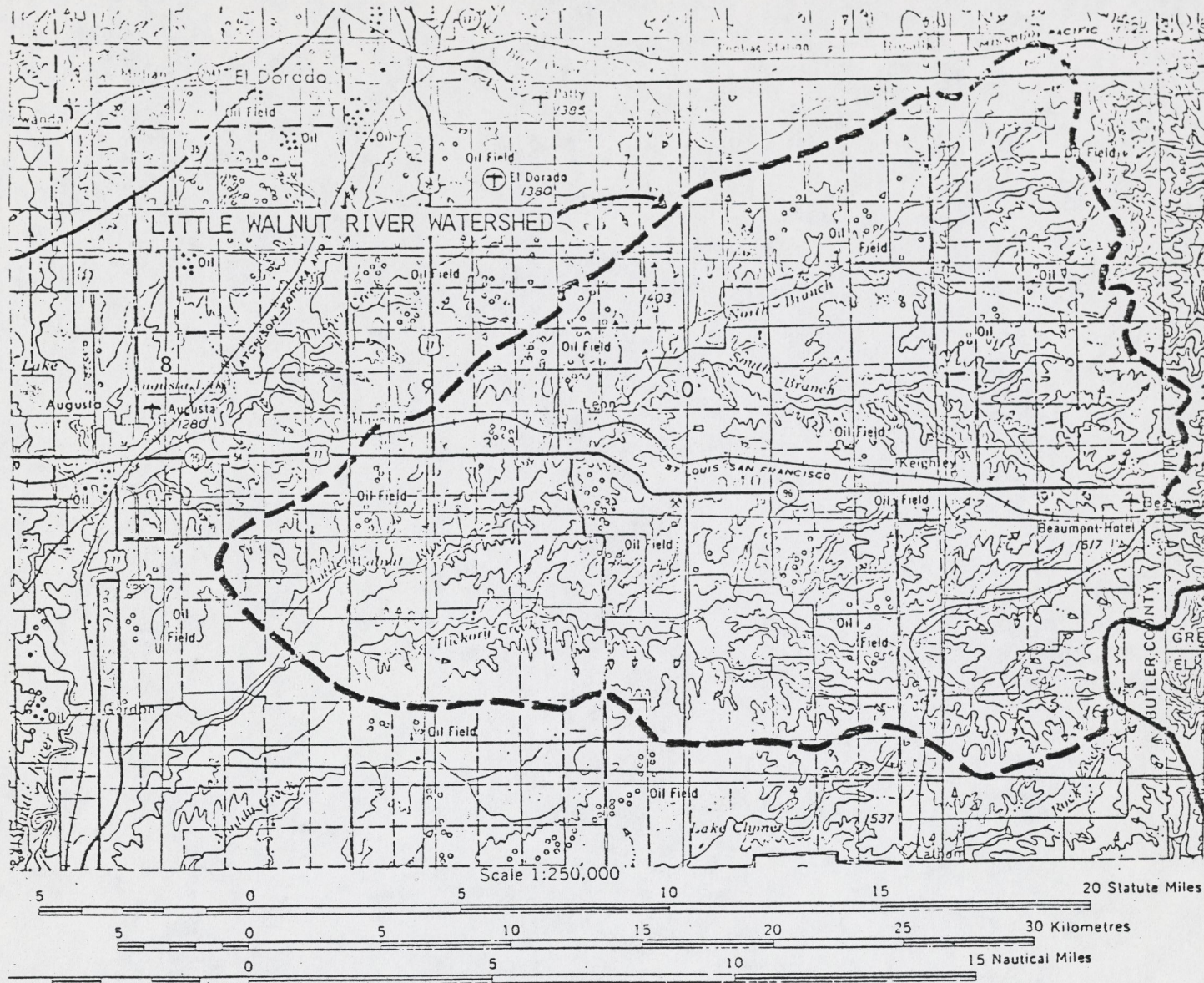


Figure 1. Elk River Watershed upstream from Longton, Kansas (Continued).

Each dam has a vegetated or rock emergency spillway to discharge runoff exceeding reservoir storage capacity. The chance of the spillway operating in any one year is 4% or less.

The question addressed herein is, "What will happen to the morphology of the Elk River as a result of changes in streamflow and sediment discharge caused by the construction of the 48 small flood retention reservoirs upstream?" Specifically, both the short and long term changes in meander pattern, channel configuration, substrate material, and pool riffle sequence are desired.

GENERAL DESCRIPTION

Climate

The climate of southeastern Kansas is sub-humid. The average annual participation is 35.07 in. The largest annual precipitation recorded at Howard, Kansas was 56.07 in. in 1961; the smallest was 18.47 in. in 1956. Normally, approximately 75 percent of the precipitation falls during the growing season, April to October (U.S. Department of Agriculture 1967).

The average growing season is 185 days, and in a normal year the area is frost free from 15 April to 17 October.

Temperatures generally average 35°F during January and 80°F during July. Extreme temperatures have been above 115°F and below -20°F.

Valley

The valley of the Elk River in the vicinity of Longton, Kansas lies in an east-west direction and is relatively straight (Figure 2). The valley floor is approximately 3500 ft wide on the average, and is some 100 to 150 ft below the level of the surrounding hills.

The north side of the valley floor from Longton to 3 mi west of Longton is terraces (Verville et al., 1958). The stream-laid deposits of gravel, sand, silt, and clay are as much as 40 ft thick. The coarse materials, predominantly chert, limestone, and sandstone gravel are commonly found in the lower zone ranging from a fraction of an inch to 8 ft in thickness. Sand is intermingled with the pebbles, some of which are 2 to 3 in. in diameter. The upper part of the deposit consists mostly of clay and silt but grades downward to more sandy material.

At Longton, the valley slope changes from 6.5 ft/mi upstream to 5.0 ft/mi downstream. The reason could be rock outcrops on the valley floor. Downstream from Longton the valley is wider and there are no terraces, at least to Elk City.

River Morphology

Shown on Figure 2, the study reach is that 50,000 ft long section of the Elk River between River Station 96+000 and 146+000. The

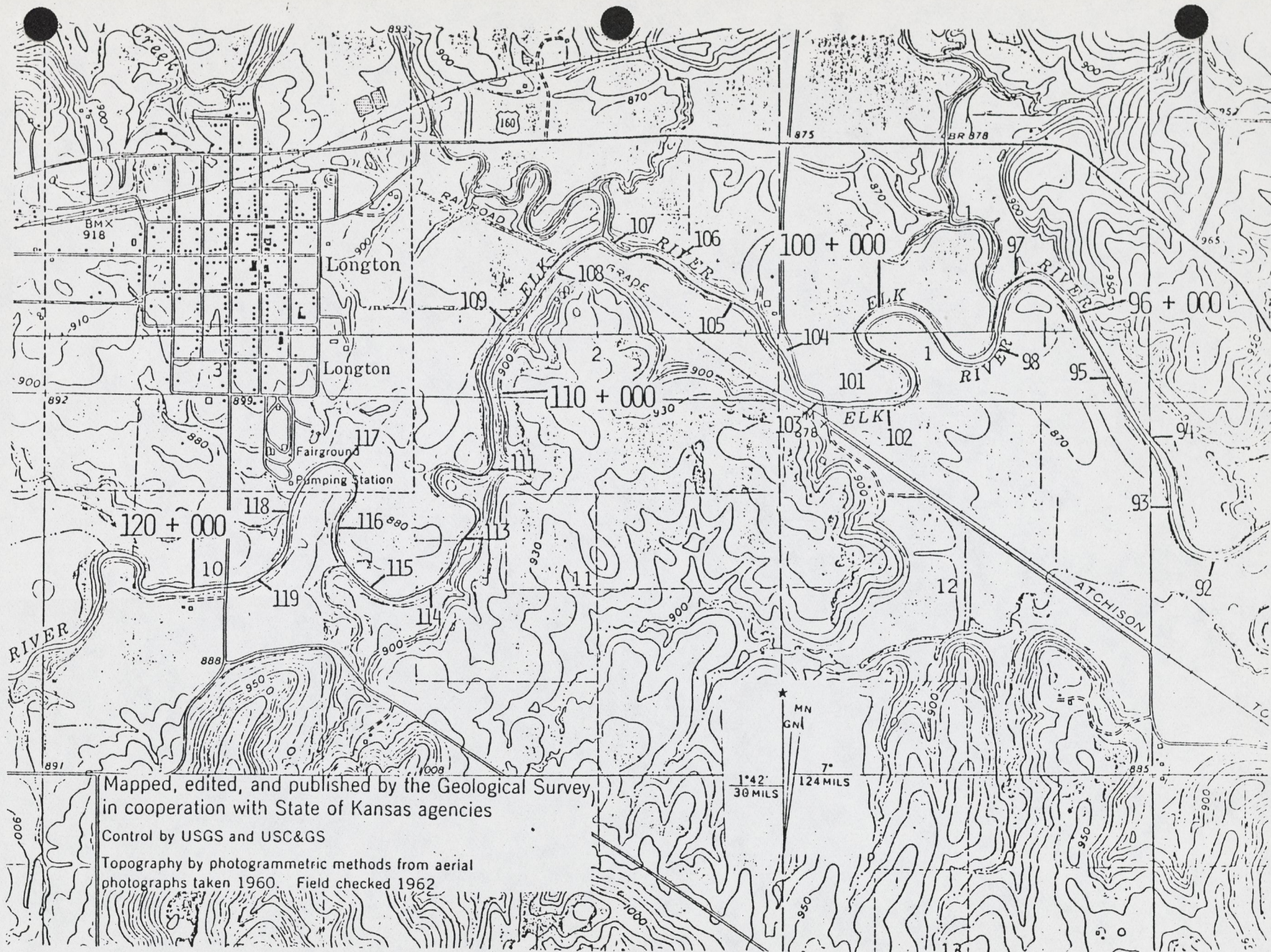


Figure 2. Elk River near Longton, Kansas

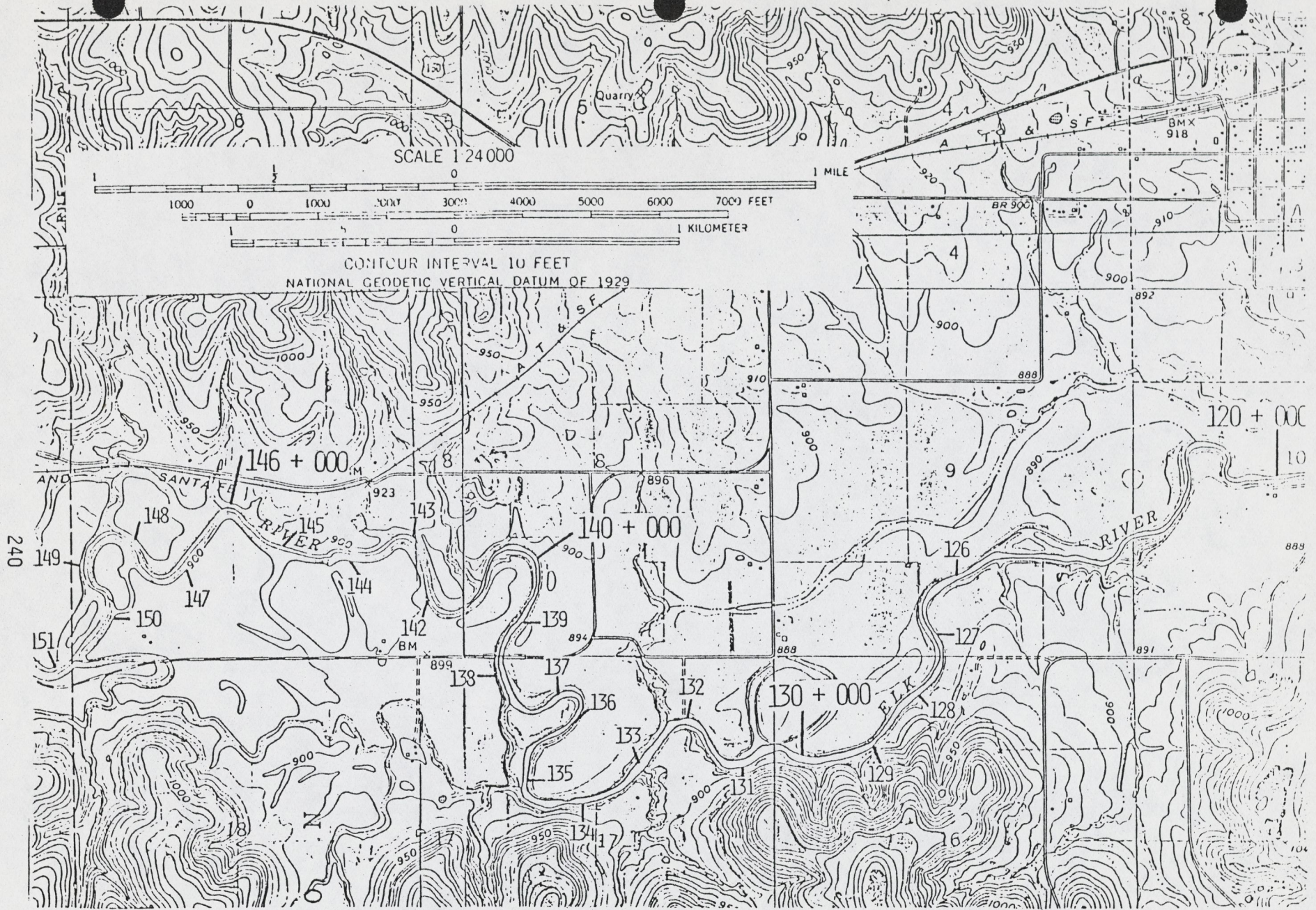


Figure 2. Elk River near Longton, Kansas (Continued).

stationing is in feet. River Station 0+000 was chosen as the U.S. gaging station downstream from Elk City. Below this gage the river cuts through the hills and enters a reservoir.

The slope of the river channel in the study reach changes from approximately 1.5 ft/mi at the downstream end and to approximately 3.5 ft/mi at the upstream end. The channel profile has a discontinuity upstream at the Elk Falls where the river drops at least 10 ft in a 1000 ft long reach.

At the upstream end of the study reach the river is against the north valley wall. Thereafter, the river crosses the valley floor to wander along the south valley wall. Downstream from Longton, the river meanders abruptly to the north side of the valley. In the 5.3 mi length of the valley, the river channel is against the valley wall for approximately 2.5 mi.

In this reach, the sinuosity of the Elk River is approximately 1.8 overall. One cutoff has occurred recently leaving a timbered loop on the floodplain at River Station 130+500.

The bankfull width of the river channel averages 200 ft with variations of 50 ft in either direction. The banks are timbered throughout the entire reach.

The bankfull depth varies from approximately 25 to 50 ft; the average is 33 ft. Thus, the width-to-depth ratio for bankfull flow in the Elk River is approximately 6.0 in the study reach.

Based on the plan, profile, and cross-sectional surveys conducted for the Elk River Watershed, the bankfull discharge for the reach is in the range from 5000 ft³/s to 10,800 ft³/s; the average is 7800 ft³/s.

Streamflow

Streamflow records have been kept for the U.S. Geological Survey gaging station on the Elk River at Elk Falls from January 1967 to the current year. The drainage area for the station is 220 mi².

In the 11 years of complete annual record, the average discharge has been 181 ft³/s which corresponds to an average runoff of 11.18 in./year from the entire catchment upstream.

Monthly records of streamflow have been compiled for the 1968 to 1976 water years inclusive. These are summarized in Table 1.

Table 1. Mean Monthly Streamflow Elk River
at Elk Falls (Units of ft³/s and %)

Oct	Nov	Dec	Jan	Feb	Mar	Apr	May	Jun	Jul	Aug	Sep	Annual
153	233	139	141	135	336	284	250	236	292	24.4	56.7	190
6.83	10.05	6.21	6.28	5.50	14.99	12.26	11.16	10.16	13.02	1.09	2.45	100

In a normal year, streamflow is a maximum in March when 15% of the annual runoff occurs on the average. August is the driest month with only 1% of the annual runoff.

A listing of all floods greater than 4000 ft³/s at Elk Falls is given in Table 2. Based on the 11 years of annual peaks, the mean annual flood is 32,800 ft³/s; the coefficient of variation is a rather large value of 1.7.

The flood of record occurred on 3 July 1976 when the momentary discharge reached an estimated 200,000 ft³/s, a value 6 times greater than the mean annual flood and 40 times greater than the minimum bankfull discharge.

In the 11-years of record, the flood peak at Elk Falls has exceeded the lowest bankfull discharge (5000 ft³/s) in the study reach 40 times. Eighteen of these 40 floods occurred in the growing season.

Sediment Yield

There has been only one suspended sediment sample taken in the Elk River in the period of record. Sediment yield has been estimated by the Soil Conservation Service in order to design the detention reservoirs. Their estimates of sediment yield range from 0.51 in. to 1.79 in. equivalent erosion in 100 years.

The development plan is to store 12,942 ac ft of sediment from a drainage area of 239.8 mi² in 100 years. If it is assumed that all sediment is stored, the equivalent sediment yield is approximately 0.01 in./year. This value corresponds to an average annual suspended sediment concentration of approximately 900 mg/l.

By way of comparison, the rate of sedimentation in Howard Lake near Howard, Kansas (Figure 1) since 1936 is equivalent to an erosion rate of 0.007 in./year.

The Soil Conservation Service (1967) reports that upland erosion is a serious problem. Floods occurring in the springtime after the thaw but before vegetative cover develops cause extreme land damage. Furthermore, the sediment deposited on the floodplain is infertile silt,

Table 2. Floods in the Elk River at Elk Falls.

Water Year	Date	Hour	Discharge ft ³ /s
1967	5 Jul		7,240
1968	7 Oct	1900	5,850
	3 Apr	1300	4,150
	25 May	1045	9,610*
1969	9 Oct	1300	5,480
	2 Nov	2400	7,900
	24 Mar	0330	5,480
	18 Apr	0515	10,100
	27 Apr	0500	10,100
	8 May	1400	4,720
	30 May	1300	8,940
	1 Jun	1300	19,100*
	24 Jun	0600	4,860
16 Sep	0800	5,160	
1970	12 Oct	2100	9,460
	1 Apr	1715	12,400
	18 Apr	2015	29,300*
	12 Jun	0600	6,380
1971	3 Jan	1645	4,580*
1972	15 Dec	1045	6,220
	18 Jul	2200	15,200*
1973	13 Nov	0915	7,050
	30 Dec	0915	8,430
	1 Feb	1330	4,460
	4 Mar	1630	7,110
	6 Mar	2000	6,170
	9 Mar	0415	9,620
	10 Mar	2130	10,700*
	25 Mar	1045	6,420
	31 Mar	0700	5,790
	15 Apr	2100	4,230
1974	11 Oct	1630	5,100
	20 Nov	1230	5,580
	4 Dec	1115	8,920
	9 Mar	0400	4,400
	10 Mar	0745	14,500*
	30 Apr	1100	4,290
	22 May	0915	6,460
	24 May	0015	5,160

Table 2. (Concluded).

Water Year	Date	Hour	Discharge ft ³ /s
1975	31 Oct	0400	12,900
	3 Nov	1000	22,400*
	31 Jan	0100	6,830
	27 Mar	1745	6,570
1976	28 Apr	1500	8,120
	3 Jul	0700	200,000*
1977	13 Apr	2130	6,420
	21 May	1345	9,100
	22 Jun	1130	24,100*
1978	24 Mar	1230	4,370
	19 May	1415	11,700*

^aThe symbol "*" denotes the peak discharge for the year.

sandy silts, and clays. From the information given by the Soil Conservation Service (1967), Verville et al. (1958), and Bell and Rowland (undated), it is concluded that the sediment yield of the Elk River catchment must be mostly silt and clay.

Bed Material

The material on the bed of the Elk River in the study reach was investigated in February 1980. The data collected indicate that the riverbed level is controlled by rock outcrops and that there is but a thin vein of coarse alluvium on the bed at other locations.

Immediately above the upstream end of the reach (RS 152+000) the river is flowing on limestone. Between here and the abandoned railroad grade (RS 108+000) the riverbed consists of Lawrence Shale overlain with alluvium and an armor layer of limestone, sandstone, and shale material in the gravel, cobble, and boulder sizes. The particles are very angular indicating they have not travelled far from their parent rocks. In the areas sampled, the depth of alluvium varied from zero (RS 138+000) to 4 ft (RS 124+000).

Downstream from the abandoned railroad grade, the river flows on a relatively thin layer of alluvium on top of the Ireland Sandstone member of the Lawrence Shale formation. At River Station 103+000, sandstone was encountered 6 in. below the bottom of the channel. At other locations an armor layer of gravel and boulder sandstone particles covered the bed.

It is apparent that the bed of the Elk River is not, in general, alluvial. The bed level is controlled by outcrops of rock or gravel and boulder particles of angular rock obtained from the outcrops but not transported far from their source. It is inferred that the bedload of this river is very small.

Bank Material

The river banks are alluvial at almost all cross sections investigated in February 1980. One exception is at River Station 103+000 where the river is against the south valley wall. Here the right bank is limestone talus of gravel and cobble sizes.

The upper (above the low-water channel) slopes of the river banks are approximately 2.5 horizontal to 1 vertical and support a growth of old timber, and grass. In locations where the upper slopes are steeper but not yet caving, only grass and willows grow. It appears that the alluvial banks are composed almost entirely of silt and clay materials.

Bank erosion was observed in February 1980 at two locations on the outside of bends. Here the caving banks were vertical and bare of vegetation.

The river banks are timbered all along the reach and, at many cross sections, there appear to be natural levees on the floodplain side, in some cases, at least 3 feet high. There are no man-made levees.

Floodplain

The valley floor of the Elk River in the study reach is on the average 3500 ft wide. The land which is floodplain is cropped with alfalfa and feed grains. There are roads and small drainage channels on the floodplain. The cross-sectional surveys indicate the surface of the floodplain is very irregular.

EFFECTS OF DEVELOPMENT

The 48 flood retention reservoirs in the upstream catchment will affect the amount and delivery of water and sediment to the Elk River. Estimates of these have been made by the Soil Conservation Service.

Water Yield

The annual water yield from the Elk River catchment will be only slightly decreased because of evaporation from the water ponded in the flood retention reservoirs.

Flood Hydrographs

Floods will be decreased greatly when the 48 flood retention reservoirs are constructed. The peak of the unit hydrograph will be reduced from 30.3 to 14.7 ft³/s per mi² per in. of runoff. That is a reduction of approximately 50 percent. Also, the volume of water in the main part of the hydrograph (first 36 hours) will be reduced approximately the same amount.

Sediment Discharge

The sediment trapped in the flood retention reservoirs will result in an estimated 58% reduction in the annual sediment load in the Elk River at Longton. As the bed-load transport is very small now, the reduction will be primarily in the suspended load which is presumed to be mostly silt and clay.

RESPONSE OF THE RIVER

Methods

The response of the Elk River to the great decreases in flood peaks and sediment discharge can be predicted on the basis of general geomorphic relations developed from experiences in many rivers in many parts of the world. Normally, I would supplement the geomorphic analysis with some mathematical modeling of the water and sediment transport, and changes in bed-material size and configuration. In this

case, the mathematical modeling is foregone for these reasons: 1) the river bed is not entirely alluvial which means that the bed load at a section may not be related to the local shear stress but controlled entirely by the amount coming in from above; 2) much of the material on the bed is very angular gravel, cobbles, and boulders which have not been transported far from their parent rock. Existing bed-load equations are for more rounded rock and not this talus material; 3) the river meanders appreciably, so average flow properties at any cross section may not represent values to be used in existing gravel transport equations (which were developed for more regular reaches of rivers); 4) it appears that the amount of bed load in the river is very small and of little consequence; and 5) there are no field data with which to calibrate a mathematical model of the sediment transport process.

The geomorphic relations employed in this paper are those developed by Schumm (1977) but modified slightly based on experience on tropical islands in the Orient where sediment concentrations reach values as large as 100,000 mg/l in the rivers and 20,000 mg/l in the irrigation canals.

Geomorphic Equations

The basic premise is that the river morphology is in "regime." That is, the width, depth, and other features have adjusted over a long period of time to conform to the stresses caused by the water and sediment load. It appears that the Elk River is in regime because there are no reports of geomorphic change caused by the enormous flood of 3 July 1976. Also, if a flood with a peak 6 times the mean annual flood could not cause widespread bank caving and channel change, one must presume the banks and bed are stable and will not respond quickly to changes in water and sediment discharge.

Schumm's expression relating river channel morphology to water discharge is

$$Q \sim \frac{b, d, \lambda}{s} \quad (1)$$

in which, Q = either the mean annual discharge, or the mean annual flood

b = bankfull width

d = bankfull depth

λ = meander wave length

S = riverbed slope

Equation 1 is not an equality but merely a short-hand method of saying that the magnitudes of the width, depth, and wave length are

directly proportional to the magnitude of the streamflow and the bed-slope is inversely proportional to the magnitude of the streamflow.

In relating channel morphology to sediment load, Schumm assumed that the percent of silt and clay in the wetted perimeter is inversely proportional to the bed load and that the total load is directly proportional to the bed load. In general, these assumptions may be valid. However, it would seem that one can expand Schumm's analysis to reflect the fact that some rivers have noncohesive beds but transport mostly silt and clay. These rivers can respond differently depending on whether it is the silt and clay load or the noncohesive load which is affected by development.

Schumm has shown that channel width and depth are related closely to the percentage of silt and clay (M) in the sediments forming the perimeter of the channel.

For the range of channels studied by Schumm

$$\frac{b}{d} = \frac{225}{M^{1.08}} \quad (2)$$

$$b = 37 \frac{Q^{0.38}}{M^{0.39}} \quad (3)$$

and $d = 0.6 M^{0.34} Q^{0.29} \quad (4)$

Here Q is the mean annual discharge in ft³/s and b and d have units of ft.

In general, it is expected that the values of the coefficients and exponents in the above equations vary somewhat with geological setting and size of river.

Equations 2, 3, and 4 indicate that, for these particular channels, the percent silt and clay was inversely proportional to the river size or

$$M = \frac{57}{Q^{0.26}} \quad (5)$$

One can argue that it is the amount of silt and clay in the banks only that determines the width of the alluvial channel and that the composition of the bed is less important in determining the width. That is, one can have the same channel width and shape with a sand bed or

gravel bed provided the banks are clay. It is the cohesion in the clay that allows clay-bank channels to withstand higher stresses developed in narrow channels. Thus, one can use the wash load Q_w (silt and clay transported in suspension) as the indicator of the effect of sediment on channel width. Furthermore, M is directly proportional to Q_w so Equation 3 can be expressed as

$$b \sim \frac{Q}{Q_w} \quad (6)$$

The same type of argument applies to depth. If the wash load is large (large concentration of silts and clays), the river channel is narrow, flood depths are large, a large amount of sediment can be deposited on the flood plain, thus building up high banks. Therefore Equation 4 can be expressed as

$$d \sim Q_w, Q \quad (7)$$

It follows that the bankfull width-to-depth ratio b/d is almost entirely dependent on the wash load.

Following this type of reasoning and by considering Q constant, an expression relating channel morphology to wash load (silt and clay carried as suspended load) can be derived

$$Q_w \sim \frac{d, P}{b, \lambda, S} \quad (8)$$

in which P = sinuosity of the channel

λ = meander wavelength

S = riverbed slope

It has been assumed that the valley slope is an independent variable, the valley having been carved by hydrologic events no longer directly influencing the river channel shape.

Equation 8 is equivalent to Schumm's if one assumes Q_w is inversely proportional to the bed-material load Q_b . In this expression, the bed slope change is accomplished by changes in sinuosity and meander wavelength and not by aggradation and degradation.

Now, it is known that for rivers with alluvial beds and fixed cross-sectional shape,

$$Q_b \sim \frac{Q, S}{d_{50}} \quad (9)$$

This is Lanes' (1955) qualitative relation between bed-material load Q_b , water discharge Q , median bed sediment d_{50} , and riverbed slope S . The expression can also be derived mathematically by relating shear stress on the bed with the transport of non-cohesive bed particles. The change in slope in Equation 9 is accomplished by aggradation or degradation and not by a change in alignment as in Equation 8.

For a fixed discharge, then

$$Q_b \sim \frac{S}{d_{50}} \quad (10)$$

Now, Equations 1, 8, and 10 form a set of relations among water and sediment discharge and alluvial river morphology. The assumptions are that the wash load and water discharge have the major influence on the cross-sectional shape, slope, and sinuosity of the channel, and the bed-material load relates closely to only the material size d_{50} and the slope S .

For the relatively straight forward cases of an increase or decrease in discharge, washload, or bed-material load alone, the response of a channel to change is

$$Q^+ \sim b^+, d^+, \lambda^+, S^- \quad (11)$$

$$Q^- \sim b^-, d^-, \lambda^-, S^+ \quad (12)$$

$$Q_w^+ \sim b^-, d^+, \lambda^-, S^-, P^+ \quad (13)$$

$$Q_w^- \sim b^+, d^-, \lambda^+, S^+, P^- \quad (14)$$

$$Q_b^+ \sim S^+, d_{50}^- \quad (15)$$

$$Q_b^- \sim S^-, d_{50}^+ \quad (16)$$

Here, a plus or minus exponent is used to indicate how, with an increase or decrease of water or sediment discharge, the various aspects of channel morphology change. The plus exponent indicates an increase and a negative exponent indicates a decrease. No change is denoted with a zero exponent.

One physical interpretation of the set of expressions given above is this. The water discharge represents processes which tend to erode the banks and bed and straighten the channel alignment. The wash load represents processes which tend to build and maintain banks and to contort the alignment. The bed-material load represents processes which tend to change mainly the level and configuration of the bed and the size of material on the bed. The physical interpretation is important because the expressions may not represent all possible sequences.

Immediate Response

The immediate effect of decreasing the flood discharges and sediment load in the Elk River system should be as follows.

1. The bed material will be coarser as the supply of bed-material load from upstream is decreased. There will be less fines on the river bed, however, little or no decrease in riverbed level in the study reach is anticipated as the armor coat is non-alluvial, angular, and very coarse in most places. Rock outcrops in the bed at other places. The riverbed slope will not decrease due to degradation. Therefore

$$Q_b \sim \frac{S^0}{d_{50}^+} \quad (17)$$

It is assumed that the tributary channels on which the dams will be constructed have essentially coarse gravel, non-alluvial beds. If the beds of some tributaries are composed of sand, these channels could supply an excess bed load to the main stream during the first few years of operation.

2. The decrease in flood discharge will decrease the processes tending to widen the channel by erosion and undercutting the banks. As the bed level will not change appreciably and

sediment deposition on the floodplain will decrease, the bank full depth should not change immediately. The meander wave length will not change rapidly as considerable time (a century or more) is required for the river to work laterally across the floodplain destroying one pattern and building another if the existing floodplain deposits are tough cohesive materials. Then it follows that the channel should narrow if the sediment required to build banks is available. That is

$$Q^- \sim \frac{b^-, d^0, \lambda^0}{S^0} \quad (18)$$

A new and more vigorous growth of vegetation may result on the banks as plant scouring will be reduced because flow velocities will be lower.

3. There will be a considerable decrease in the amount of wash load carried by the river. The turbidity of the water will be much less. Wash load is the material from which new banks are made. New deposits on the banks could be facilitated by more luxuriant vegetation on the banks. It follows then, from the geomorphic expression relating wash load and channel shape, that the channel should widen initially. That is

$$Q_w^- \sim \frac{d^0, p^0}{b^+, \lambda^0, S^0} \quad (19)$$

This is in contradiction with the conclusion drawn from the geomorphic expression relating discharge and channel shape. However, the influence of Q_w on the width-to-depth ratio is much more pronounced than the influence of the water discharge. Therefore, the channel should tend to widen.

It is concluded that the short-term response will be a channel which is deeper than required and which is slightly too narrow now. The channel will not degrade but the bed will become cleaner as the finer material is removed leaving even more armor coat than is currently on the bed. The pool and ripple sequence will remain unchanged.

Long-term Response

If left unimpeded over a very long period of time, the Elk River would move its channel laterally across the floodplain eroding one bank and building the other. Even though the farmers would not tolerate the

river destroying their cropland by moving laterally, it is useful to estimate the change in river form which would result from migration. Because of the changes in water and sediment discharges, the new river will have a different form than the current river. The differences can be estimated as follows.

1. The long-term response to the decrease in bed-material load will be the same as the short-term response. That is, the river will not be able to degrade because the bed material is very coarse and there are rock outcrops controlling the bed level. Therefore, the slope will not change due to degradation, but the bed material will become coarser. That is, the new regime will be

$$Q_b^- \sim \frac{S^0}{d_{50}^+} \quad (20)$$

2. The long-term response to the decrease in water discharge and wash load can be determined by noting that the combined response is

$$Q^- Q_w^- \sim \frac{b^+, d^+, P^+, \lambda^+}{S^+} \quad (21)$$

That is, the new river will be shallower (flood discharges are not as large so the floodplain will not build up as high as before), less sinuous, and will have larger bed slope (due to less meandering, not aggradation). There should be fewer pools and ripples. There are questions concerning whether the channel will become wider (b^+) or narrower (b^-) and whether the meander wave length will be longer (λ^+) or shorter (λ^-).

First the width. It is my opinion that the vegetation on the banks will be the added stabilizing factor which will cause the new river to be narrower. Vegetation is effective in slowing down the velocity and trapping fine sediment on the bank.

The meander wave length must increase because the sinuosity will decrease and the slope will increase (due to strengthening, not aggradation).

In order to predict the magnitude of the anticipated long-term changes in river form, the geomorphic equations developed by Schumm

(1977) are used here. These equations were developed from rivers with width-to-depth ratios from 2 to 300 so represent a wide range of river forms. The equations are

$$\frac{b_f}{b_o} = \frac{Q_f}{Q_o}^{0.38} \frac{Q_{wo}}{Q_{wf}}^{0.39} \quad (22)$$

$$\frac{d_f}{d_o} = \frac{Q_{wf}}{Q_{wo}}^{0.34} \frac{Q_f}{Q_o}^{0.29} \quad (23)$$

$$\frac{P_f}{P_o} = \frac{Q_{wf}}{Q_{wo}}^{0.25} \quad (24)$$

and

$$\frac{\lambda_f}{\lambda_o} = \frac{Q_f}{Q_o}^{0.48} \frac{Q_{wo}}{Q_{wf}}^{0.74} \quad (25)$$

Here the subscripts "o" and "f" refers to conditions before and after development, respectively.

Because the valley slope remains changed, it follows that

$$\frac{S_f}{S_o} = \frac{P_o}{P_f} \quad (26)$$

Using the values,

$$b_o = 200 \text{ ft}$$

$$d_o = 33 \text{ ft}$$

$$P_o = 1.8$$

$$\frac{Q_f}{Q_o} = 0.50$$

$$\frac{Q_{wf}}{Q_{wo}} = 0.58$$

the following estimates of the changed river form are obtained.

$$b_f = 190 \text{ ft} \quad (5\% \text{ smaller})$$

$$d_o = 22 \text{ ft} \quad (32\% \text{ smaller})$$

$$P_f = 1.6 \quad (13\% \text{ smaller})$$

Currently, there are 17 meanders with a wide range of wave lengths in the study reach. As wave length is inversely proportional to number of meanders, one anticipates there will be one less meander after development. Similarly there should be a reduction of approximately 10 percent in the number of pools and ripples.

The rate at which the Elk River could move laterally across its valley will be much slower after development. Since there is very little evidence of bank caving now, it is anticipated that it will be decades before one will notice any response. The geological survey indicates that the alluvium on the valley floor is primarily silt and clay but if the river was to encounter non-cohesive materials in the former alluvial deposits, response would be much faster. If any local area becomes unstable, for example, if the river cuts into a sandy deposit, this area will probably be stabilized immediately at the request of the local land owners.

Other Considerations

Changes in river form due to changes in wash load and vegetal factors are not so well documented in geomorphic literature as the response due to changes in discharge and bed-material load. Therefore, some caution is warranted. Two studies should be done to improve estimates made above. They are to

1. Develop geomorphic equations based on the data of rivers in southeastern Kansas, and
2. Study the effect 93 flood retention reservoirs in the Little Walnut River catchment have had on the Walnut River. This catchment is immediately west of the Elk River catchment (Figure 1). Also, the vegetation along the river banks must be maintained. There is a question concerning what will happen when the large trees die of old age. A study of the vegetation succession for this river is warranted.

SUMMARY

The anticipated response of the Elk River near Longton to the construction of 48 flood retention reservoirs in the upstream catchment is as follows.

1. In the short term, the material on the riverbed will become coarser.
2. If left unimpeded in the long-term, the river would become slightly narrower, much shallower, slightly less sinuous, slightly steeper (but with no aggradation) and the number of ripples and pools would decrease slightly. However, it is anticipated that the land owners will not tolerate lateral migration of the river. Therefore, in the long term, the Elk River will become slightly narrower and will be left with its large remnant bankfull depth developed during a period before dams when water and sediment discharges were much greater.
3. Very little effort will be required to keep the Elk River in its stable regime even after development...if the vegetation on the banks can be maintained. More effort should be given to vegetation studies along this river and land owners should be informed of the importance of river bank vegetation.

REFERENCES

- U.S. Department of Agriculture, Soil Conservation Service. 1967. General Plan, Elk River Watershed, Joint District No. 47; Elk, Greenwood, Butler, Chautauqua and Wilson Counties, Kansas. Lincoln, Neb., January.
- Verville, G. J. et al. 1958. Geology, Mineral Resources, and Groundwater Resources of Elk County, Kansas. State Geological Survey of Kansas, Vol. 14, University of Kansas, Lawrence, Kansas, July.
- Bell, E. L., and H. T. Rowland. Soil Survey of Chautauqua County, Kansas. U.S. Department of Agriculture, Soil Conservation Service.
- Schumm, S. A., 1977. The Fluvial System. John Wiley & Sons, New York, 338 p.
- Lane, E. W., 1955. The Importance of Fluvial Morphology in Hydraulic Engineering. ASCE Proceedings, Vol. 81, No. 745, 17 p.

APPENDIX A

CHANNEL CHANGE WORKSHOP: PROBLEM NUMBER 1
YAMPA RIVER, COLORADO

THE QUESTION

The question relative to the Yampa River is "What will happen to the morphology of the stream channel as a result of the changes in stream flows and sediment discharge caused by the construction of a number of reservoirs upstream of a reach of stream?" Specifically, for both a short and a long time after construction of the reservoirs:

1. What will be the meander pattern?
2. What will be the configuration of the channel?
3. What will be the substrate material?
4. What will be the pool riffle sequence?

The Yampa River is located in Northwestern Colorado. All the reaches of interest are in the lower part of the river. Data are supplied for four reaches but only two are to be considered in any detail. These two are

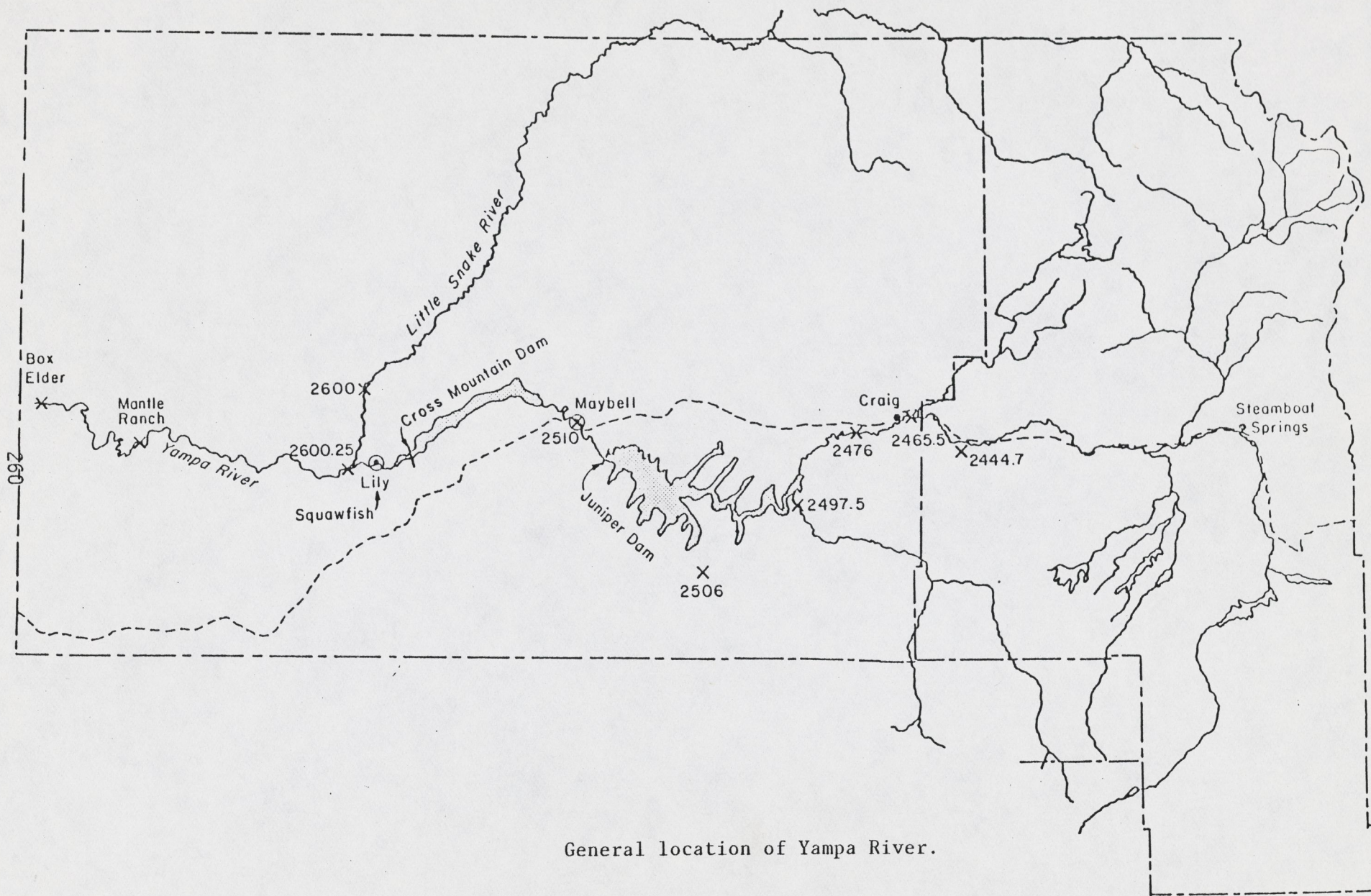
1. The Box Elder Reach, and
2. The Lily Park Reach.

Lily Park Reach is located just above the Little Snake River and the Box Elder Reach is located just above the junction of the Yampa River with the Green River. Data were obtained from the U.S. Geological Survey and from the U.S. Fish and Wildlife Service. The full problem set is available from the Instream Flow Group, U.S. Fish and Wildlife Service.

CONTENTS¹

- General location of Yampa River Basin, Colorado
- Monthly Streamflow Data for Yampa River Basin, Colorado
- Annual Peak Flow Frequency Analysis
- Average Monthly Discharge for Yampa River Basin, Colorado
- Average Daily Discharge for Yampa River Basin, Colorado
- Yampa River Cross Section

¹Only selected data are provided in this appendix. Contact the Instream Flow Group of the Fish and Wildlife Service for the complete data set.



General location of Yampa River.

YAMPA RIVER BELOW PROPOSED JUNIPER RESERVOIR
SIMULATION OF MONTHLY FLOWS
UNITS OF DISCHARGE ARE CUSECS

YEAR	OCT	NOV	DEC	JAN	FEB	MAR	APR	MAY	JUNE	JULY	AUG	SEPT	ANNUAL
1927	0.	0.	0.	0.	0.	624.	200.	200.	200.	200.	200.	0.	137.
1928	0.	0.	0.	272.	401.	200.	200.	200.	200.	200.	200.	200.	171.
1929	0.	0.	0.	465.	446.	0.	200.	200.	200.	200.	200.	200.	177.
1930	200.	200.	214.	499.	532.	399.	200.	200.	200.	200.	200.	0.	253.
1931	0.	0.	353.	511.	496.	0.	200.	200.	200.	0.	0.	0.	162.
1932	0.	0.	0.	0.	0.	0.	200.	200.	200.	200.	200.	0.	84.
1933	0.	0.	0.	0.	0.	485.	200.	200.	200.	200.	0.	0.	108.
1934	0.	0.	0.	0.	0.	0.	411.	200.	0.	0.	0.	0.	51.
1935	0.	0.	0.	0.	0.	0.	508.	200.	200.	0.	0.	0.	75.
1936	0.	0.	0.	0.	0.	0.	200.	200.	200.	200.	0.	0.	67.
1937	0.	0.	0.	0.	0.	0.	386.	200.	200.	0.	0.	0.	65.
1938	0.	0.	0.	0.	0.	551.	200.	200.	200.	200.	0.	0.	114.
1939	0.	0.	0.	0.	0.	482.	200.	200.	200.	0.	0.	0.	91.
1940	0.	0.	0.	0.	0.	0.	292.	200.	200.	0.	0.	0.	57.
1941	0.	0.	0.	0.	0.	0.	404.	200.	200.	0.	0.	0.	67.
1942	0.	0.	647.	0.	0.	477.	200.	200.	200.	200.	0.	0.	152.
1943	0.	0.	0.	0.	0.	0.	200.	200.	200.	0.	0.	0.	50.
1944	0.	0.	0.	0.	0.	0.	489.	200.	200.	0.	0.	0.	74.
1945	0.	0.	0.	0.	0.	0.	383.	200.	200.	200.	0.	0.	82.
1946	0.	0.	0.	0.	0.	541.	200.	200.	200.	0.	0.	0.	95.
1947	0.	0.	0.	0.	0.	481.	200.	200.	200.	200.	0.	0.	108.
1948	0.	0.	0.	774.	698.	646.	200.	200.	200.	200.	0.	0.	241.
1949	0.	0.	0.	0.	0.	0.	200.	200.	200.	200.	200.	0.	84.
1950	0.	0.	0.	728.	0.	0.	310.	200.	200.	0.	0.	0.	121.
1951	0.	0.	0.	0.	0.	609.	206.	200.	200.	200.	0.	0.	119.
1952	0.	0.	0.	0.	750.	0.	200.	200.	200.	200.	200.	200.	150.
1953	0.	0.	0.	0.	0.	0.	480.	200.	200.	0.	0.	0.	73.
1954	0.	0.	0.	0.	0.	0.	390.	200.	0.	0.	0.	0.	49.
1955	0.	0.	0.	748.	0.	0.	344.	200.	200.	0.	0.	0.	125.
1956	0.	0.	0.	0.	0.	0.	200.	200.	200.	200.	0.	0.	67.
1957	247.	0.	0.	0.	0.	0.	324.	200.	200.	200.	200.	200.	131.
1958	200.	200.	200.	200.	229.	301.	234.	200.	200.	200.	200.	200.	214.
1959	0.	200.	0.	659.	0.	0.	424.	200.	200.	0.	0.	0.	191.
1960	0.	672.	0.	762.	0.	0.	200.	200.	200.	0.	0.	0.	170.
1961	0.	0.	0.	0.	0.	0.	571.	200.	200.	0.	0.	0.	80.
1962	502.	0.	679.	0.	475.	440.	200.	200.	200.	200.	200.	0.	258.
1963	0.	0.	0.	0.	0.	621.	277.	200.	200.	0.	0.	0.	109.
1964	0.	0.	0.	0.	0.	0.	470.	200.	200.	0.	0.	0.	72.
1965	0.	0.	0.	0.	0.	656.	0.	200.	200.	200.	200.	0.	123.
1966	200.	0.	200.	351.	0.	474.	200.	200.	200.	0.	0.	0.	154.
1967	0.	0.	0.	0.	0.	0.	413.	200.	200.	0.	0.	0.	67.
1968	0.	0.	0.	0.	703.	0.	377.	200.	200.	200.	0.	0.	136.
1969	0.	0.	0.	0.	0.	0.	200.	200.	200.	200.	0.	0.	67.

YAMPA RIVER BELOW PROPOSED JUNIPER RESERVOIR

(CONTINUED)

1970	0.	0.	640.	0.	0.	609.	263.	200.	200.	200.	200.	200.	211.
1971	0.	200.	200.	295.	426.	247.	200.	200.	200.	200.	200.	200.	213.
1972	200.	198.	0.	200.	200.	200.	200.	200.	200.	0.	0.	0.	132.
1973	0.	0.	0.	0.	730.	0.	368.	200.	200.	200.	200.	0.	154.
1974	0.	0.	0.	710.	0.	0.	200.	200.	200.	200.	200.	0.	144.
1975	0.	0.	0.	0.	0.	0.	410.	200.	200.	200.	0.	0.	84.
1976	0.	0.	0.	0.	0.	0.	422.	200.	200.	0.	0.	0.	68.
AVERAGE	31.	33.	63.	143.	123.	181.	283.	200.	192.	108.	60.	28.	120.
Q/QANN	2.186	2.291	4.428	10.126	7.880	12.769	19.336	14.114	13.113	7.622	4.234	1.912	100.000
COV VAR	3.006	3.299	2.679	1.900	1.934	1.382	.413	0.000	.206	.932	1.543	2.504	.461
SKEW	3.489	4.476	2.916	1.558	1.679	.838	.522	1	-4.841	-.166	.900	2.140	.840
MAXIMUM	502.	672.	679.	774.	750.	656.	571.	200.	200.	200.	200.	200.	258.
MINIMUM	0.	0.	0.	0.	0.	0.	0.	200.	0.	0.	0.	0.	49.

252

A-4

STATISTICAL PARAMETERS FOR STATION CP18
SIMULATION OF MONTHLY FLOWS

YAMPA RIVER BELOW PROPOSED JUNIPER RESERVOIR

MONTH	ARTH. AVERAGE	LOG PARAMETERS				
		MEAN	VARIANCE	STD. DEV.	SKEW	
1	OCT	30.98	-2.3540	3.1249	1.7677	2.4158
2	NOV	33.40	-2.3534	3.1329	1.7700	2.4192
3	DEC	62.74	-2.1152	4.2026	2.0500	1.9218
4	JAN	143.48	-1.4139	6.6131	2.5716	1.0169
5	FEB	122.52	-1.6379	6.0033	2.4502	1.2602
6	MAR	180.94	-.8523	7.6895	2.7730	.5140
7	APR	283.12	2.3237	.6137	.7834	-6.6479
8	MAY	200.00	2.3010	.0000	.0000	1.0312
9	JUNE	192.00	2.0890	1.1011	1.0493	-4.8413
10	JULY	108.00	-.1374	7.1227	2.6688	-.1655
11	AUG	60.00	-1.4097	6.0216	2.4539	.9001
12	SEPT	28.00	-2.2579	3.4524	1.8581	2.1397
13	ANNUAL	120.27	2.0363	.0389	.1972	.1026

SAMPLE SIZE 50 YEARS

LOG NORMAL DISTRIBUTION STATION CP18
SIMULATION OF MONTHLY FLOWS

YAMPA RIVER BELOW PROPOSED JUNIPER RESERVOIR

YEARS FLOW IS NOT EXCEEDED

MONTH	9 IN 10	2 IN 3	1 IN 2	1 IN 5	1 IN 10	1 IN 20	Q2 - Q10
OCT	1.	0.	0.	0.	0.	0.	0.
NOV	1.	0.	0.	0.	0.	0.	0.
DEC	3.	0.	0.	0.	0.	0.	0.
JAN	76.	0.	0.	0.	0.	0.	0.
FEB	32.	0.	0.	0.	0.	0.	0.
MAR	504.	2.	0.	0.	0.	0.	0.
APR	2128.	457.	211.	46.	21.	11.	190.
MAY	200.	200.	200.	200.	200.	200.	0.
JUNE	2718.	346.	123.	16.	6.	2.	117.
JULY	1923.	10.	1.	0.	0.	0.	1.
AUG	54.	0.	0.	0.	0.	0.	0.
SEPT	1.	0.	0.	0.	0.	0.	0.
ANNUAL	195.	132.	109.	74.	61.	52.	48.

Q2 - Q10 IS THE 1 IN 2 YEAR FLOW MINUS THE 1 IN 10 YEAR FLOW

LOG-PEARSON TYPE III DISTRIBUTION FOR STATION CPIA
 YAMPA RIVER BELOW PROPOSED JUNIPER RESERVOIR
 SIMULATION OF MONTHLY FLOWS

DATA SKEW K FACTORS

MONTH	SKEW	RETURN PERIOD					
		10 YRS	6.67 YRS	2 YRS	1.25 YRS	1.11 YRS	1.05 YRS
OCT	2.42	1.2639	.1041	-.3494	-.7272	-.7988	-.8239
NOV	2.42	1.2643	.1046	-.3491	-.7277	-.7996	-.8249
DEC	1.92	1.3117	.1788	-.2914	-.7904	-.9255	-.9918
JAN	1.02	1.3398	.3085	-.1613	-.8523	-1.1312	-1.3231
FEB	1.26	1.3406	.2862	-.1860	-.8464	-1.0986	-1.2653
MAR	.51	1.3222	.3731	-.0806	-.8559	-1.2181	-1.4955
APR	-6.65	.6600	.5184	.3960	-.4200	-1.1800	-2.0030
MAY	1.03	1.3397	.3104	-.1590	-.8526	-1.1339	-1.3282
JUNE	-4.84	.6600	.5184	.3960	-.4200	-1.1800	-2.0030
JULY	-.17	1.2535	.4531	.0389	-.8279	-1.3038	-1.7090
AUG	.90	1.3390	.3197	-.1480	-.8540	-1.1470	-1.3530
SEPT	2.14	1.2972	.1514	-.3142	-.7698	-.8793	-.9279
ANNUAL	.10	1.2917	.4183	-.0166	-.8459	-1.2703	-1.6168

264

9-V

FREQUENCY OF FLOWS

YEARS FLOW IS NOT EXCEEDED

MONTH	9 IN 10	2 IN 3	1 IN 2	1 IN 5	1 IN 10	1 IN 20	Q2 - Q10
OCT	1.	0.	0.	0.	0.	0.	0.
NOV	1.	0.	0.	0.	0.	0.	0.
DEC	4.	0.	0.	0.	0.	0.	0.
JAN	108.	0.	0.	0.	0.	0.	0.
FEB	44.	0.	0.	0.	0.	0.	0.
MAR	652.	2.	0.	0.	0.	0.	0.
APR	693.	537.	430.	99.	25.	6.	405.
MAY	200.	200.	200.	200.	200.	200.	0.
JUNE	605.	429.	320.	44.	7.	1.	312.
JULY	1614.	12.	1.	0.	0.	0.	1.
AUG	75.	0.	0.	0.	0.	0.	0.
SEPT	1.	0.	0.	0.	0.	0.	0.
ANNUAL	195.	131.	108.	74.	61.	52.	47.

Q2 - Q10 IS THE 1 IN 2 YEAR FLOW MINUS THE 1 IN 10 YEAR FLOW

LITTLE SNAKE RIVER NEAR LILY, COLORADO
 MEASURED FLOWS IN CUBIC FEET PER SECOND
 UNITS OF DISCHARGE ARE CUSECS

YEAR	OCT	NOV	DEC	JAN	FEB	MAR	APR	MAY	JUNE	JULY	AUG	SEPT	ANNUAL
1927	142.	84.	81.	81.	72.	228.	1835.	3899.	1875.	597.	53.	77.	755.
1928	189.	363.	244.	179.	191.	785.	1468.	4230.	2067.	115.	13.	16.	825.
1929	303.	165.	120.	150.	150.	950.	2000.	5539.	4116.	581.	167.	260.	1212.
1930	255.	147.	85.	70.	120.	255.	1283.	1471.	973.	54.	94.	83.	407.
1931	101.	25.	25.	67.	145.	620.	2074.	2268.	1367.	455.	85.	15.	604.
1932	176.	260.	140.	205.	210.	1135.	2344.	4505.	2919.	529.	56.	43.	1046.
1933	136.	182.	90.	16.	18.	390.	1116.	3056.	3642.	254.	10.	3.	743.
1934	1.	34.	65.	16.	18.	200.	465.	477.	37.	0.	0.	0.	110.
1935	0.	0.	33.	65.	36.	179.	397.	1431.	1666.	114.	28.	55.	334.
1936	2.	59.	130.	98.	108.	311.	1233.	2824.	958.	65.	237.	3.	505.
1937	59.	105.	49.	49.	18.	211.	1175.	3394.	2052.	732.	82.	103.	672.
1938	71.	114.	111.	128.	155.	320.	1197.	3391.	2059.	184.	42.	171.	663.
1939	93.	183.	175.	109.	137.	544.	902.	2196.	646.	6.	0.	6.	418.
1940	43.	47.	36.	22.	46.	225.	841.	2722.	767.	8.	4.	29.	359.
1941	202.	70.	35.	40.	56.	338.	779.	2947.	1265.	87.	534.	145.	545.
1942	323.	242.	178.	118.	112.	447.	1801.	2247.	1986.	176.	4.	0.	636.
1943	48.	70.	47.	66.	93.	230.	1237.	1399.	2274.	144.	43.	7.	470.
1944	17.	66.	73.	65.	72.	127.	622.	2446.	2737.	242.	7.	0.	540.
1945	80.	59.	78.	66.	90.	160.	640.	2689.	2957.	823.	213.	73.	662.
1946	102.	136.	86.	69.	139.	325.	1504.	1826.	1058.	68.	50.	4.	447.
1947	154.	179.	91.	67.	95.	534.	966.	3150.	1919.	348.	84.	120.	645.
1948	177.	141.	72.	60.	60.	85.	640.	2304.	1052.	100.	2.	0.	393.
1949	22.	101.	115.	114.	118.	248.	1335.	3597.	2797.	396.	20.	9.	741.
1950	193.	126.	101.	78.	80.	285.	1172.	2588.	2374.	290.	11.	16.	610.
1951	53.	102.	73.	61.	145.	239.	597.	1970.	1452.	164.	15.	4.	407.
1952	243.	92.	81.	71.	83.	89.	3259.	4817.	2927.	241.	123.	31.	1005.
1953	31.	51.	65.	75.	75.	308.	431.	1441.	1828.	84.	60.	2.	371.
1954	9.	94.	72.	75.	118.	216.	885.	1112.	305.	25.	0.	41.	246.
1955	45.	63.	68.	57.	59.	272.	587.	1600.	961.	77.	58.	2.	322.
1956	3.	69.	146.	115.	89.	1240.	1172.	2611.	1200.	38.	82.	1.	567.
1957	7.	63.	54.	59.	80.	210.	591.	2438.	3632.	1100.	129.	45.	701.
1958	103.	150.	109.	94.	165.	309.	777.	3548.	1666.	69.	6.	17.	527.
1959	29.	62.	61.	58.	83.	253.	421.	1128.	1002.	146.	64.	270.	294.
1960	174.	162.	85.	51.	68.	151.	1474.	1534.	1222.	53.	3.	4.	414.
1961	34.	44.	44.	36.	54.	279.	320.	1016.	795.	5.	10.	57.	225.
1962	213.	159.	155.	139.	452.	1260.	2663.	2727.	1412.	249.	6.	0.	786.
1963	14.	53.	59.	46.	124.	232.	483.	1460.	747.	18.	76.	57.	241.
1964	7.	51.	40.	51.	57.	81.	390.	2338.	1918.	294.	21.	4.	439.
1965	6.	75.	102.	92.	95.	191.	957.	2591.	2724.	612.	178.	314.	652.
1966	224.	146.	130.	135.	96.	1119.	1006.	1542.	616.	28.	4.	0.	423.
1967	54.	67.	89.	63.	80.	341.	524.	1691.	2298.	448.	31.	59.	479.
1968	75.	65.	104.	101.	128.	262.	591.	2882.	3243.	367.	94.	29.	652.
1969	91.	105.	92.	91.	87.	429.	1482.	2664.	1271.	296.	30.	43.	558.

265

LOG-PEARSON TYPE III DISTRIBUTION FOR STATION 09260000
 LITTLE SNAKE RIVER NEAR LILY, COLORADO
 MEASURED FLOWS IN CUBIC FEET PER SECOND

DATA SKEW K FACTORS

MONTH	SKFW	RETURN PERIOD					
		10 YRS	6.67 YRS	2 YRS	1.25 YRS	1.11 YRS	1.05 YRS
OCT	-3.26	.6600	.5184	.3960	-.4200	-1.1800	-2.0030
NOV	-5.97	.6600	.5184	.3960	-.4200	-1.1800	-2.0030
DEC	-.40	1.2307	.4686	.0664	-.8158	-1.3171	-1.7506
JAN	-.94	1.1154	.5184	.1736	-.7502	-1.3406	-1.8872
FEB	-.39	1.2298	.4691	.0673	-.8154	-1.3175	-1.7519
MAR	.39	1.3016	.4063	-.0342	-.8502	-1.2571	-1.5839
APR	.11	1.2907	.4195	-.0148	-.8455	-1.2716	-1.6193
MAY	-.85	1.1372	.5116	.1563	-.7633	-1.3395	-1.8678
JUNE	-2.55	.7345	.5333	.3723	-.4881	-1.2304	-2.0125
JULY	-1.66	.9602	.5443	.2735	-.6534	-1.3217	-1.9755
AUG	-2.36	.7863	.5407	.3543	-.5301	-1.2576	-2.0114
SEPT	-1.77	.9377	.5457	.2855	-.6383	-1.3157	-1.9833
ANNUAL	-1.04	1.0945	.5238	.1889	-.7373	-1.3404	-1.9035

266

A-8

FREQUENCY OF FLOWS

YEARS FLOW IS NOT EXCEEDED

MONTH	9 IN 10	2 IN 3	1 IN 2	1 IN 5	1 IN 10	1 IN 20	Q2 - Q10
OCT	184.	138.	107.	20.	4.	1.	103.
NOV	245.	192.	156.	38.	10.	3.	145.
DEC	153.	106.	88.	57.	45.	37.	43.
JAN	136.	99.	83.	50.	37.	28.	46.
FEB	216.	131.	100.	56.	40.	30.	60.
MAR	725.	399.	297.	172.	131.	106.	166.
APR	1891.	1170.	921.	582.	461.	380.	460.
MAY	3955.	2964.	2515.	1646.	1261.	988.	1254.
JUNE	2747.	2353.	2078.	1070.	604.	331.	1474.
JULY	629.	327.	214.	50.	17.	6.	196.
AUG	132.	75.	49.	6.	1.	0.	48.
SEPT	159.	49.	22.	1.	0.	0.	22.
ANNUAL	841.	658.	570.	383.	295.	232.	275.

Q2 - Q10 IS THE 1 IN 2 YEAR FLOW MINUS THE 1 IN 10 YEAR FLOW

STATISTICAL PARAMETERS FOR STATION 09260000 LITTLE SNAKE RIVER NEAR LILY, COLORADO
 MEASURED FLOWS IN CUBIC FEET PER SECOND

	MONTH	ARTH. AVERAGE	LOG PARAMETERS			
			MEAN	VARIANCE	STD. DEV.	SKEW
1	OCT	101.46	1.6760	.7961	.8923	-3.2628
2	NOV	112.75	1.8974	.5565	.7460	-5.9743
3	DEC	94.32	1.9284	.0436	.2088	-.3977
4	JAN	84.88	1.8763	.0535	.2313	-.9400
5	FEB	117.64	1.9814	.0822	.2868	-.3922
6	MAR	383.96	2.4829	.0840	.2898	.3931
7	APR	1077.96	2.9677	.0573	.2394	.1131
8	MAY	2569.34	3.3693	.0402	.2004	-.8482
9	JUNE	1885.27	3.1930	.1121	.3348	-2.5457
10	JULY	275.04	2.1431	.4656	.6823	-1.6609
11	AUG	64.90	1.3364	.9985	.9992	-2.3635
12	SEPT	50.24	.9710	1.7200	1.3115	-1.7708
13	ANNUAL	569.20	2.7205	.0349	.1867	-1.0405

SAMPLE SIZE 50 YEARS

LOG NORMAL DISTRIBUTION STATION 09260000 LITTLE SNAKE RIVER NEAR LILY, COLORADO
 MEASURED FLOWS IN CUBIC FEET PER SECOND

YEARS FLOW IS NOT EXCEEDED

MONTH	9 IN 10	2 IN 3	1 IN 2	1 IN 5	1 IN 10	1 IN 20	Q2 - Q10
OCT	660.	114.	47.	8.	3.	2.	44.
NOV	714.	165.	79.	19.	9.	5.	70.
DEC	157.	104.	85.	57.	46.	38.	39.
JAN	149.	95.	75.	48.	38.	31.	37.
FEB	223.	127.	96.	55.	41.	32.	55.
MAR	715.	405.	304.	173.	129.	101.	175.
APR	1882.	1176.	928.	584.	458.	375.	470.
MAY	4229.	2853.	2340.	1587.	1295.	1096.	1045.
JUNE	4190.	2171.	1560.	815.	580.	439.	979.
JULY	1042.	273.	139.	37.	19.	10.	120.
AUG	414.	58.	22.	3.	1.	0.	21.
SEPT	449.	34.	9.	1.	0.	0.	9.
ANNUAL	912.	632.	525.	366.	303.	259.	223.

Q2 - Q10 IS THE 1 IN 2 YEAR FLOW MINUS THE 1 IN 10 YEAR FLOW

LITTLE SNAKE RIVER NEAR LILY, COLORADO

(CONTINUED)

1970	148.	135.	86.	94.	190.	241.	651.	3503.	2979.	464.	43.	63.	718.
1971	183.	167.	133.	141.	159.	612.	1607.	3370.	3606.	584.	25.	55.	887.
1972	123.	147.	154.	140.	424.	752.	808.	1697.	1625.	102.	7.	3.	497.
1973	73.	153.	106.	100.	91.	127.	1061.	3572.	2517.	586.	109.	84.	717.
1974	98.	166.	132.	106.	113.	189.	1073.	4140.	2540.	281.	48.	7.	743.
1975	57.	104.	78.	71.	96.	246.	409.	2605.	2526.	758.	119.	76.	597.
1976	88.	134.	139.	124.	330.	419.	653.	2374.	1685.	296.	68.	9.	527.
AVERAGE	101.	113.	94.	85.	118.	384.	1078.	2569.	1885.	275.	65.	50.	569.
Q/QANN	1.513	1.627	1.406	1.266	1.599	5.725	15.555	38.311	27.204	4.101	.968	.725	100.000
COV VAR	.834	.585	.460	.468	.728	.782	.579	.412	.506	.926	1.359	1.422	.380
SKEW	.777	1.314	.970	.777	2.430	1.752	1.389	.537	.324	1.152	3.433	2.273	.546
MAXIMUM	323.	363.	244.	205.	452.	1260.	3259.	5539.	4116.	1100.	534.	314.	1212.
MINIMUM	0.	0.	25.	16.	18.	81.	320.	477.	37.	0.	0.	0.	110.

268

A-10

YAMPA RIVER AT MAYHELL, COLORADO
 MEASURED FLOWS IN CUBIC FEET PER SECOND
 UNITS OF DISCHARGE ARE CUSECS

YEAR	OCT	NOV	DEC	JAN	FEB	MAR	APR	MAY	JUNE	JULY	AUG	SEPT	ANNUAL
1927	305.	305.	330.	320.	310.	590.	3189.	8261.	6118.	1564.	508.	356.	1851.
1928	499.	673.	540.	420.	520.	1445.	2879.	9963.	5733.	1401.	470.	362.	2083.
1929	409.	476.	380.	400.	400.	1900.	4941.	11270.	8953.	2566.	792.	972.	2795.
1930	697.	570.	480.	340.	380.	785.	4349.	4396.	4022.	569.	623.	437.	1469.
1931	682.	691.	505.	325.	360.	740.	2750.	3864.	2974.	468.	193.	152.	1142.
1932	300.	300.	247.	177.	195.	915.	3361.	8282.	6167.	2073.	631.	259.	1915.
1933	341.	392.	245.	170.	300.	470.	1890.	4975.	7573.	822.	269.	163.	1465.
1934	183.	191.	195.	115.	340.	535.	1567.	2450.	548.	20.	27.	28.	517.
1935	121.	195.	180.	262.	368.	455.	1173.	3886.	6336.	1174.	283.	139.	1213.
1936	159.	234.	215.	240.	270.	440.	4285.	7634.	4204.	730.	352.	136.	1579.
1937	210.	201.	180.	175.	260.	600.	1673.	6436.	4098.	1208.	288.	177.	1297.
1938	315.	353.	376.	357.	434.	818.	2854.	6889.	6100.	1121.	311.	401.	1696.
1939	247.	314.	355.	252.	254.	1216.	2964.	6136.	2926.	338.	144.	217.	1284.
1940	304.	250.	162.	154.	249.	563.	2412.	6076.	3264.	391.	74.	91.	1169.
1941	345.	312.	248.	221.	296.	639.	1546.	7338.	4036.	811.	339.	191.	1356.
1942	654.	474.	391.	343.	345.	815.	4009.	5874.	5524.	952.	226.	84.	1941.
1943	171.	247.	208.	188.	244.	753.	3190.	3850.	4539.	1143.	323.	144.	1249.
1944	166.	239.	199.	168.	179.	292.	735.	5064.	5827.	1039.	131.	33.	1174.
1945	142.	268.	208.	197.	177.	381.	1497.	7143.	6611.	2648.	915.	333.	1716.
1946	274.	319.	263.	227.	330.	645.	3614.	3584.	3831.	704.	261.	149.	1181.
1947	333.	413.	319.	225.	269.	1233.	2760.	8007.	5327.	1929.	527.	278.	1809.
1948	352.	445.	624.	610.	644.	924.	3277.	7368.	4035.	837.	319.	97.	1531.
1949	263.	278.	258.	250.	288.	711.	3223.	6868.	7270.	1953.	334.	178.	1824.
1950	391.	335.	252.	245.	277.	448.	2239.	4410.	5497.	1273.	218.	191.	1314.
1951	250.	287.	268.	235.	268.	533.	1852.	5356.	5299.	1761.	475.	214.	1403.
1952	334.	265.	202.	228.	252.	306.	4033.	8394.	7971.	1186.	542.	243.	1997.
1953	190.	208.	222.	237.	220.	401.	1218.	3602.	6116.	842.	396.	102.	1145.
1954	151.	264.	189.	237.	272.	382.	1728.	3398.	1436.	273.	122.	183.	721.
1955	374.	273.	204.	202.	191.	456.	2009.	4881.	3355.	503.	248.	67.	1067.
1956	133.	263.	322.	277.	246.	489.	3598.	6518.	4358.	539.	280.	64.	1426.
1957	126.	213.	187.	206.	234.	467.	2108.	7156.	11430.	5819.	1052.	450.	2459.
1958	469.	483.	405.	328.	504.	675.	2716.	8931.	5539.	573.	178.	165.	1751.
1959	216.	246.	215.	220.	262.	350.	1501.	4306.	4827.	783.	351.	205.	1124.
1960	665.	532.	377.	235.	221.	663.	4035.	4675.	4496.	592.	155.	97.	1394.
1961	197.	254.	211.	199.	215.	309.	948.	3790.	3272.	356.	131.	534.	869.
1962	1001.	566.	353.	322.	743.	733.	6496.	7145.	5119.	1841.	295.	113.	2059.
1963	288.	288.	217.	213.	391.	466.	1324.	4081.	2475.	270.	215.	199.	870.
1964	117.	201.	137.	160.	221.	1128.	5428.	4916.	1348.	317.	177.	177.	1194.
1965	172.	231.	273.	270.	266.	285.	2626.	6280.	7648.	2439.	753.	501.	1813.
1966	604.	421.	326.	335.	280.	1427.	2044.	3858.	1848.	285.	123.	48.	970.
1967	268.	218.	206.	192.	228.	691.	1486.	4063.	5305.	1766.	347.	237.	1252.
1968	273.	246.	195.	222.	244.	454.	1494.	5584.	7832.	1532.	598.	256.	1578.
1969	354.	325.	278.	283.	284.	419.	4173.	6510.	3732.	1216.	340.	337.	1523.

269

A-11

YAMPA RIVER AT MAYBELL, COLOPADO

(CONTINUED)

1970	450.	412.	360.	350.	394.	489.	1374.	8302.	7449.	1992.	442.	287.	1864.
1971	470.	437.	345.	374.	384.	1081.	4649.	6401.	7756.	1901.	320.	254.	2030.
1972	306.	355.	310.	347.	436.	1175.	2116.	4248.	4872.	538.	151.	197.	1253.
1973	400.	382.	351.	305.	284.	428.	1626.	7689.	6022.	2128.	517.	205.	1701.
1974	250.	338.	357.	308.	271.	577.	3775.	9695.	6208.	1236.	314.	89.	1957.
1975	207.	280.	138.	220.	298.	458.	1566.	5439.	7270.	3388.	509.	180.	1666.
1976	247.	298.	272.	246.	343.	531.	1463.	5011.	3712.	997.	357.	165.	1190.
270 AVERAGE	328.	335.	286.	262.	312.	656.	2589.	6015.	5236.	1277.	362.	223.	1492.
Q/QANN	1.864	1.845	1.624	1.491	1.615	3.729	14.252	34.214	28.818	7.262	2.057	1.226	100.000
COV VAR	.549	.363	.367	.328	.360	.520	.478	.324	.385	.767	.588	.720	.290
SKEW	1.568	1.235	1.129	1.436	1.837	1.636	.821	.497	.322	2.291	1.203	2.384	.456
MAXIMUM	1001.	691.	624.	610.	743.	1900.	6496.	11270.	11430.	5819.	1052.	972.	2795.
MINIMUM	117.	191.	137.	115.	160.	221.	735.	2450.	548.	20.	27.	28.	517.

STATISTICAL PARAMETERS FOR STATION 09251000
MEASURED FLOWS IN CUBIC FEET PER SECOND

YAMPA RIVER AT MAYBELL, COLORADO

	MONTH	ARTH. AVERAGE	MEAN	LOG PARAMETERS		
				VARIANCE	STD. DEV.	SKEW
1	OCT	327.80	2.4602	.0477	.2184	.2398
2	NOV	335.22	2.5008	.0206	.1434	.5849
3	DEC	285.60	2.4296	.0225	.1502	.2929
4	JAN	262.18	2.3979	.0180	.1343	.1505
5	FEB	311.60	2.4710	.0186	.1365	.7031
6	MAR	655.58	2.7690	.0399	.1997	.4250
7	APR	2589.26	3.3642	.0447	.2115	-.1107
8	MAY	6015.30	3.7565	.0207	.1439	-.1812
9	JUNE	5235.52	3.6774	.0472	.2173	-1.9386
10	JULY	1276.77	2.9807	.1472	.3837	-1.6247
11	AUG	361.72	2.4789	.0831	.2883	-1.0081
12	SEPT	222.73	2.2530	.0903	.3004	-.4325
13	ANNUAL	1492.17	3.1548	.0180	.1340	-.6993

SAMPLE SIZE 50 YEARS

LOG NORMAL DISTRIBUTION STATION 09251000
MEASURED FLOWS IN CUBIC FEET PER SECOND

YAMPA RIVER AT MAYBELL, COLORADO

YEARS FLOW IS NOT EXCEEDED

MONTH	9 IN 10	2 IN 3	1 IN 2	1 IN 5	1 IN 10	1 IN 20	Q2 - Q10
OCT	550.	358.	289.	189.	151.	126.	137.
NOV	484.	365.	317.	240.	208.	184.	109.
DEC	419.	312.	269.	201.	173.	152.	96.
JAN	372.	285.	250.	193.	168.	150.	82.
FEB	443.	338.	296.	227.	198.	176.	98.
MAR	1059.	716.	587.	399.	326.	276.	262.
APR	4319.	2851.	2313.	1535.	1239.	1038.	1074.
MAY	8729.	6579.	5708.	4318.	3732.	3309.	1975.
JUNE	9037.	5897.	4758.	3122.	2505.	2089.	2253.
JULY	2969.	1397.	957.	455.	308.	224.	648.
AUG	705.	400.	301.	172.	129.	101.	173.
SEPT	435.	241.	179.	100.	74.	57.	105.
ANNUAL	2121.	1630.	1428.	1101.	962.	860.	467.

Q2 - Q10 IS THE 1 IN 2 YEAR FLOW MINUS THE 1 IN 10 YEAR FLOW

LOG-PEARSON TYPE III DISTRIBUTION FOR STATION 09251000
 YAMPA RIVER AT MAYBELL, COLORADO
 MEASURED FLOWS IN CUBIC FEET PER SECOND

DATA SKEW. K FACTORS.

MONTH	SKEW	RETURN PERIOD					
		10 YRS	6.67 YRS	2 YRS	1.25 YRS	1.11 YRS	1.05 YRS
OCT	.24	1.2974	.4115	-.0266	-.8484	-1.2628	-1.5979
NOV	.54	1.3179	.3820	-.0686	-.8552	-1.2287	-1.5190
DEC	.29	1.2926	.4173	-.0181	-.8463	-1.2692	-1.6139
JAN	.15	1.2869	.4238	-.0084	-.8440	-1.2761	-1.6306
FEB	.70	1.3328	.3465	-.1155	-.8570	-1.1835	-1.4241
MAR	.43	1.3150	.3868	-.0620	-.8545	-1.2345	-1.5318
APR	-.11	1.2464	.4585	.0482	-.8246	-1.3081	-1.7232
MAY	-.18	1.2556	.4515	.0362	-.8289	-1.3025	-1.7049
JUNE	-1.94	.8790	.5466	.3144	-.5986	-1.2971	-1.9991
JULY	-1.62	.9512	.5451	.2785	-.6472	-1.3195	-1.9788
AUG	-1.01	1.0877	.5256	.1938	-.7331	-1.3401	-1.9087
SEPT	-.43	1.2052	.4829	.0938	-.8026	-1.3264	-1.7895
ANNUAL	-.70	1.1829	.4940	.1161	-.7899	-1.3330	-1.8191

272

A-14

FREQUENCY OF FLOWS

YEARS FLOW IS NOT EXCEEDED

MONTH	9 IN 10	2 IN 3	1 IN 2	1 IN 5	1 IN 10	1 IN 20	Q2 - Q10
OCT	554.	355.	285.	188.	153.	129.	132.
NOV	490.	359.	310.	239.	211.	192.	99.
DEC	420.	311.	267.	201.	173.	154.	94.
JAN	372.	285.	249.	193.	168.	151.	81.
FEB	450.	330.	285.	226.	204.	189.	81.
MAR	1076.	702.	571.	397.	333.	290.	238.
APR	4245.	2892.	2368.	1548.	1223.	999.	1145.
MAY	8652.	6629.	5776.	4337.	3707.	3244.	2070.
JUNE	7386.	6254.	5568.	3526.	2486.	1750.	3082.
JULY	2217.	1548.	1224.	540.	298.	167.	925.
AUG	620.	427.	343.	185.	124.	85.	219.
SEPT	412.	250.	191.	103.	72.	52.	120.
ANNUAL	2057.	1663.	1480.	1119.	947.	815.	534.

Q2 - Q10 IS THE 1 IN 2 YEAR FLOW MINUS THE 1 IN 10 YEAR FLOW

YAMPA RIVER BELOW PROPOSED CROSS MOUNTAIN DAM
SIMULATED MONTHLY FLOWS
UNITS OF DISCHARGE ARE CUSECS

YEAR	OCT	NOV	DEC	JAN	FEB	MAR	APR	MAY	JUNE	JULY	AUG	SEPT	ANNUAL
1927	57.	52.	57.	59.	57.	517.	344.	1731.	3421.	3226.	2342.	507.	1038.
1928	517.	380.	501.	568.	556.	344.	344.	2479.	3357.	3191.	2336.	1464.	1342.
1929	401.	581.	627.	597.	597.	344.	344.	2859.	3923.	3408.	2394.	1587.	1477.
1930	600.	600.	663.	678.	627.	490.	494.	2158.	3043.	3054.	2369.	1022.	1322.
1931	607.	724.	725.	681.	602.	344.	344.	1720.	2005.	1720.	382.	246.	844.
1932	0.	52.	42.	33.	36.	166.	344.	1720.	3256.	3318.	2365.	960.	1030.
1933	571.	565.	658.	375.	55.	353.	344.	1720.	2855.	3087.	457.	577.	974.
1934	698.	550.	34.	22.	62.	97.	362.	1720.	1393.	0.	0.	0.	412.
1935	0.	32.	30.	48.	67.	83.	431.	1720.	2305.	2427.	440.	528.	679.
1936	442.	39.	37.	44.	51.	80.	344.	1720.	2941.	3083.	1205.	574.	885.
1937	650.	605.	31.	72.	48.	109.	344.	1720.	2424.	2552.	385.	479.	785.
1938	577.	62.	67.	64.	79.	423.	344.	1720.	3186.	3134.	1069.	409.	933.
1939	615.	564.	373.	46.	47.	344.	344.	1720.	2570.	2097.	462.	565.	816.
1940	394.	42.	27.	29.	46.	102.	344.	1720.	2240.	1885.	464.	554.	657.
1941	377.	54.	43.	41.	54.	116.	344.	1720.	2577.	2776.	344.	436.	744.
1942	381.	448.	568.	211.	63.	344.	344.	1720.	3097.	3110.	592.	566.	959.
1943	661.	598.	35.	34.	45.	137.	344.	1720.	2119.	2293.	424.	574.	752.
1944	414.	40.	34.	31.	33.	54.	344.	1720.	2426.	2493.	458.	569.	722.
1945	323.	46.	36.	36.	34.	71.	344.	1720.	3000.	3422.	1766.	510.	949.
1946	606.	611.	209.	41.	59.	418.	344.	1720.	2256.	1720.	417.	569.	751.
1947	453.	72.	55.	41.	49.	344.	344.	1759.	3292.	3292.	2051.	462.	1025.
1948	529.	606.	439.	689.	689.	663.	344.	1828.	3044.	3090.	456.	576.	1082.
1949	675.	284.	45.	46.	53.	129.	344.	1720.	3177.	3290.	2304.	564.	1060.
1950	513.	621.	243.	670.	165.	82.	344.	1720.	2323.	2480.	448.	551.	851.
1951	419.	50.	47.	43.	49.	506.	344.	1720.	2504.	3256.	447.	571.	835.
1952	462.	58.	36.	42.	665.	176.	344.	1720.	3729.	3153.	2349.	1443.	1183.
1953	673.	698.	684.	47.	40.	73.	393.	1720.	2231.	2180.	407.	575.	814.
1954	447.	46.	32.	44.	49.	70.	344.	1720.	1720.	812.	0.	0.	442.
1955	5.	47.	35.	692.	138.	82.	344.	1720.	2021.	1720.	409.	147.	617.
1956	0.	45.	57.	51.	45.	88.	344.	1720.	2641.	3036.	384.	578.	753.
1957	687.	377.	33.	38.	43.	85.	344.	1720.	3912.	4027.	2440.	1480.	1272.
1958	605.	597.	639.	654.	582.	434.	344.	2254.	3309.	3035.	2276.	1435.	1352.
1959	681.	686.	688.	691.	293.	64.	397.	1720.	2120.	1720.	403.	344.	820.
1960	57.	585.	238.	698.	160.	120.	344.	1720.	2407.	2585.	465.	579.	833.
1961	330.	44.	36.	36.	40.	57.	492.	1720.	1720.	1505.	0.	0.	501.
1962	493.	588.	592.	209.	344.	344.	2362.	3234.	3234.	3263.	2297.	1191.	1276.
1963	688.	696.	690.	60.	70.	513.	344.	1720.	1721.	1517.	0.	0.	673.
1964	0.	34.	23.	26.	29.	41.	393.	1720.	2328.	2179.	447.	572.	652.
1965	412.	40.	48.	50.	49.	555.	344.	1720.	2624.	3384.	2393.	344.	1006.
1966	482.	601.	618.	612.	652.	344.	344.	1720.	1834.	1720.	213.	0.	763.
1967	0.	36.	36.	36.	42.	125.	344.	1720.	2206.	2644.	436.	524.	683.
1968	500.	42.	34.	41.	620.	200.	344.	1720.	2890.	3214.	625.	553.	900.
1969	617.	642.	135.	52.	52.	75.	344.	1720.	2641.	3165.	488.	540.	877.

273

YAMPA RIVER BELOW PROPOSED CROSS MOUNTAIN DAM

(CONTINUED)

1970	559.	612.	662.	350.	72.	504.	344.	1720.	3405.	3297.	2337.	1457.	1284.
1971	524.	580.	614.	606.	588.	344.	344.	2078.	3709.	3274.	2302.	1451.	1372.
1972	432.	66.	62.	55.	657.	190.	344.	1720.	2982.	3328.	2344.	498.	1060.
1973	611.	581.	254.	641.	190.	105.	344.	1872.	3444.	3162.	2301.	876.	1205.
1974	652.	644.	671.	147.	55.	83.	377.	1720.	2784.	3549.	1991.	506.	1106.
1975	620.	614.	164.	45.	63.	97.	344.	1720.	2092.	2345.	398.	572.	760.
274 AVERAGE	449.	352.	259.	226.	199.	233.	356.	1805.	2703.	2678.	1104.	634.	921.
Q/QANN	4.141	3.137	2.390	2.084	1.673	2.150	3.175	16.637	24.105	24.679	10.172	5.658	100.000
COV VAR	.487	.786	1.047	1.190	1.195	.743	.094	.128	.229	.295	.842	.662	.275
SKEW	-1.086	-.121	.657	.940	1.202	.701	3.271	3.117	.052	-1.191	.470	.856	.284
MAXIMUM	698.	724.	725.	698.	689.	663.	494.	2859.	3923.	4027.	2440.	1587.	1477.
MINIMUM	0.	32.	23.	22.	29.	41.	344.	1720.	1393.	0.	0.	0.	412.

STATISTICAL PARAMETERS FOR STATION CP19
SIMULATED MONTHLY FLOWS

YAMPA RIVER BELOW PROPOSED CROSS MOUNTAIN DAM

MONTH	ARTH. AVERAGE	MEAN	LOG PARAMETERS			
			VARIANCE	STD. DEV.	SKEW	
1	OCT	449.33	2.0540	3.0867	1.7569	-2.5446
2	NOV	351.78	2.2895	.3038	.5512	-.2931
3	DEC	259.33	2.0920	.3194	.5651	.2546
4	JAN	226.16	2.0115	.3000	.5478	.6300
5	FEB	199.20	2.0179	.2258	.4752	.8152
6	MAR	233.24	2.2389	.1200	.3464	.0940
7	APR	355.98	2.5498	.0013	.0361	3.0898
8	MAY	1805.31	3.2537	.0023	.0477	2.8707
9	JUNE	2702.82	3.4200	.0109	.1042	-.4342
10	JULY	2677.92	3.2883	.8585	.9266	-6.7818
11	AUG	1103.71	2.4546	2.8190	1.6790	-2.9153
12	SEPT	634.39	2.2039	3.1835	1.7842	-2.6458
13	ANNUAL	920.96	2.9473	.0156	.1249	-.4480

SAMPLE SIZE 49 YEARS

275

A-17

LOG NORMAL DISTRIBUTION STATION CP19
SIMULATED MONTHLY FLOWS

YAMPA RIVER BELOW PROPOSED CROSS MOUNTAIN DAM

YEARS FLOW IS NOT EXCEEDED

MONTH	9 IN 10	2 IN 3	1 IN 2	1 IN 5	1 IN 10	1 IN 20	Q2 - Q10
OCT	20246.	642.	113.	4.	1.	0.	113.
NOV	991.	336.	195.	67.	38.	24.	156.
DEC	655.	216.	124.	41.	23.	15.	100.
JAN	517.	176.	103.	36.	20.	13.	82.
FEB	424.	167.	104.	41.	26.	17.	79.
MAR	482.	244.	173.	89.	62.	47.	111.
APR	395.	368.	355.	331.	319.	309.	36.
MAY	2065.	1880.	1794.	1635.	1558.	1497.	235.
JUNE	3578.	2915.	2630.	2149.	1933.	1772.	697.
JULY	29931.	4850.	1942.	322.	126.	58.	1816.
AUG	40463.	1496.	285.	11.	2.	0.	283.
SEPT	30996.	932.	160.	5.	1.	0.	159.
ANNUAL	1281.	1002.	886.	695.	613.	552.	273.

Q2 - Q10 IS THE 1 IN 2 YEAR FLOW MINUS THE 1 IN 10 YEAR FLOW

LUG-PEARSON TYPE III DISTRIBUTION FOR STATION CP19
 YAMPA RIVER BELOW PROPOSED CROSS MOUNTAIN DAM
 SIMULATED MONTHLY FLOWS

DATA SKEW K FACTORS

MONTH	SKEW	RETURN PERIOD					
		10 YRS	6.67 YRS	2 YRS	1.25 YRS	1.11 YRS	1.05 YRS
OCT	-2.54	.7343	.5332	.3724	-.4879	-1.2302	-2.0124
NOV	-.29	1.2440	.4601	.0511	-.8235	-1.3095	-1.7276
DEC	.25	1.2961	.4131	-.0243	-.8478	-1.2646	-1.6024
JAN	.63	1.3265	.3631	-.0942	-.8567	-1.2048	-1.4679
FEB	.82	1.3355	.3351	-.1296	-.8562	-1.1686	-1.3933
MAR	.09	1.2707	.4392	.0160	-.8364	-1.2914	-1.6713
APR	3.09	1.1800	.0202	-.3960	-.6360	-.6600	-.6650
MAY	2.87	1.2199	.0561	-.3783	-.6766	-.7176	-.7287
JUNE	-.43	1.2055	.4827	.0935	-.8027	-1.3263	-1.7891
JULY	-6.78	.6600	.5184	.3960	-.4200	-1.1800	-2.0030
AUG	-2.92	2.0312	.7635	.4163	-.3997	-.7190	-.8649
SEPT	-2.65	.7121	.5295	.3803	-.4687	-1.2164	-2.0109
ANNUAL	-.45	1.2077	.4816	.0913	-.8038	-1.3256	-1.7860

276

A-18

FREQUENCY OF FLOWS

YEARS FLOW IS NOT EXCEEDED

MONTH	9 IN 10	2 IN 3	1 IN 2	1 IN 5	1 IN 10	1 IN 20	Q2 - Q10
OCT	2208.	979.	511.	16.	1.	0.	510.
NOV	944.	349.	208.	68.	37.	22.	171.
DEC	668.	212.	120.	41.	24.	15.	96.
JAN	547.	162.	91.	35.	22.	16.	69.
FEB	449.	150.	90.	41.	29.	23.	61.
MAR	478.	246.	176.	89.	62.	46.	114.
APR	391.	355.	343.	336.	336.	336.	7.
MAY	2051.	1805.	1721.	1665.	1658.	1656.	63.
JUNE	3513.	2953.	2690.	2169.	1913.	1712.	777.
JULY	7940.	5869.	4521.	793.	157.	27.	4364.
AUG	732713.	5451.	1424.	61.	18.	10.	1407.
SEPT	2992.	1408.	763.	23.	1.	0.	762.
ANNUAL	1254.	1017.	909.	703.	605.	530.	304.

Q2 - Q10 IS THE 1 IN 2 YEAR FLOW MINUS THE 1 IN 10 YEAR FLOW

PSM J407 VER 3.4
(REV 10/22/79)

U. S. GEOLOGICAL SURVEY
ANNUAL PEAK FLOW FREQUENCY ANALYSIS
FOLLOWING WRC GUIDELINES BULL. 17-A.

EXECUTION BEGINNING AT DATE, TIME = 5/13/80 134

INPUT FORMAT = 1 WATSTORE PEAK FILE RETRIEVAL

EXPLANATION OF PEAK DISCHARGE QUALIFICATION CODES

J407 FILE MEANING

D	3	DAM FAILURE, NON-RECURRENT FLOW ANOMALY
G	8	DISCHARGE GREATER THAN STATED VALUE
X	3+8	BOTH OF THE ABOVE
L	4	DISCHARGE LESS THAN STATED VALUE
K	6 OR C	KNOWN EFFECT OF REGULATION OR URBANIZATION
H	7	HISTORIC PEAK

REPORT TROUBLE TO WATSTORE USER ASSISTANCE.

U. S. GEOLOGICAL SURVEY
ANNUAL PEAK FLOW FREQUENCY ANALYSIS
FOLLOWING WRC GUIDELINES HULL. 17-A.

RUN-DATE 5/13/80 AT 134 SEQ 1.0001

OPTIONS IN EFFECT -- PLOT NOBC LGPT NODR PPOS NORS EXPR CLIM

STATION - 09251000/USGS YAMPA RIVER NEAR MAYRELL, CO. 1904-1978 09251000/USGS

INPUT DATA SUMMARY

-- YEARS OF RECORD --		HISTORIC	GENERALIZED	SKEW	GAGE BASE	USER-SET OUTLIER CRITERIA	
SYSTEMATIC	HISTORIC	PEAKS	SKEW	OPTION	DISCHARGE	HIGH OUTLIER	LOW OUTLIER
65	0	0	-0.300	WRC WEIGHTED	0.0	--	--

***** NOTICE -- PRELIMINARY MACHINE COMPUTATIONS. *****
***** USER RESPONSIBLE FOR ASSESSMENT AND INTERPRETATION. *****

WCF1341-NO SYSTEMATIC PEAKS WERE BELOW GAGE BASE. 0.0
WCF1951-NO LOW OUTLIERS WERE DETECTED BELOW CRITERION. 2895.6
WCF1631-NO HIGH OUTLIERS OR HISTORIC PEAKS WERE NOTED.

ANNUAL FREQUENCY CURVE PARAMETERS -- LOG-PEARSON TYPE III

	FLOOD BASE DISCHARGE	FLOOD BASE EXCEEDANCE PROBABILITY	LOGARITHMIC MEAN	LOGARITHMIC STANDARD DEVIATION	LOGARITHMIC SKEW
SYSTEMATIC RECORD	0.0	1.0000	3.9871	0.1350	-0.676
W R C ESTIMATE	0.0	1.0000	3.9871	0.1350	-0.501

ANNUAL FREQUENCY CURVE ORDINATES -- DISCHARGES AT SELECTED EXCEEDANCE PROBABILITIES

ANNUAL EXCEEDANCE PROBABILITY	W R C ESTIMATE	SYSTEMATIC RECORD	'EXPECTED- PROBABILITY' ESTIMATE	95-PCT CONFIDENCE LIMITS FOR W R C ESTIMATES	
				LOWER	UPPER
0.9950	3772.2	3589.6	3556.9	3167.3	4312.9
0.9900	4212.8	4057.1	4038.8	3598.4	4757.3
0.9500	5592.6	5524.9	5505.0	4980.4	6132.1
0.9000	6434.5	6416.9	6361.7	5837.2	6968.1
0.8000	7551.0	7588.4	7510.8	6973.3	8088.3
0.5000	9961.1	10050.5	9961.1	9346.7	10627.6
0.2000	12667.4	12669.8	12714.7	11816.2	13737.1
0.1000	14165.0	14040.0	14263.5	13114.6	15551.3
0.0400	15797.9	15462.5	15966.8	14494.1	17585.2
0.0200	16861.6	16346.1	17114.4	15376.8	18936.0
0.0100	17817.6	17110.1	18144.9	16161.5	20165.6
0.0050	18687.7	17779.4	19122.5	16869.3	21296.4
0.0020	19730.6	18547.9	20226.0	17710.5	22665.2

U. S. GEOLOGICAL SURVEY
ANNUAL PEAK FLOW FREQUENCY ANALYSIS
FOLLOWING WRC GUIDELINES BULL. 17-A.

RUN-DATE 5/13/80 AT 134 SEQ 1.0001

STATION - 09251000/USGS YAMPA RIVER NEAR MAYRELL, CO. 1904-1978 09251000/USGS

***** NOTICE -- PRELIMINARY MACHINE COMPUTATIONS. *****
***** USER RESPONSIBLE FOR ASSESSMENT AND INTERPRETATION. *****

INPUT DATA LISTING

EMPIRICAL FREQUENCY CURVES -- WEIHULL PLOTTING POSITIONS

WATER YEAR	DISCHARGE	CODES	WATER YEAR	RANKED DISCHARGE	SYSTEMATIC RECORD	W R C ESTIMATE
1904	8050.0		1917	17900.0	0.0152	0.0152
1905	11400.0		1921	17700.0	0.0303	0.0303
1916	11700.0		1920	16000.0	0.0455	0.0455
1917	17900.0		1957	15700.0	0.0606	0.0606
1918	10500.0		1974	15400.0	0.0758	0.0758
1919	7670.0		1929	14400.0	0.0909	0.0909
1920	16000.0		1952	13800.0	0.1061	0.1061
1921	17700.0		1928	13700.0	0.1212	0.1212
1922	10800.0		1970	12700.0	0.1364	0.1364
1923	10900.0		1947	12400.0	0.1515	0.1515
1924	7810.0		1958	12200.0	0.1667	0.1667
1925	6640.0		1932	12100.0	0.1818	0.1818
1926	9090.0		1938	12100.0	0.1970	0.1970
1927	11800.0		1973	12100.0	0.2121	0.2121
1928	13700.0		1927	11800.0	0.2273	0.2273
1929	14400.0		1965	11800.0	0.2424	0.2424
1930	7980.0		1916	11700.0	0.2576	0.2576
1931	6500.0		1941	11700.0	0.2727	0.2727
1932	12100.0		1975	11700.0	0.2879	0.2879
1933	11200.0		1978	11600.0	0.3030	0.3030
1934	4080.0		1962	11500.0	0.3182	0.3182
1935	9870.0		1905	11400.0	0.3333	0.3333
1936	10600.0		1968	11400.0	0.3485	0.3485
1937	10000.0		1948	11300.0	0.3636	0.3636
1938	12100.0		1933	11200.0	0.3788	0.3788
1939	7860.0		1923	10900.0	0.3939	0.3939
1940	9170.0		1945	10900.0	0.4091	0.4091
1941	11700.0		1922	10800.0	0.4242	0.4242
1942	9930.0		1936	10600.0	0.4394	0.4394
1943	9280.0		1918	10500.0	0.4545	0.4545
1944	9080.0		1971	10300.0	0.4697	0.4697
1945	10900.0		1953	10100.0	0.4848	0.4848
1946	6850.0		1937	10000.0	0.5000	0.5000
1947	12400.0		1964	9990.0	0.5152	0.5152
1948	11300.0		1942	9930.0	0.5303	0.5303
1949	9730.0		1935	9870.0	0.5455	0.5455
1950	8210.0		1956	9870.0	0.5606	0.5606
1951	8870.0		1949	9730.0	0.5758	0.5758
1952	13800.0		1943	9280.0	0.5909	0.5909
1953	10100.0		1940	9170.0	0.6061	0.6061

-- CONTINUED --

U. S. GEOLOGICAL SURVEY
ANNUAL PEAK FLOW FREQUENCY ANALYSIS
FOLLOWING WRC GUIDELINES BULL. 17-A.

RUN-DATE 5/13/80 AT 134 SEQ 1.0001

STATION - 09251000/USGS YAMPA RIVER NEAR MAYBELL, CO. 1904-1978 09251000/USGS

***** NOTICE -- PRELIMINARY MACHINE COMPUTATIONS. *****
***** USER RESPONSIBLE FOR ASSESSMENT AND INTERPRETATION. *****

INPUT DATA LISTING

EMPIRICAL FREQUENCY CURVES -- WEIBULL PLOTTING POSITIONS

WATER YEAR	DISCHARGE	CODES	WATER YEAR	RANKED DISCHARGE	SYSTEMATIC RECORD	W R C ESTIMATE
-- CONTINUED --						
1954	5480.0		1926	9090.0	0.6212	0.6212
1955	7000.0		1944	9080.0	0.6364	0.6364
1956	9870.0		1967	8890.0	0.6515	0.6515
1957	15700.0		1972	8890.0	0.6667	0.6667
1958	12200.0		1951	8870.0	0.6818	0.6818
1959	6690.0		1969	8290.0	0.6970	0.6970
1960	8000.0		1950	8210.0	0.7121	0.7121
1961	6350.0		1904	8050.0	0.7273	0.7273
1962	11500.0		1960	8000.0	0.7424	0.7424
1963	6290.0		1930	7980.0	0.7576	0.7576
1964	9990.0		1939	7860.0	0.7727	0.7727
1965	11800.0		1924	7810.0	0.7879	0.7879
1966	6900.0		1919	7670.0	0.8030	0.8030
1967	8890.0		1976	7450.0	0.8182	0.8182
1968	11400.0		1955	7000.0	0.8333	0.8333
1969	8290.0		1966	6900.0	0.8485	0.8485
1970	12700.0		1946	6850.0	0.8636	0.8636
1971	10300.0		1959	6690.0	0.8788	0.8788
1972	8890.0		1925	6640.0	0.8939	0.8939
1973	12100.0		1931	6500.0	0.9091	0.9091
1974	15400.0		1961	6350.0	0.9242	0.9242
1975	11700.0		1963	6290.0	0.9394	0.9394
1976	7450.0		1954	5480.0	0.9545	0.9545
1977	3620.0		1934	4080.0	0.9697	0.9697
1978	11600.0		1977	3620.0	0.9848	0.9848

U. S. GEOLOGICAL SURVEY
ANNUAL PEAK FLOW FREQUENCY ANALYSIS
FOLLOWING WRC GUIDELINES BULL. 17-A.

RUN-DATE 5/13/80 AT 134 SEQ 1.0001

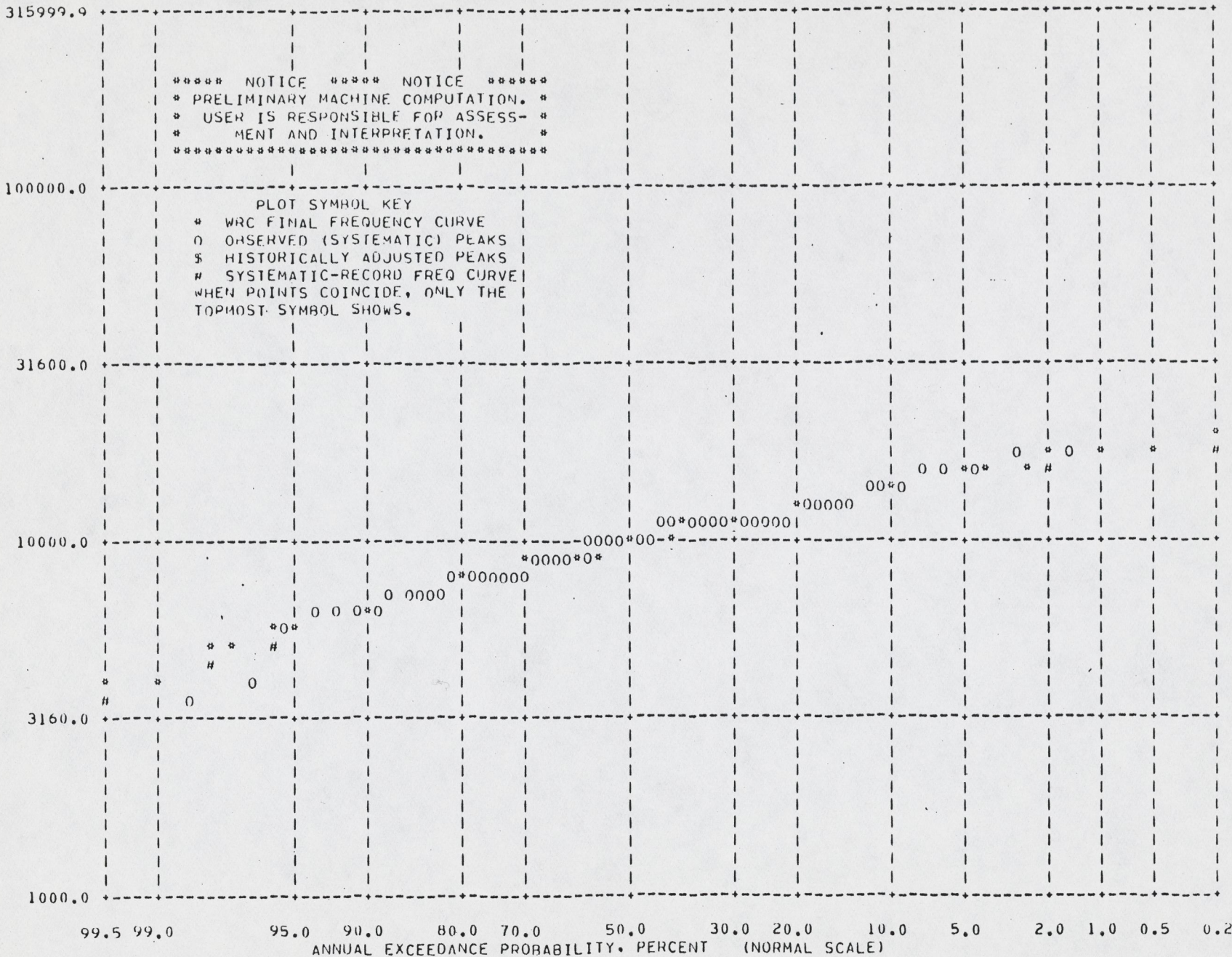
STATION - 09251000/USGS

YAMPA RIVER NEAR MAYRELL, CO.

1904-1978

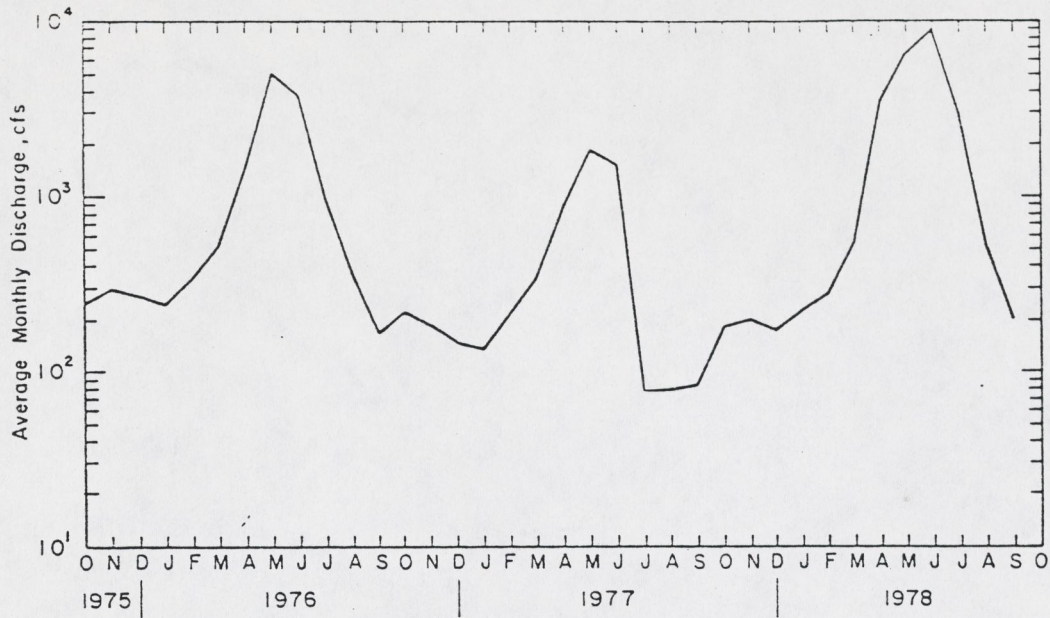
09251000/USGS

ANNUAL PEAK MAGNITUDE / LOG SCALE /



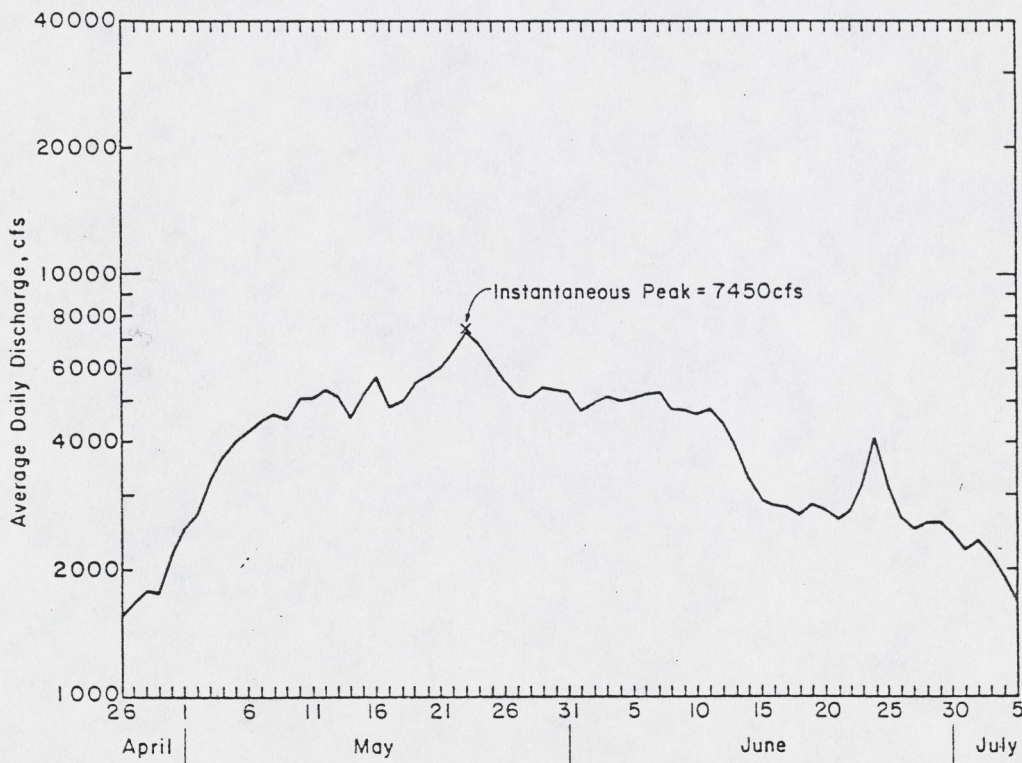
281

A-23

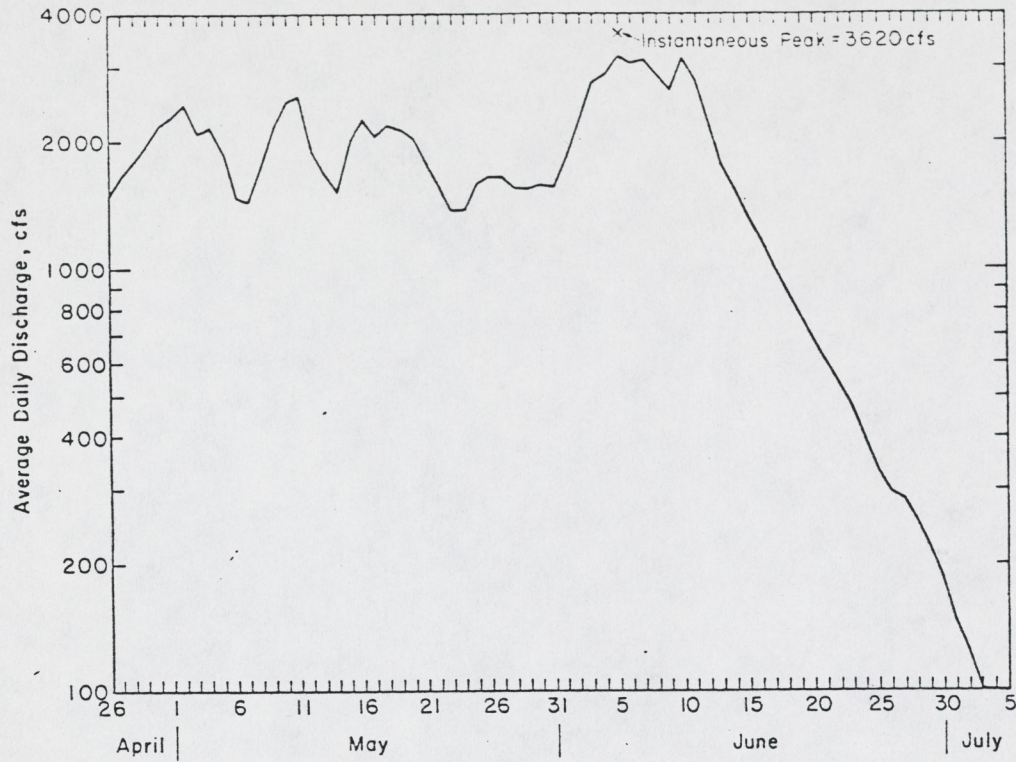


Average monthly discharge for Yampa River near Maybell, Colorado (October 1975-September 1978).

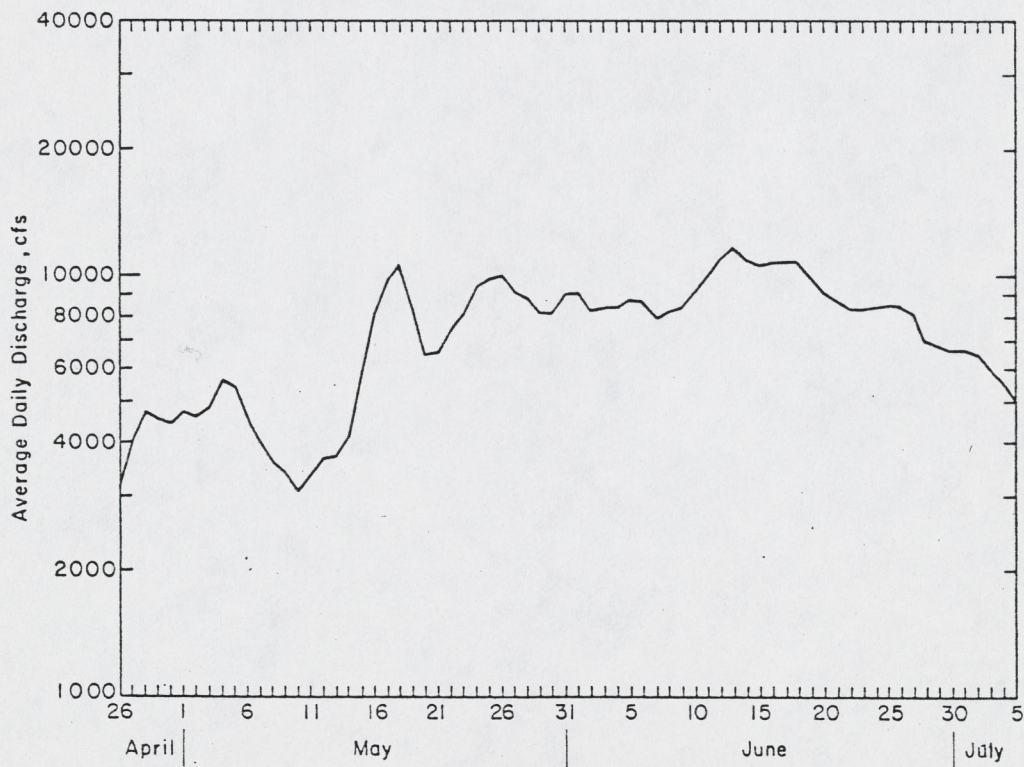
ANNUAL PEAK FLOW FREQUENCY ANALYSIS



Average daily discharges for Yampa River near Maybell, Colorado (April 26-July 5, 1976).



Average daily discharges for Yampa River near Maybell, Colorado (April 26-July 5, 1977).



Average daily discharges for Yampa River near Maybell, Colorado (April 26-July 5, 1978).

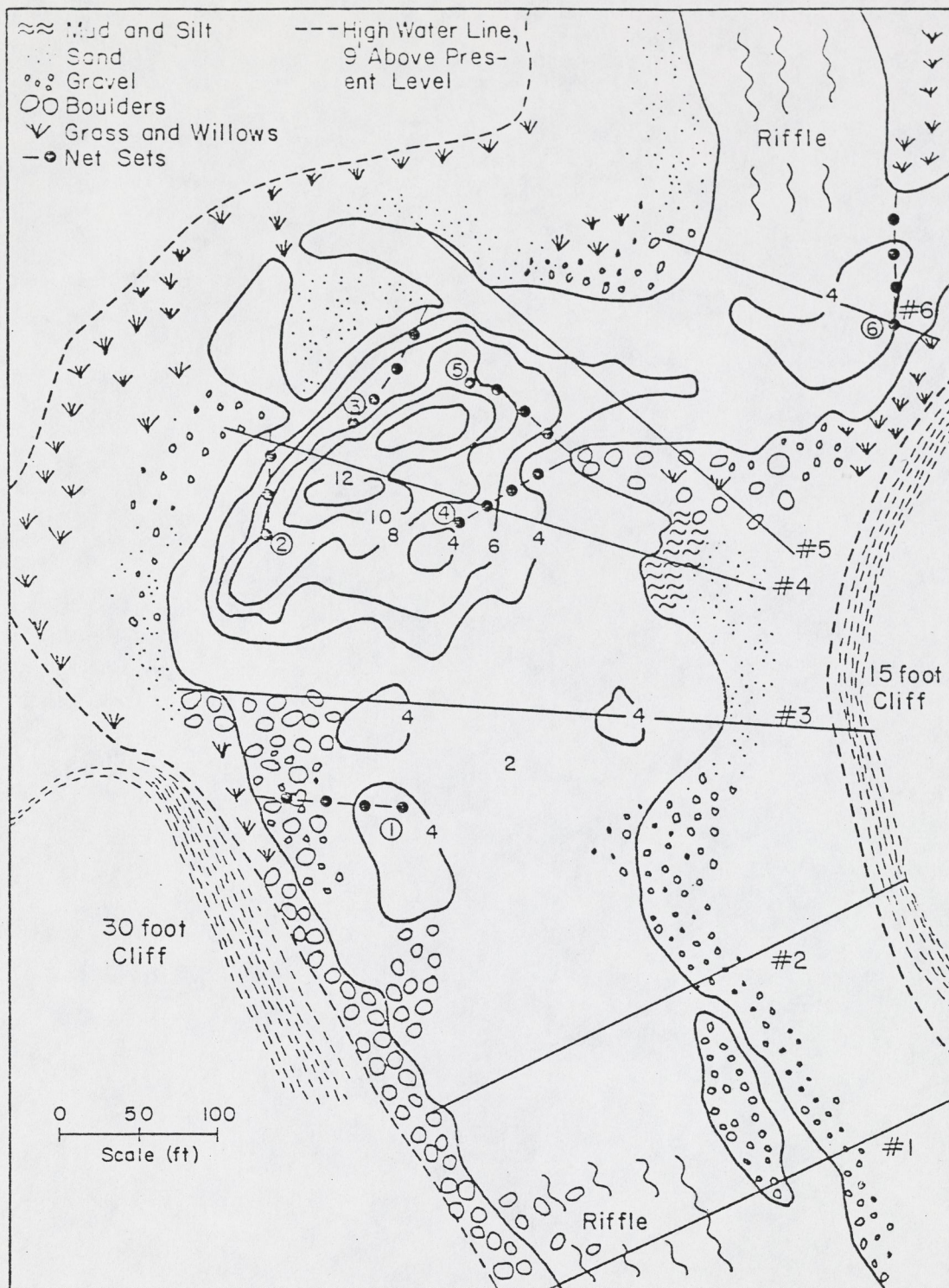
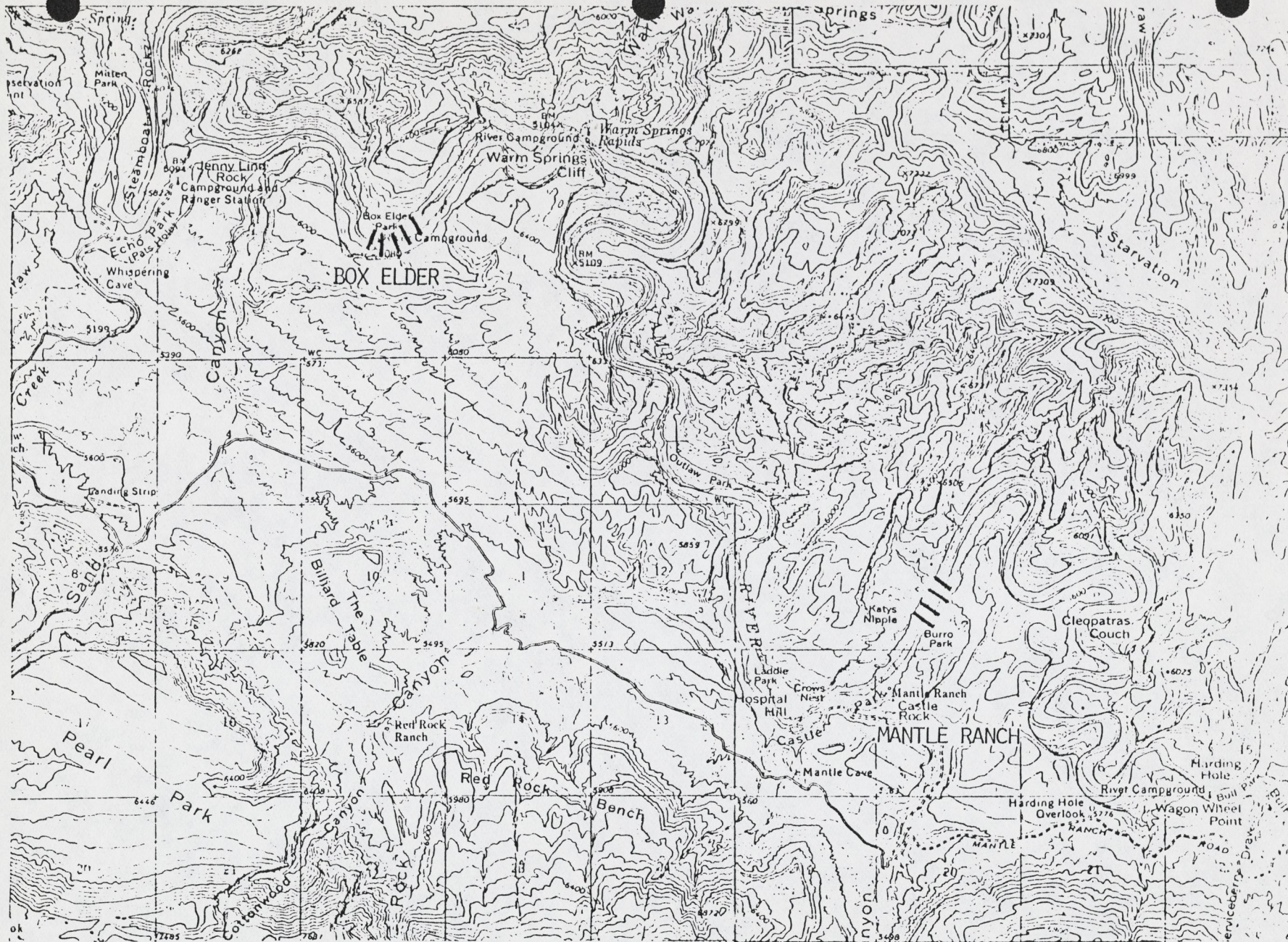
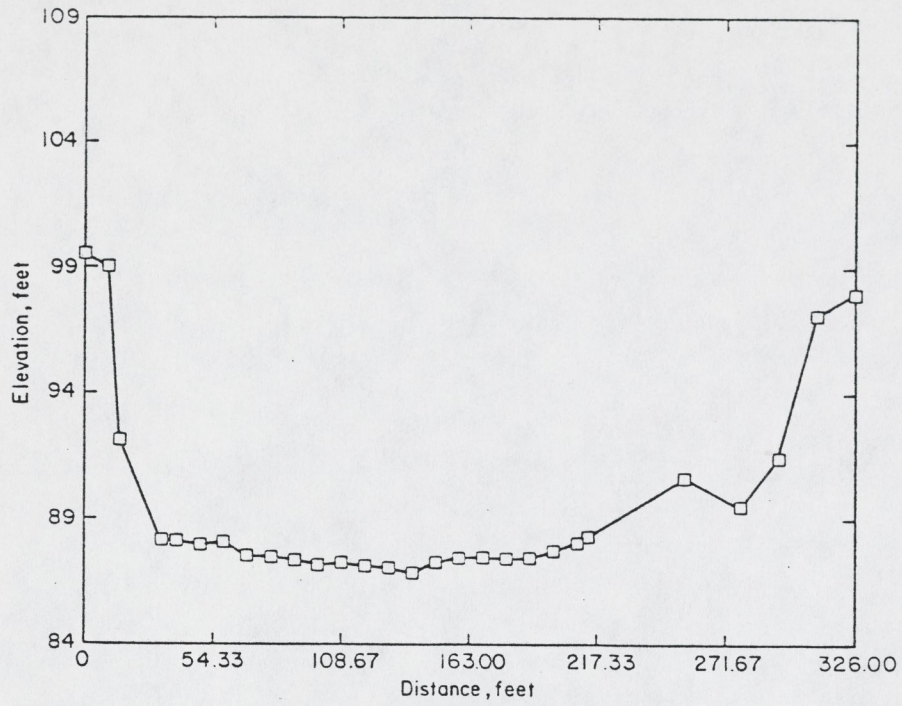


Diagram of a pool riffle habitat used by Colorado squawfish on the Yampa River in Lily Park. (Contours indicate depth in feet, 30 August 1976).

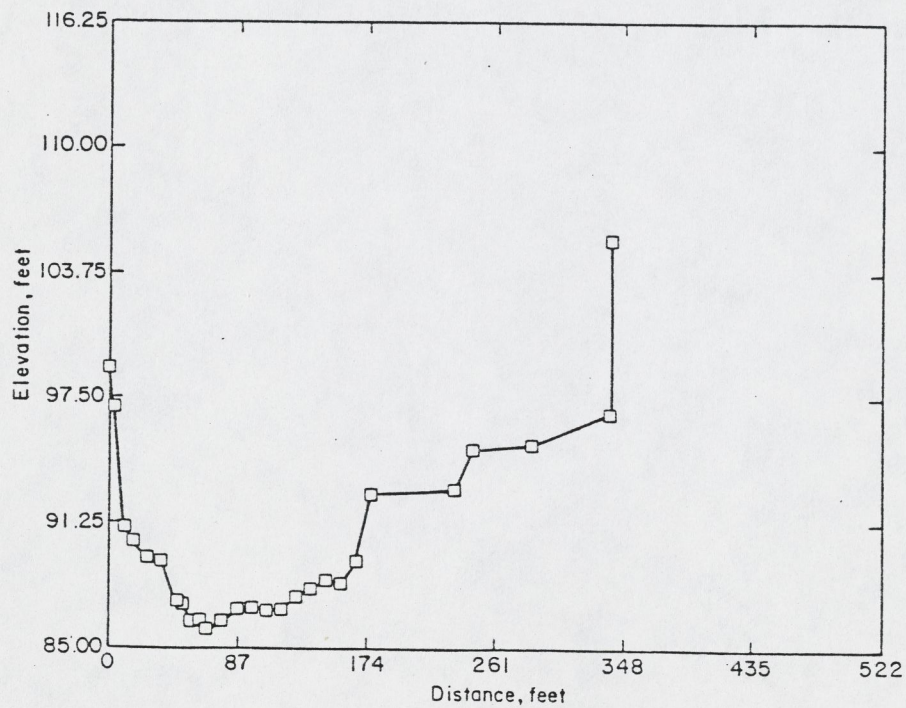


Location of Box Elder and Mantle Ranch reaches.

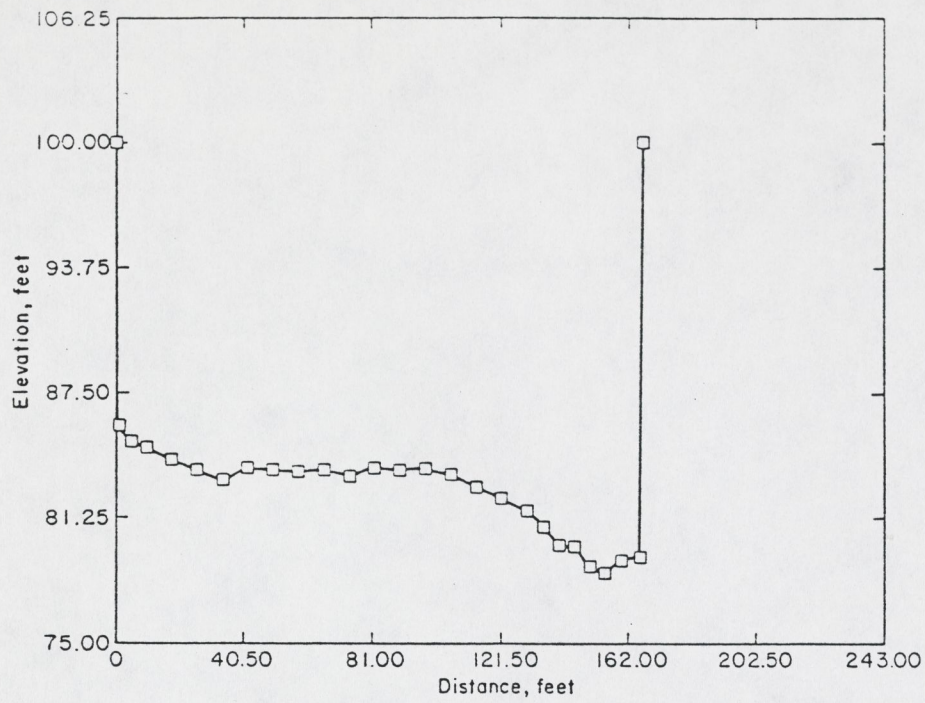
YAMPA RIVER CROSS SECTIONS



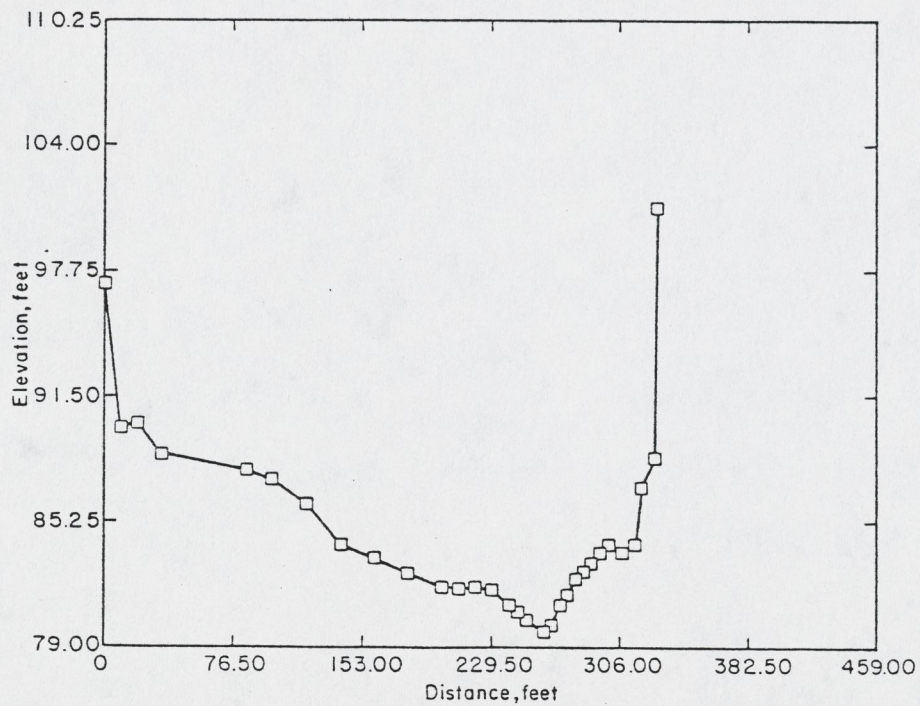
Cross section of Yampa River at Maybell Reach.



Cross section of Yampa River at Lily Park.



Cross section of Yampa River at Dinosaur National Monument near Mantle Ranch.



Cross section of Yampa River at Dinosaur National Monument at Box Elder Reach.

APPENDIX B

CHANNEL CHANGE WORKSHOP: PROBLEM NUMBER 2
POPLAR CREEK, CALIFORNIA

THE QUESTION

The question relative to the Poplar Creek is "What will happen to the morphology of the stream channel as a result of the changes in streamflows and sediment discharge caused by the construction of a reservoir upstream of a reach of stream?" Specifically, for both a short and a long time after construction of the reservoir:

1. What will be the meander pattern?
2. What will be the configuration of the channel?
3. What will be the substrate material?
4. What will be the pool riffle sequence?

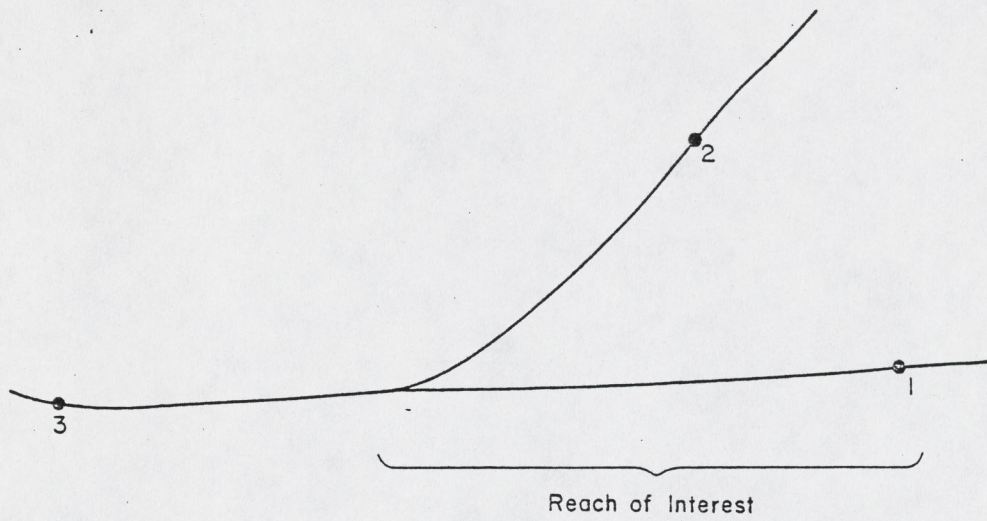
Poplar Creek is located in a semi-arid region of northern California. Poplar is not the actual name of the stream. The problem as presented here has been abstracted from the data for the actual project to the extent it no longer is representative of the concern about the actual project. In order to make it clear the comments on the workshop problem are not necessarily applicable to the actual project, the name of the stream has been changed.

The reach of interest is the reach of the stream from just below Dutch Gulch dam site to the junction with the south fork. Special attention should be given to the reach just downstream of Dry Creek.

CONTENTS¹

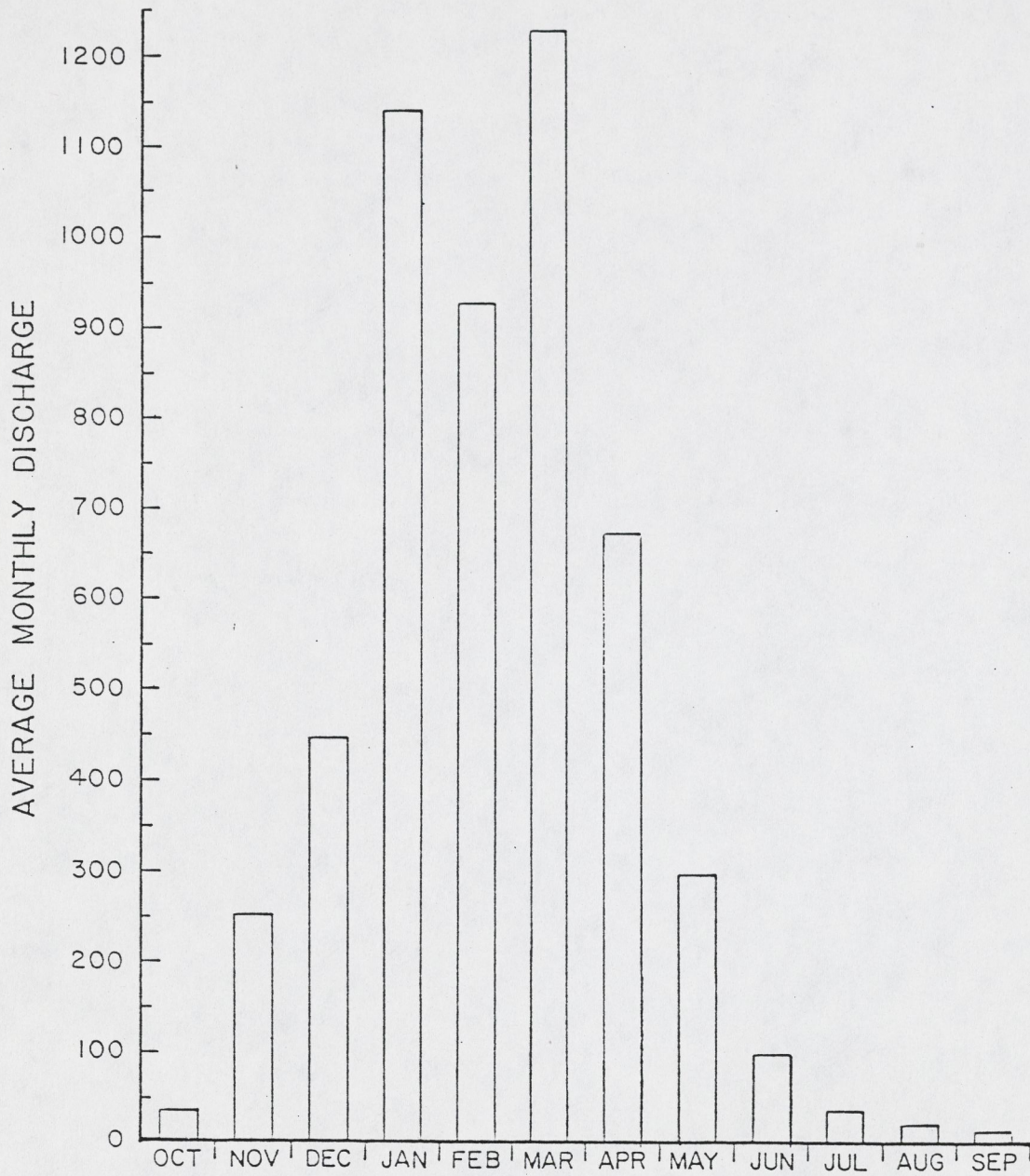
- Location of Reach of Interest
- Monthly Flows
- Cross Sections of Reach of Interest
- Peak Flows and Flow Duration Curves
- Bed Material

¹Only selected data are provided in this appendix. Contact the Instream Flow Group of the Fish and Wildlife Service for the complete data set.

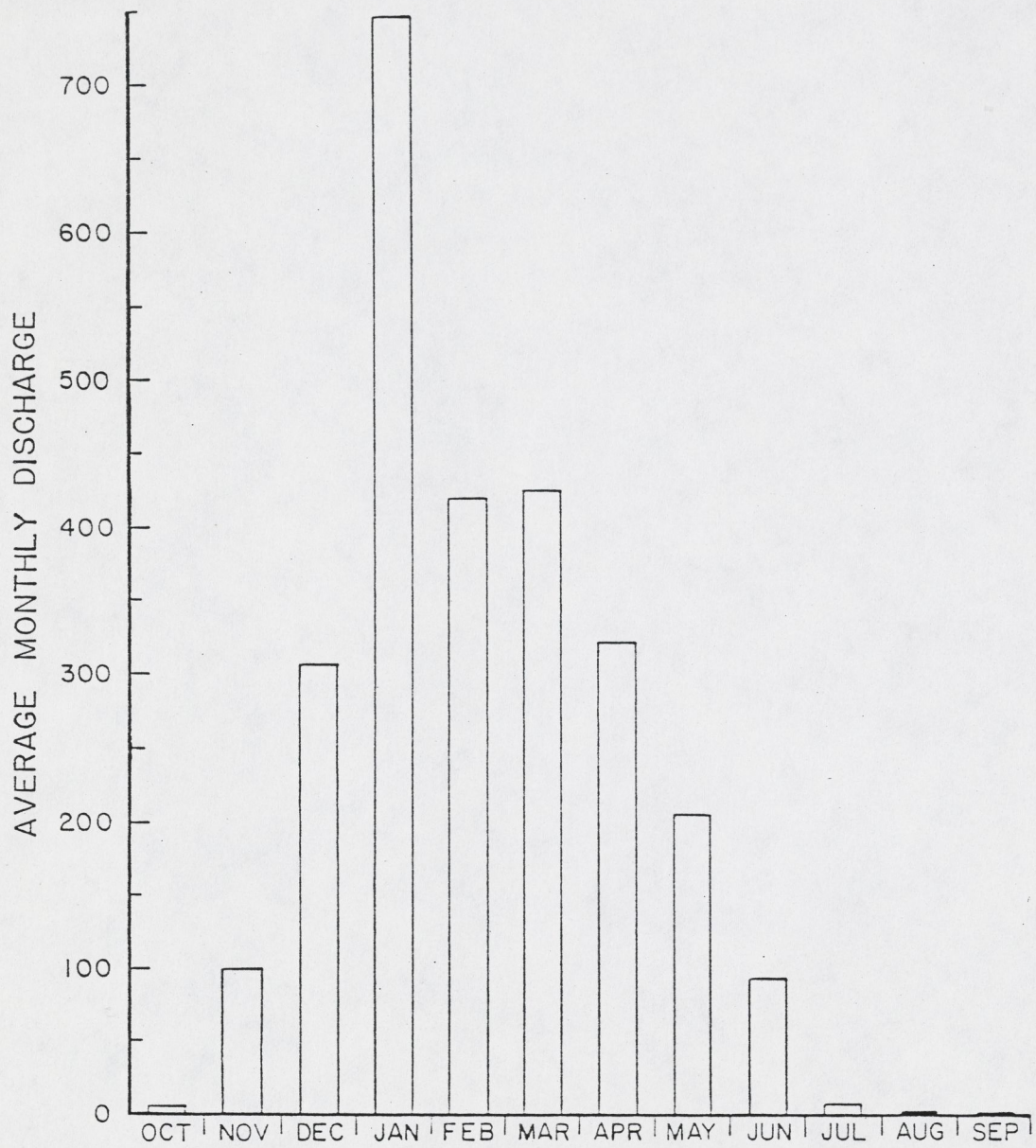


Skematic diagram showing relative location of gaging stations.

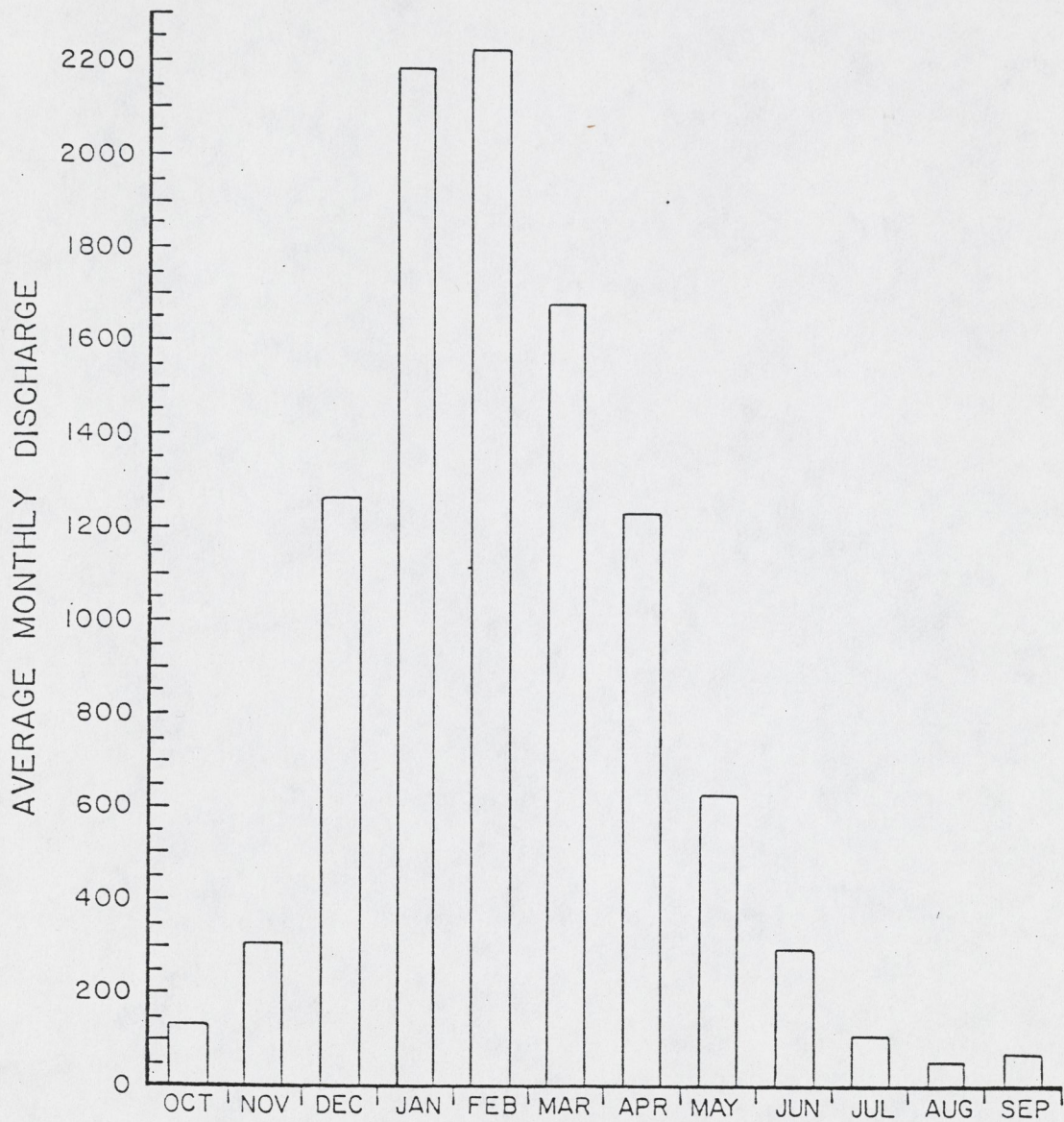
MONTHLY FLOW



Poplar Creek at Station 1 (1972-1978)



Poplar Creek at Station 2 (1963-1978)



Poplar Creek at Station 3 (1941-1979)

POPLAR CREEK AT STATION 1
 CHANNEL CHANGE WORKSHOP PROBLEM
 UNITS OF DISCHARGE ARE CUSECS

YEAR	OCT	NOV	DEC	JAN	FEB	MAR	APR	MAY	JUNE	JULY	AUG	SEPT	ANNUAL
1972	26.	89.	201.	362.	370.	618.	248.	126.	51.	10.	4.	5.	175.
1973	63.	462.	668.	1871.	1855.	1466.	646.	298.	95.	33.	20.	15.	619.
1974	65.	1102.	1604.	3155.	795.	1883.	1620.	419.	153.	95.	49.	21.	916.
1975	32.	61.	158.	215.	1547.	2853.	1000.	590.	197.	73.	28.	21.	560.
1976	70.	114.	146.	84.	305.	298.	240.	95.	19.	7.	28.	21.	118.
1977	17.	35.	28.	52.	45.	78.	43.	63.	17.	1.	0.	10.	32.
1978	17.	99.	699.	2925.	1722.	1811.	1027.	412.	138.	63.	17.	23.	743.
1979	24.	45.	53.	289.	764.	813.	453.	347.	93.	26.	13.	14.	242.
AVERAGE	39.	251.	445.	1119.	925.	1227.	660.	294.	95.	38.	20.	16.	426.
Q/QANN	.780	4.842	8.866	22.317	16.819	24.477	12.729	5.857	1.841	.767	.393	.311	100.000
COV VAR	.580	1.477	1.207	1.183	.753	.767	.798	.632	.681	.908	.780	.404	.766
SKEW	.506	2.206	1.711	.878	.254	.496	.766	.146	.205	.590	.704	-.804	.264
MAXIMUM	70.	1102.	1604.	3155.	1855.	2853.	1620.	590.	197.	96.	49.	23.	916.
MINIMUM	17.	35.	28.	52.	45.	78.	43.	63.	17.	1.	0.	5.	32.

B-6

STATISTICAL PARAMETERS FOR STATION 11375810 POPLAR CREEK AT STATION 1
CHANNEL CHANGE WORKSHOP PROBLEM

MONTH	ARTH. AVERAGE	MEAN	LOG PARAMETERS			
			VARIANCE	STD. DEV.	SKEW	
1	OCT	39.14	1.5261	.0667	.2583	.2088
2	NOV	250.93	2.0870	.2633	.5131	1.1211
3	DEC	444.59	2.3382	.3541	.5950	-.0609
4	JAN	1119.11	2.6535	.4743	.6887	.0705
5	FEB	925.43	2.7734	.2903	.5388	-1.3737
6	MAR	1227.45	2.9024	.2659	.5157	-1.1586
7	APR	659.61	2.6368	.2520	.5020	-1.1302
8	MAY	293.71	2.3622	.1255	.3542	-.6264
9	JUNE	95.38	1.8464	.1671	.4088	-.7648
10	JULY	38.45	1.2634	.5628	.7502	-1.4250
11	AUG	19.70	.7246	2.3693	1.5392	-2.5942
12	SEPT	16.13	1.1612	.0575	.2397	-1.5946
13	ANNUAL	425.61	2.4472	.2463	.4963	-.9517

SAMPLE SIZE 8 YEARS

LOG NORMAL DISTRIBUTION STATION 11375810 POPLAR CREEK AT STATION 1
CHANNEL CHANGE WORKSHOP PROBLEM

YEARS FLOW IS NOT EXCEEDED

MONTH	9 IN 10	2 IN 3	1 IN 2	1 IN 5	1 IN 10	1 IN 20	Q2 - Q10
OCT	72.	43.	34.	20.	16.	13.	18.
NOV	556.	203.	122.	45.	27.	17.	95.
DEC	1262.	392.	218.	69.	38.	23.	180.
JAN	3439.	889.	450.	118.	59.	33.	391.
FEB	2911.	1010.	593.	209.	121.	77.	472.
MAR	3660.	1329.	799.	294.	174.	113.	624.
APR	1907.	712.	433.	164.	98.	65.	335.
MAY	655.	327.	230.	116.	81.	60.	149.
JUNE	235.	105.	70.	32.	21.	15.	49.
JULY	168.	38.	18.	4.	2.	1.	16.
AUG	499.	24.	5.	0.	0.	0.	5.
SEPT	29.	18.	14.	9.	7.	6.	7.
ANNUAL	1212.	457.	280.	107.	65.	43.	215.

Q2 - Q10 IS THE 1 IN 2 YEAR FLOW MINUS THE 1 IN 10 YEAR FLOW

LOG-PEARSON TYPE III DISTRIBUTION FOR STATION 11375A10
 POPLAR CREEK AT STATION 1
 CHANNEL CHANGE WORKSHOP PROBLEM

DATA SKEW K FACTORS

MONTH	SKEW	RETURN PERIOD					
		10 YRS	6.67 YRS	2 YRS	1.25 YRS	1.11 YRS	1.05 YRS
OCT	1.21	1.3002	.4081	-.0316	-.8496	-1.2591	-1.5886
NOV	1.12	1.3408	.2948	-.1766	-.8488	-1.1114	-1.2878
DEC	-.06	1.2653	.4436	.0232	-.8337	-1.2955	-1.6835
JAN	.07	1.2735	.4367	.0120	-.8378	-1.2890	-1.6647
FEB	-1.37	1.0350	.5356	.2289	-.7011	-1.3359	-1.9414
MAR	-1.16	1.0769	.5278	.2012	-.7266	-1.3396	-1.9162
APR	-1.13	1.0706	.5290	.2055	-.7229	-1.3393	-1.9205
MAY	-.63	1.1705	.4993	.1278	-.7826	-1.3352	-1.8337
JUNE	-.76	1.1593	.5036	.1376	-.7761	-1.3371	-1.8457
JULY	-1.43	1.0000	.5404	.2505	-.6788	-1.3300	-1.9592
AUG	-2.59	.7457	.5351	.3685	-.4978	-1.2372	-2.0129
SEPT	-1.59	.9927	.5413	.2548	-.6742	-1.3287	-1.9525
ANNUAL	-.95	1.1179	.5177	.1717	-.7517	-1.3405	-1.8852

296

B-8

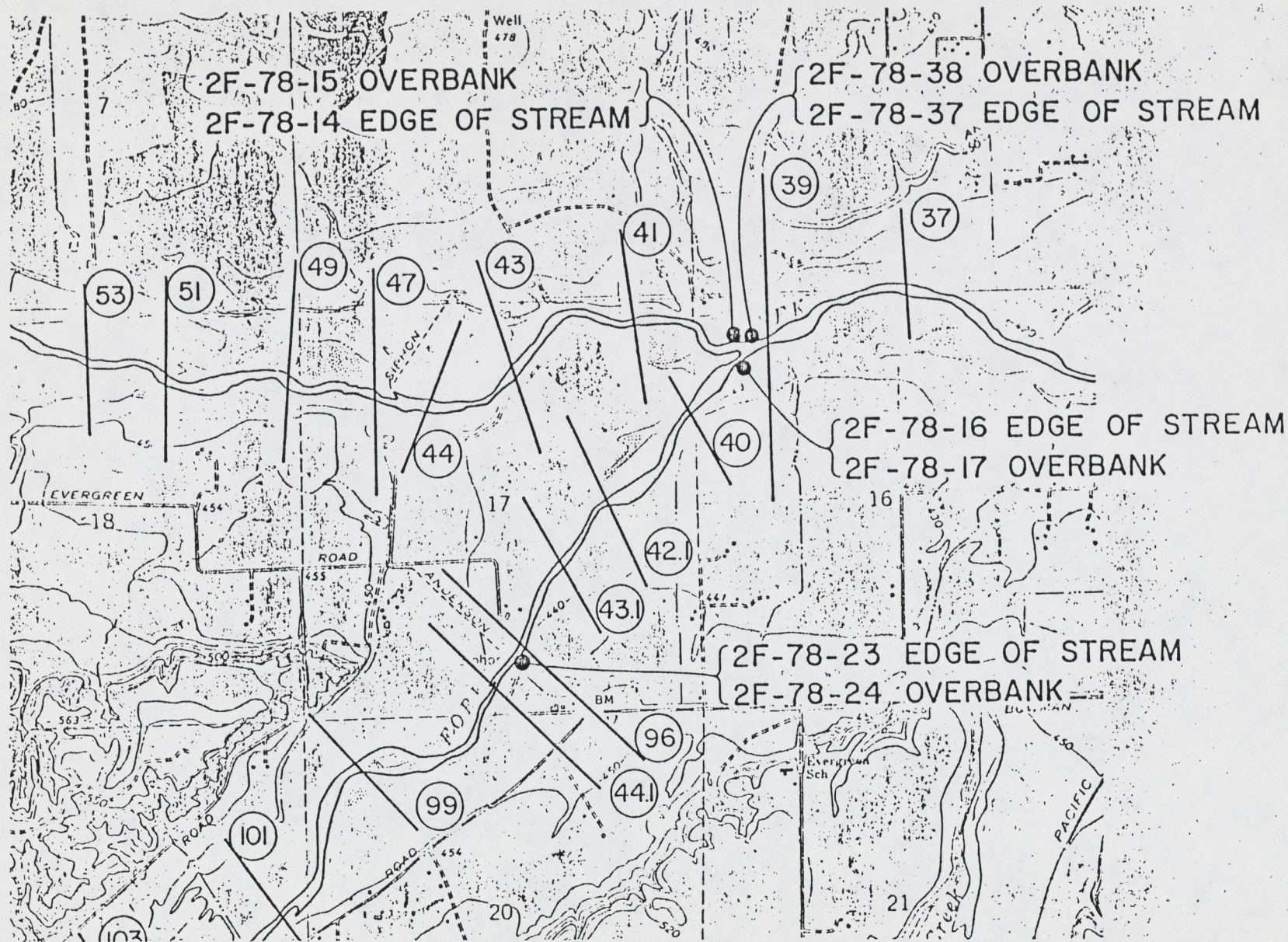
FREQUENCY OF FLOWS

YEARS FLOW IS NOT EXCEEDED

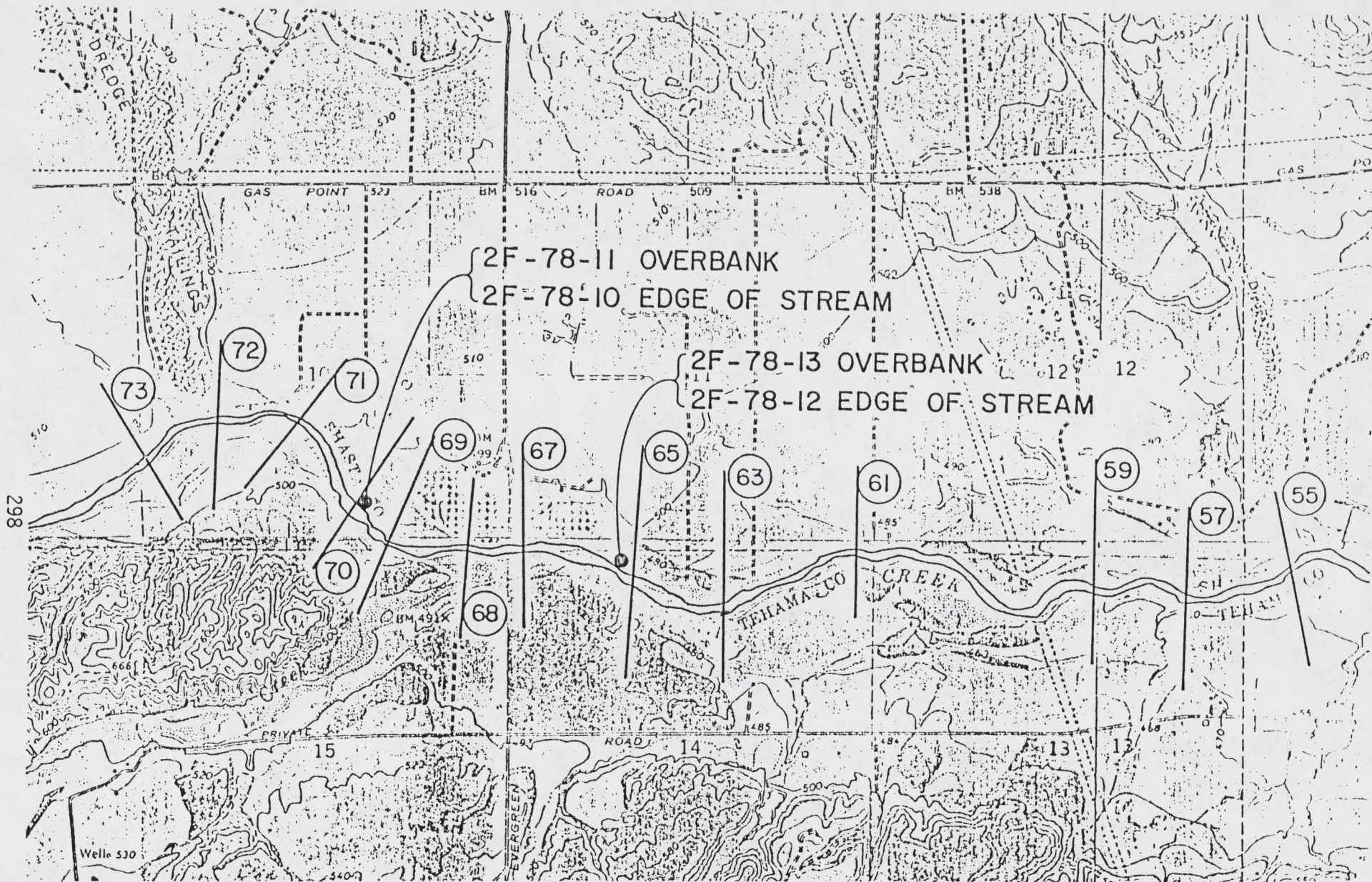
MONTH	9 IN 10	2 IN 3	1 IN 2	1 IN 5	1 IN 10	1 IN 20	Q2 - Q10
OCT	73.	43.	33.	20.	16.	13.	17.
NOV	596.	173.	99.	45.	33.	27.	66.
DEC	1234.	400.	225.	70.	37.	22.	188.
JAN	3393.	900.	459.	119.	58.	32.	401.
FEB	2143.	1153.	788.	249.	113.	53.	675.
MAR	2869.	1495.	1014.	337.	163.	82.	852.
APR	1494.	799.	549.	188.	92.	47.	457.
MAY	598.	346.	256.	122.	77.	52.	178.
JUNE	209.	113.	80.	34.	20.	12.	60.
JULY	103.	47.	28.	6.	2.	1.	25.
AUG	75.	35.	20.	1.	0.	0.	20.
SEPT	25.	20.	17.	10.	7.	5.	10.
ANNUAL	1005.	506.	341.	119.	61.	32.	280.

Q2 - Q10 IS THE 1 IN 2 YEAR FLOW MINUS THE 1 IN 10 YEAR FLOW

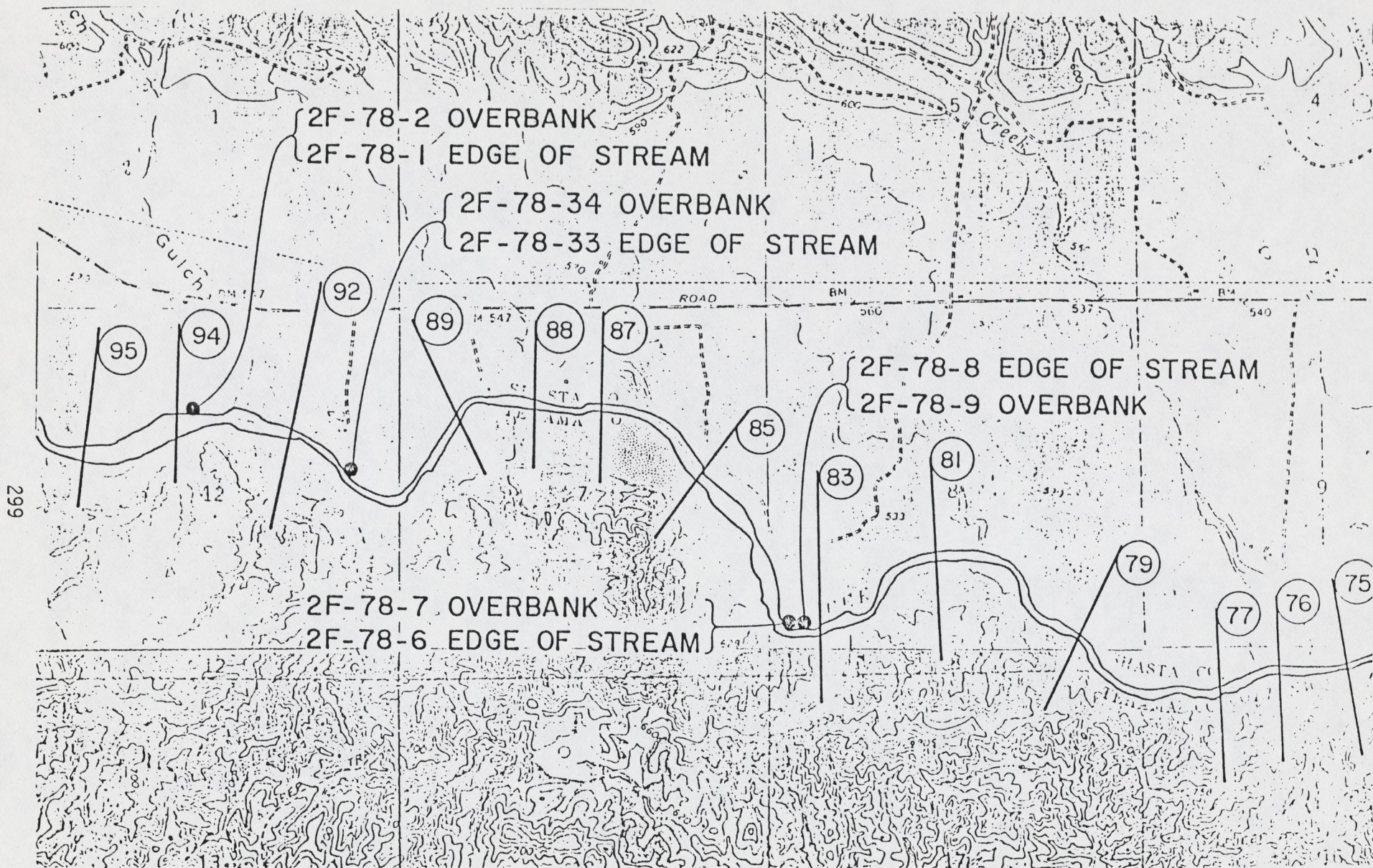
CROSS SECTIONS OF REACH OF INTEREST



Approximate location of cross sections and soil samples (scale 1:24,000 approx., April 1980).



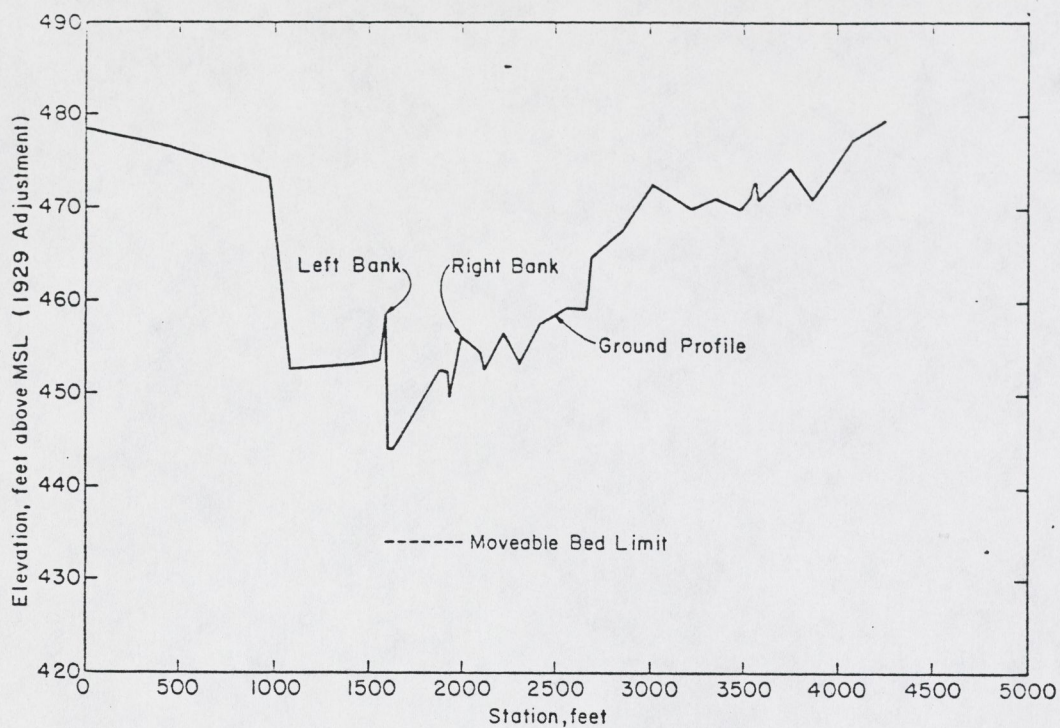
Approximate location of cross sections and soil samples (scale 1:24,000 approx., April 1980, continued).



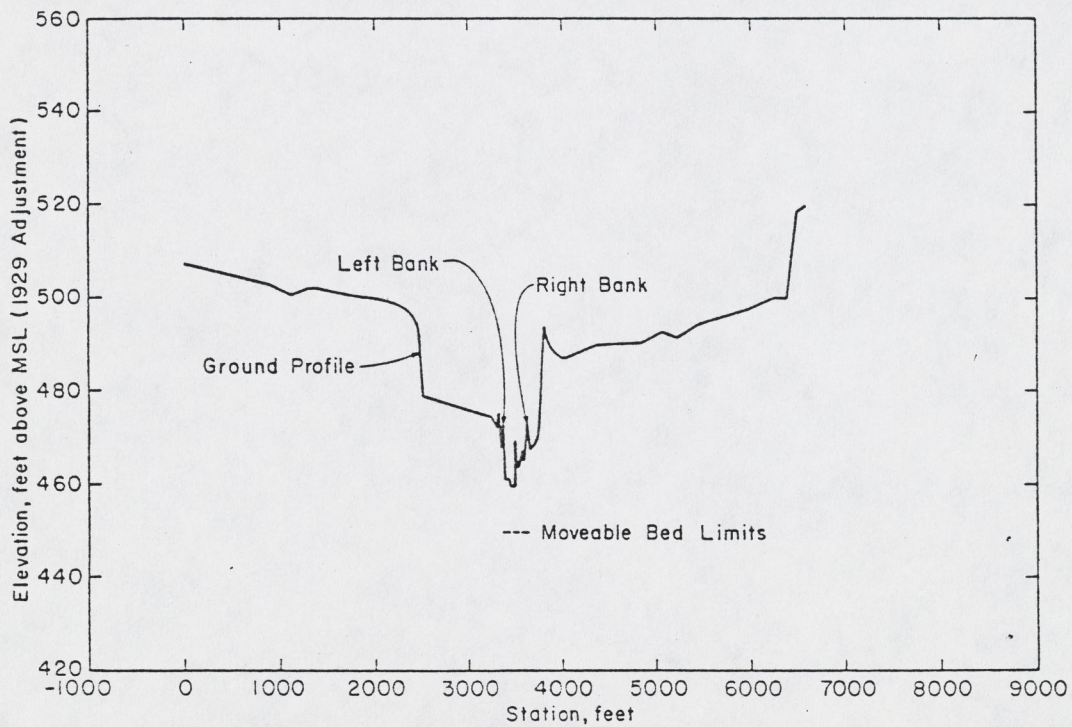
299

B-11

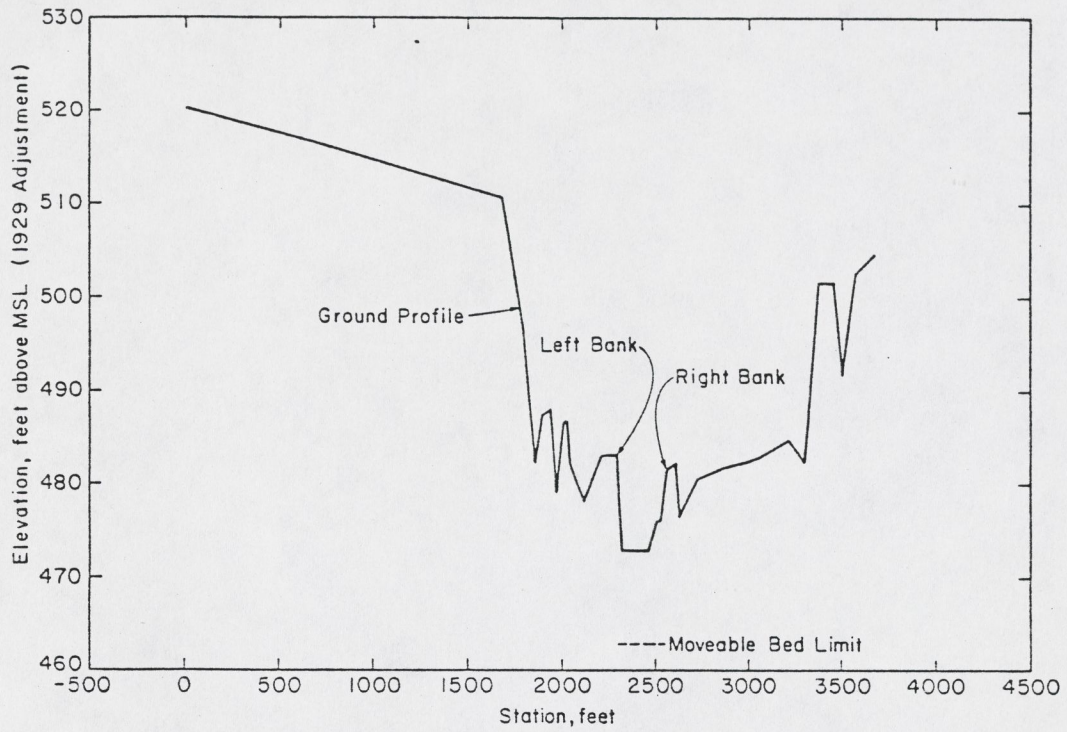
Approximate location of cross sections and soil samples (scale 1:24,000 approx., April 1980, continued).



Station compared to elevation for Section 57.



Station compared to elevation for Section 67.



Station compared to elevation for Section 72.

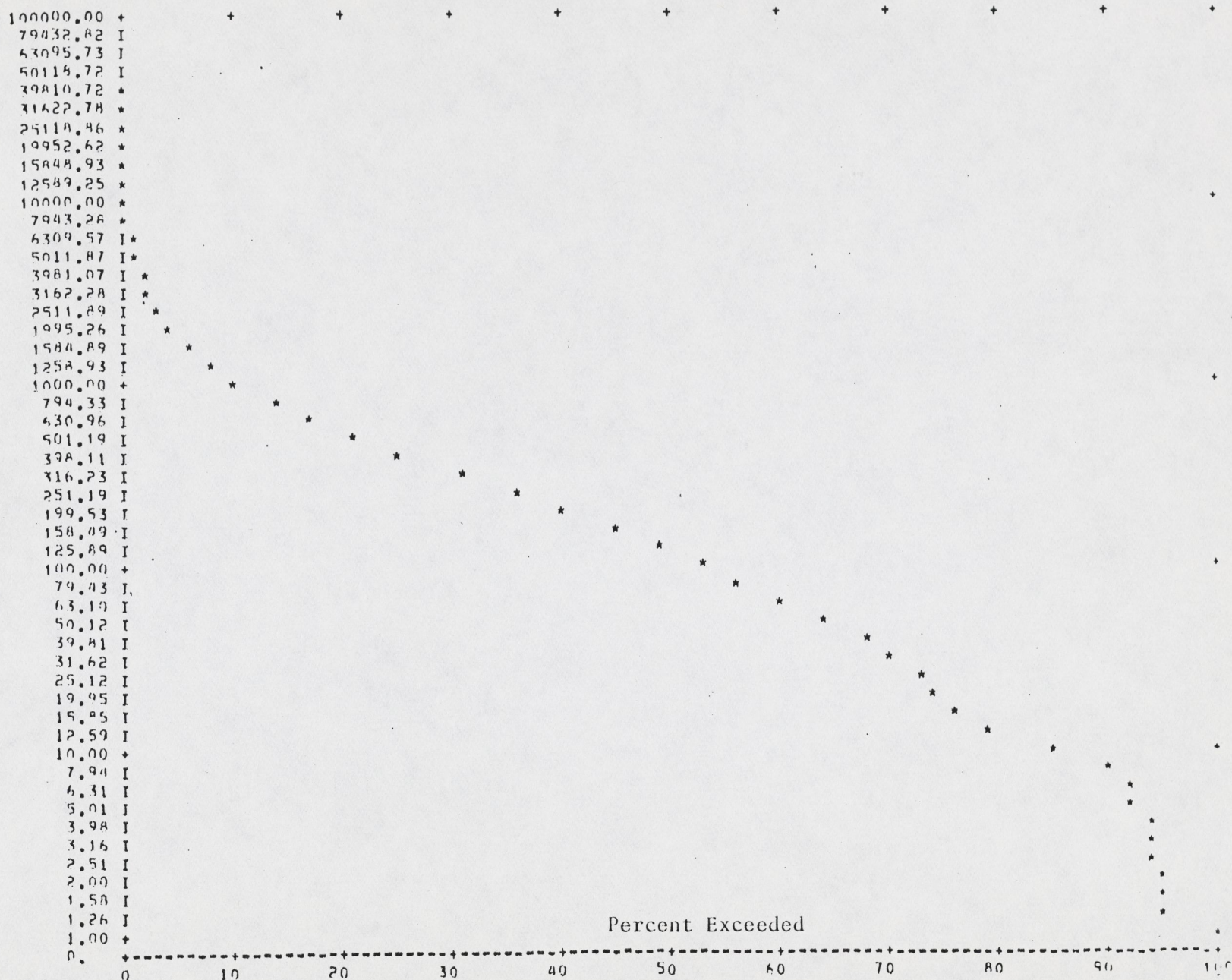
PEAK FLOWS AND FLOW DURATION CURVES

NATURAL CONDITIONS

JDA NUMBER 1

FLOW (CF3)

FLOW DURATION CURVE



303

B-15

NATURAL CONDITIONS

JOB NUMBER 1

STREAM FLOW DATA INPUT

FL CARD	FTYPE FILE	FMTYPE STD	PUNCH NO
FH CARD	TERM .999.00	ITVALS 1000000	NVALS 0

UNADJUSTED FLOW DURATION CURVE -- FLOWS IN CFS AND EXCEEDENCE PROBABILITIES AS DECIMAL FRACTIONS

PE	1.0000	.9467	.9467	.9467	.9416	.9365	.9365	.9238	.9162	.9035
QQ	1.00	1.26	1.58	2.00	2.51	3.16	3.98	5.01	6.31	7.94
PE	.8451	.7918	.7639	.7436	.7257	.6979	.6777	.6408	.6033	.5402
QQ	10.00	12.59	15.85	19.95	25.12	31.62	39.81	50.12	63.10	79.43
PE	.5271	.4931	.4472	.3975	.3553	.3110	.2504	.2098	.1750	.1162
QQ	100.00	125.89	158.49	199.53	251.19	316.23	398.11	501.19	630.96	794.33
PE	.1000	.0754	.0561	.0418	.0306	.0216	.0150	.0097	.0059	.0031
QQ	1000.00	1258.93	1584.89	1995.26	2511.89	3162.28	3981.07	5011.87	6309.57	7943.28
PE	.0015	.0006	.0003	.0002	.0001	.0000	0.			
QQ	10000.00	12589.25	15848.93	19952.62	25118.86	31622.78	39810.72			

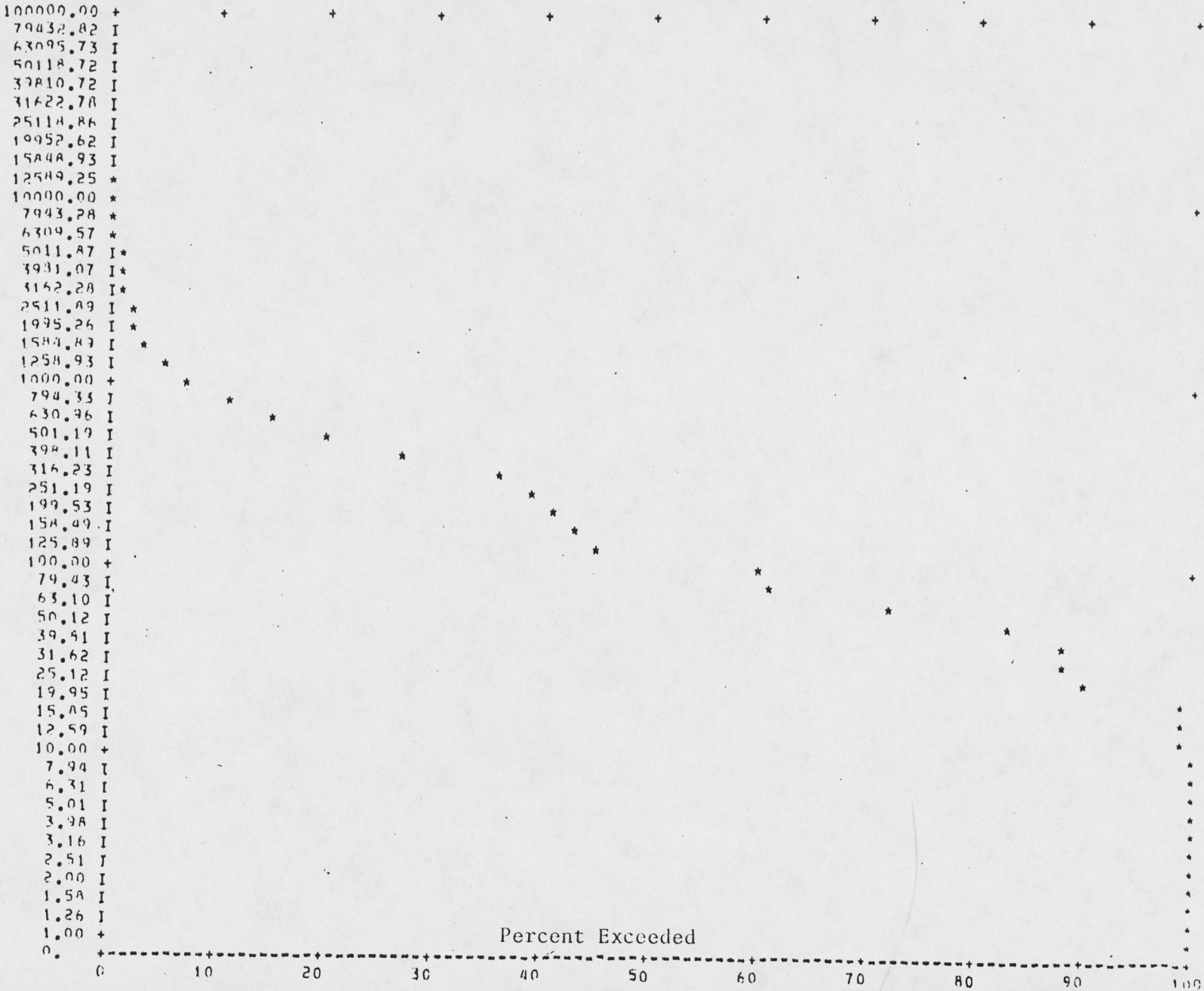
304

JOB NUMBER 1

PROJECT CONDITIONS

FLOW (CFS)

FLOW DURATION CURVE



305

B-17

JOB NUMBER 1

PROJECT CONDITIONS

STREAM FLOW DATA INPUT

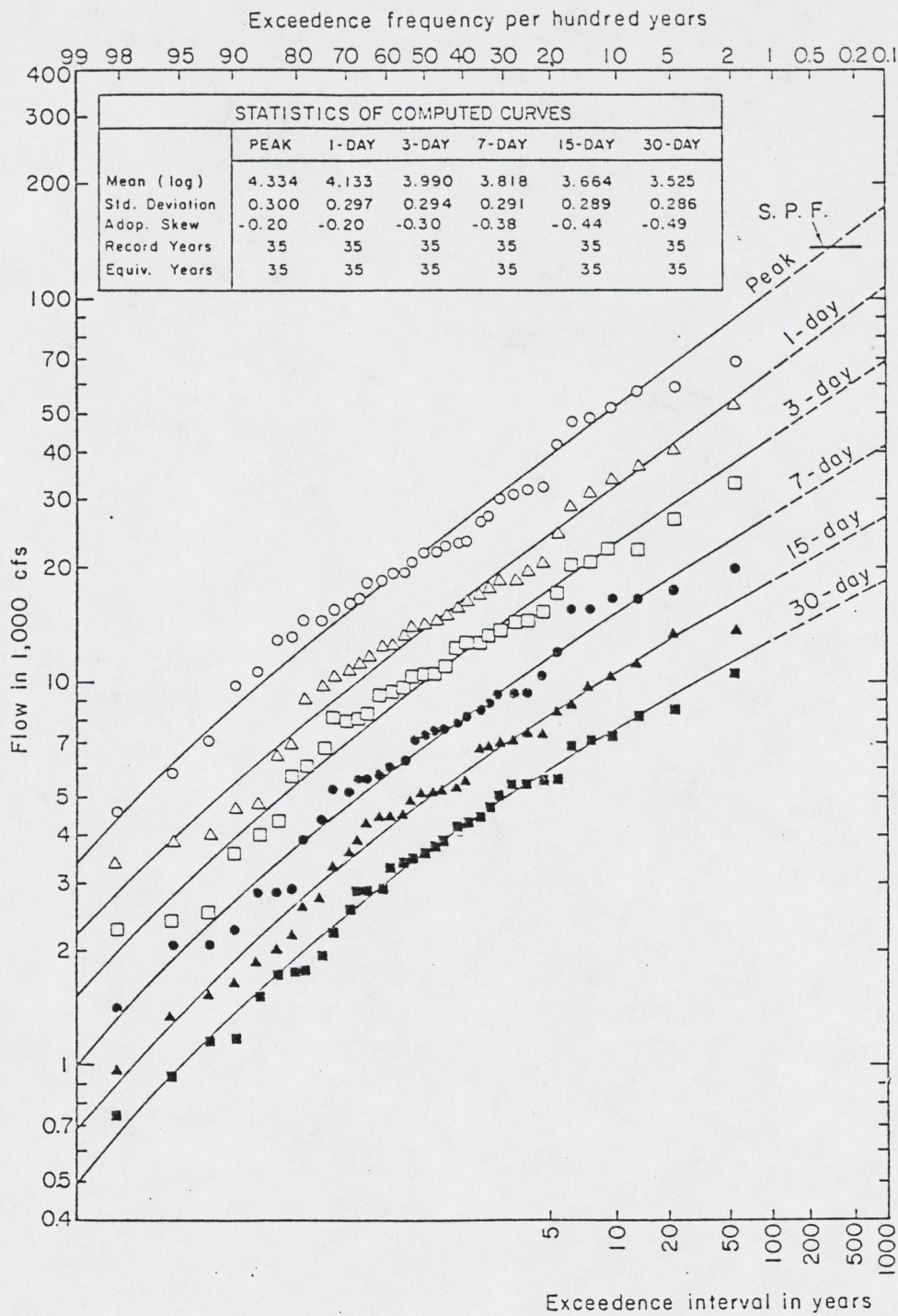
FL CARD	FTYPE FILE	FMTYPE STD	PUNCH NO
FH CARD	TERM -999.00	ITVALS 1000000	NYALS 0

UNADJUSTED FLOW DURATION CURVE -- FLOWS IN CFS AND EXCEEDENCE PROBABILITIES AS DECIMAL FRACTIONS

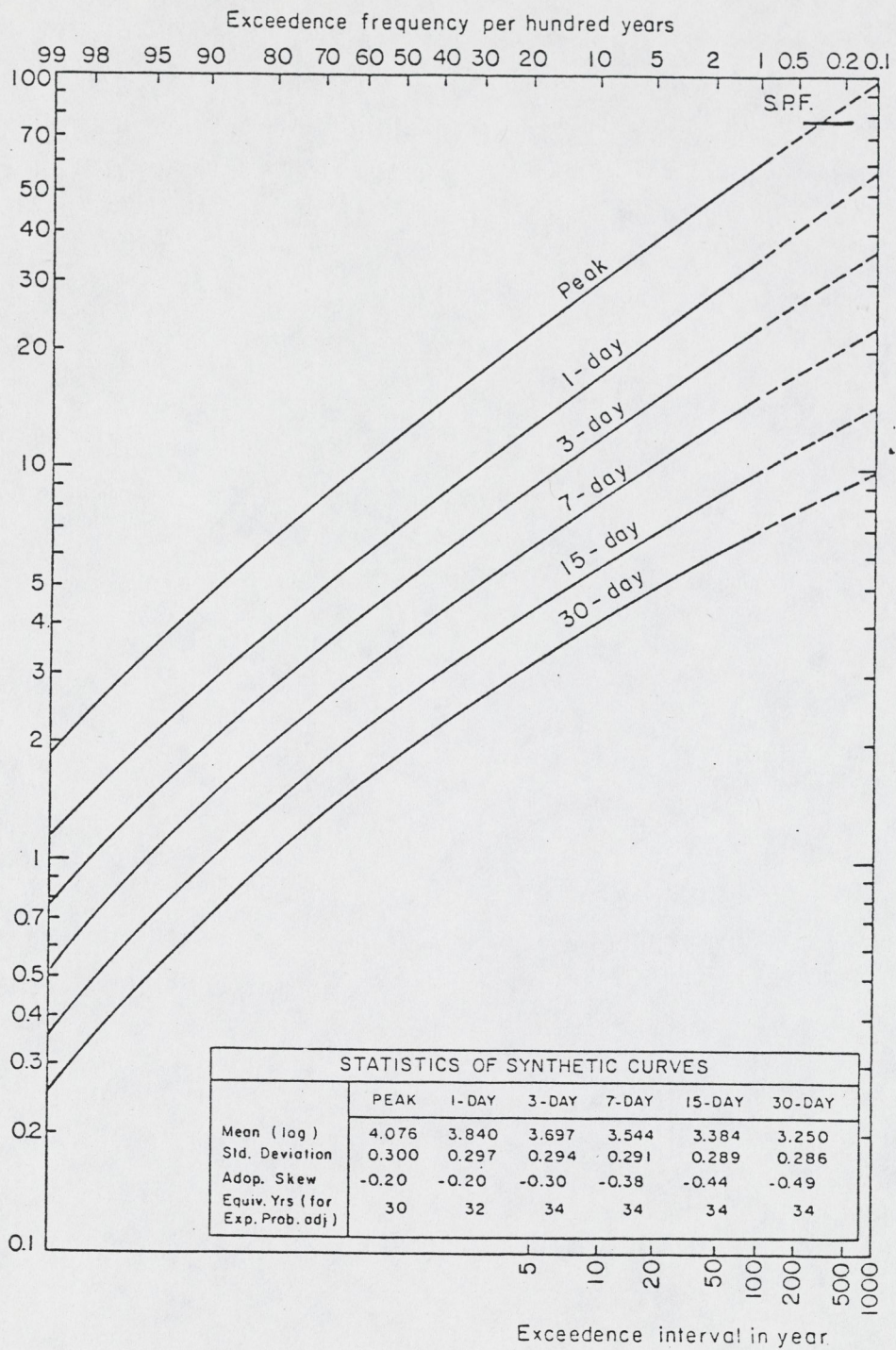
306

PE	1.0000	1.0000	1.0000	1.0000	.9975	.9975	.9975	.9975	.9975	.9975
QQ	1.00	1.26	1.58	2.00	2.51	3.16	3.98	5.01	6.31	7.94
PE	.9975	.9902	.9902	.9902	.8970	.8847	.8823	.8295	.7233	.6114
QQ	10.00	12.59	15.65	19.95	25.12	31.62	39.81	50.12	63.10	79.43
PE	.6032	.4534	.4349	.4123	.3941	.3567	.2672	.1994	.1538	.1053
QQ	100.00	125.89	158.49	199.53	251.19	316.23	398.11	501.19	630.96	794.33
PE	.0654	.0457	.0350	.0248	.0190	.0145	.0108	.0075	.0049	.0032
QQ	1000.00	1258.93	1584.89	1995.26	2511.89	3162.28	3981.07	5011.67	6309.57	7943.28
PE	.0007	0.								
QQ	10000.00	12589.25								

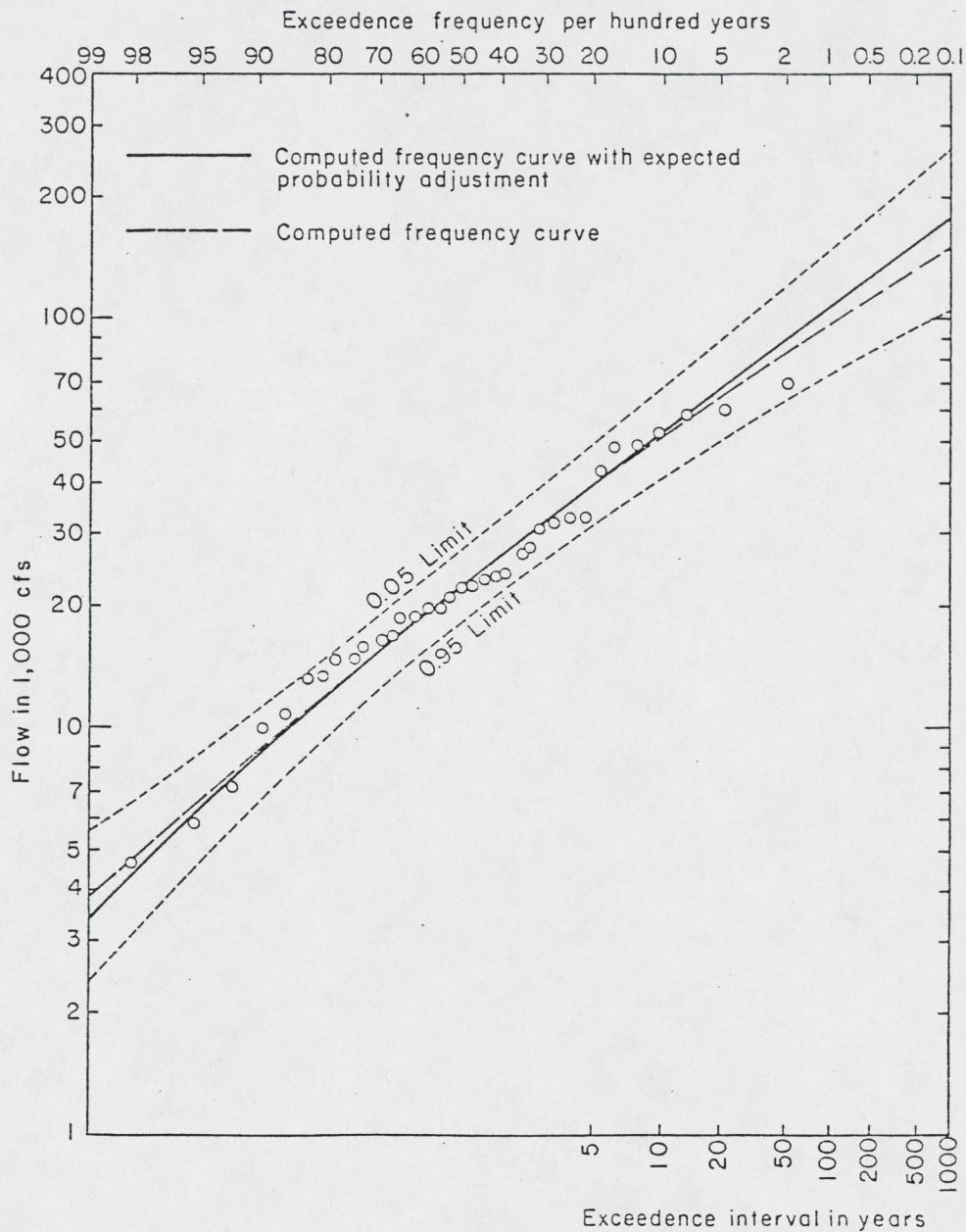
B-1-8



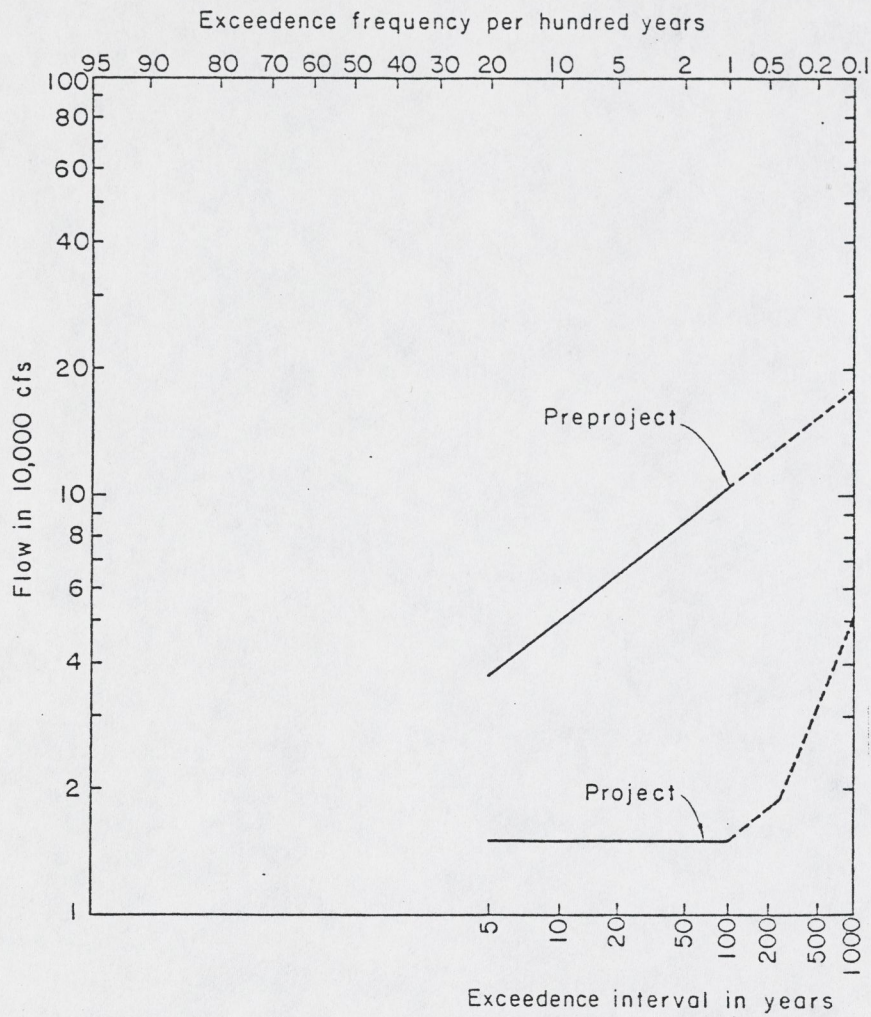
Flow frequency Poplar Creek near Site 1 (Corps of Engineers, April 1977).
 Drainage area: 927 sq. mi. Period of record: 1941-1975.



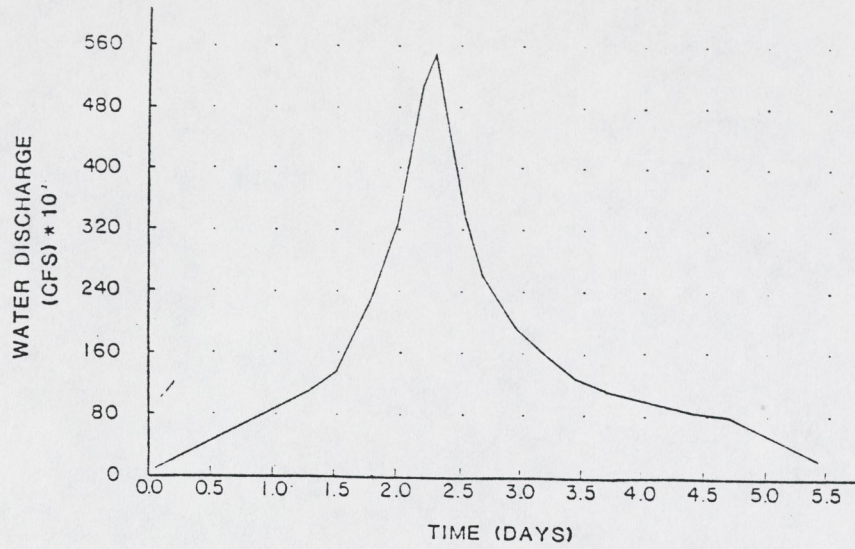
Flow frequency Poplar Creek at dam site above Site 1 (Corps of Engineers, April 1977).
 Drainage area: 394.2 sq. mi.



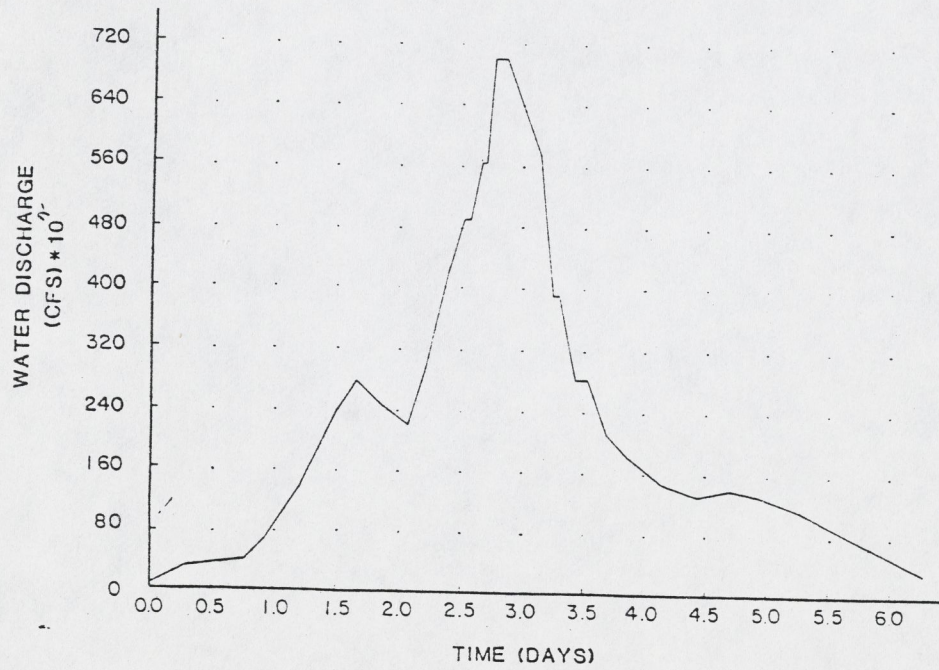
Peak flow frequency Poplar Creek at Site 3 (Corps of Engineers, April 1977).
 Drainage area: 927.0 sq. mi. Period of record: 1941-1975.



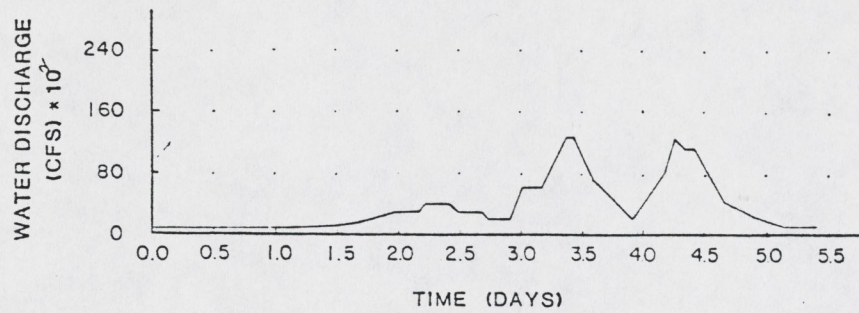
Peak flow frequency Poplar Creek near Site 1 for preproject and postproject conditions (Corps of Engineers, July 1969).



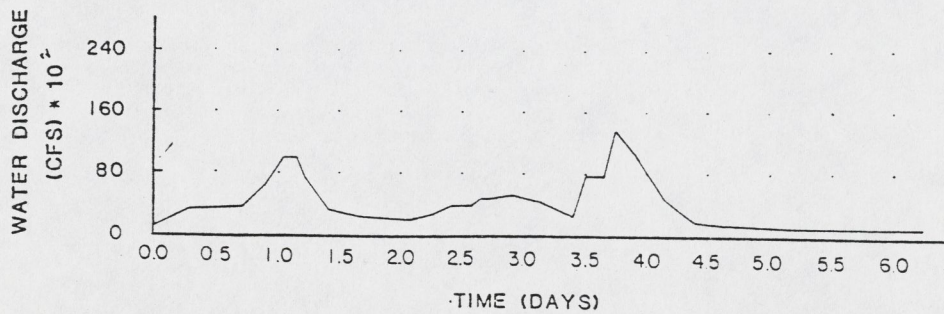
Poplar Creek preproject flood (1970).



Poplar Creek preproject flood (1974).



Poplar Creek project flood (1970).



Poplar Creek project flood (1974).

BED MATERIAL

BED MATERIAL

Total Depth of Hole: 1.0 ft each
 Date Started: 18 July 78
 Date Completed: 19 July 78

Size and Type of Bit: Shovel
 Manufacturer's Designation of Drill:
 Ames Middleweight #2

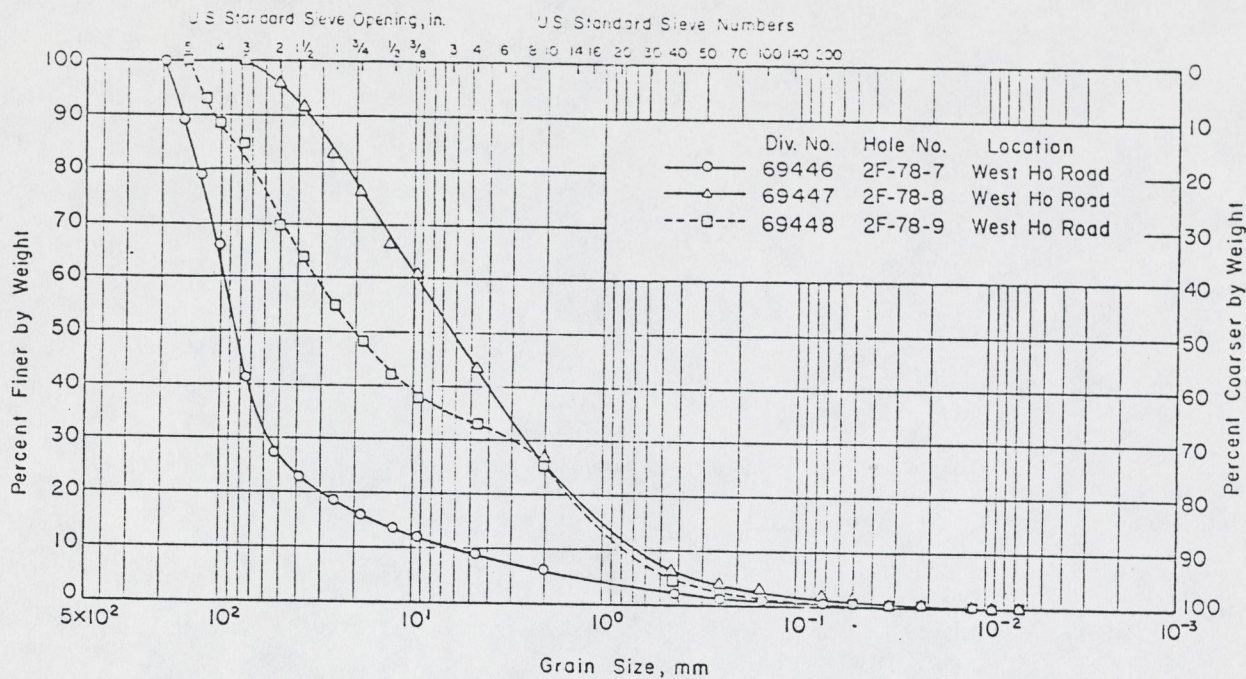
DEPTH	CLASSIFICATION OF MATERIALS	TEST HOLE NUMBERS
0.0 0.5 1.0	sandy gravel, 75% rounded gravel, 25% fine to coarse subangular sand, to 4" maximum.	2F-78-8
0.0 0.5 1.0	sandy gravel with scattered cobbles. 75% rounded gravel, 25% fine to coarse subangular sand, to 5" maximum.	2F-78-9
0.0 0.5 1.0	cobbly sandy gravel, 60% rounded gravel, 30% fine to coarse angular sand, 10% rounded cobbles.	2F-78-10
0.0 0.5 1.0	sandy gravel with cobbles, 90% rounded cobbles 10% fine-coarse, angular sand, cobbles to 6" maximum.	2F-78-11
0.0 0.5 1.0	sandy, gravelly cobbles, 40% rounded cobbles, 30% rounded gravels, 30% fine to coarse angular sand.	2F-78-12
0.0 0.5 1.0	sandy gravel with scattered cobbles, 90% rounded gravels, 10% fine-coarse angular sand, Cobbles to maximum 6".	2F-78-13
	19 July 78	
0.0 0.5 1.0	Sandy gravel, 65% gravel, 35% sand, Gravel maximum dimension 5" but would go through 3" square.	2F-78-14

Drilling log for Poplar Creek sedimentation study

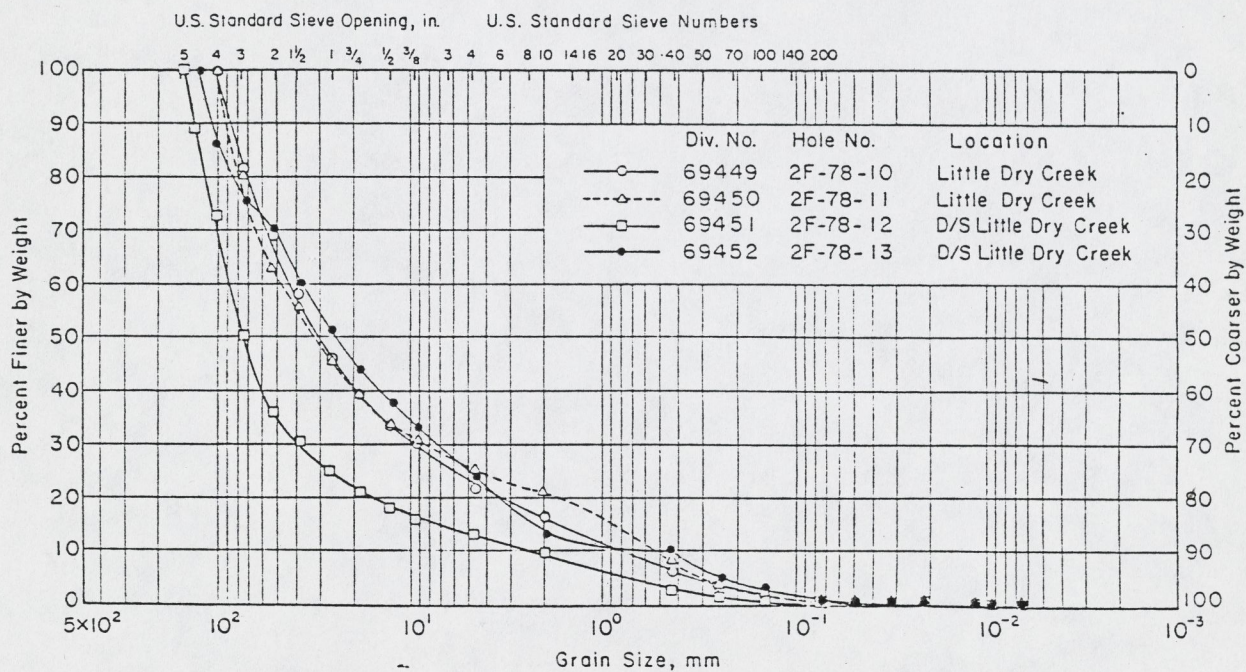
MECHANICAL ANALYSIS--% FINER

DIVISION SERIAL NO.	HOLE NO.	CO-ORD OR STA.	Gravel				Sand				Fines		
			6/3	2/1 1/2	1/3/4	1/2 3/8	No. 4	No. 10	No. 40	No. 80	No. 200		
			69440	2F-78-1	Dutch Gulch	100	98/97	94/91	88/84	74	57		19
69441	2F-78-2	Dutch Gulch		100/96	94/86	76/67	52	37	8	3	0.5		2 Sacks
69442	2F-78-3	Dutch Gulch Bridge	82/58	45/40	32/27	22/20	16	9	3	2	1	42% Cobbles	4 Sacks
69443	2F-78-4	Dutch Gulch Bridge	94/80	74/72	68/66	64/63	62	50	8	5	4	20% Cobbles	3 Sacks
69444	2F-78-5	Dutch Gulch Bridge	100	97/91	79/73	65/61	52	35	3	2	1		2 Sacks
69445	2F-78-6	West Ho Road	91/64	55/49	41/35	30/27	21	16	5	4	2	36% Cobbles	4 Sacks
69446	2F-78-7	West Ho Road	89/41	27/23	18/15	13/12	9	6	2	2	1	59% Cobbles	4 Sacks
69447	2F-78-8	West Ho Road	100	97/93	83/76	68/61	43	27	8	5	1		2 Sacks
69448	2F-78-9	West Ho Road	100/85	71/65	56/48	42/39	33	26	5	3	1	15% Cobbles	2 Sacks
69449	2F-78-10	Little Dry Cr.	100/82	69/58	46/40	34/30	22	17	6	4	1	18% Cobbles	3 Sacks
69450	2F-78-11	Little Dry Cr.	100/80	63/56	46/40	35/31	26	22	9	4	1	20% Cobbles	4 Sacks
69451	2F-78-12	D/S Little Dry Cr.	100/50	36/31	25/21	18/16	13	10	3	2	1	50% Cobbles	4 Sacks
69452	2F-78-13	D/S Little Dry Cr.	100/75	70/60	51/45	38/33	23	13	10	5	1	25% Cobbles	2 Sacks
69453	2F-78-14	Confluence of N&S Forks	100/98	90/79	64/52	43/37	29	21	7	3	0.5	2% Cobbles	2 Sacks
69454	2F-78-15	Confluence of N&S Forks	80/41	31/28	23/19	16/14	10	7	3	1	0	59% Cobbles	4 Sacks
69455	2F-78-16	Confluence of N&S Forks		100/93	82/72	62/54	39	29	9	6	2		2 Sacks
69456	2F-78-17	Confluence of N&S Forks	100/78	73/70	62/56	48/42	27	14	6	3	2	22% Cobbles	2 Sacks

Soil test result summary for Poplar Creek, August 1978.



Sedimentation study (Poplar Creek, 2F-78-7 to 2f-78-9).



Sedimentation study (Poplar Creek, 2f-78-10 to 2F-78-13).

APPENDIX C

CHANNEL CHANGE WORKSHOP: PROBLEM NUMBER 3
ELK RIVER, KANSAS

THE QUESTION

The question relative to the Elk River is "What will happen to the morphology of the stream channel as a result of the changes in stream-flows and sediment discharge caused by the construction of a number of small flood retaining reservoirs upstream of a reach of stream?" Specifically, for both a short and long time after construction of the reservoirs:

1. What will be the meander pattern?,
2. What will be the configuration of the channel?,
3. What will be the substrate material?, and
4. What will be the pool riffle sequence?

Elk River is located in southeastern Kansas in an area with sub-humid climate. The materials in this appendix describe the basin and present the data on both pre and post project conditions. The soils data is for Chautauqua County, the county immediately below Elk County. The Elk River is located in the southern part of Elk County.

The reach of interest is the reach of the river near Longton, Kansas; specifically between Cross Sections 11-15 and 3-3 as shown on the enclosed plan.

Most of the data enclosed were obtained from the Kansas State Office of the Soil Conservation Service. Other data were obtained from U.S. Geological Survey reports. The geology data is from a Kansas Geological Survey report.

CONTENTS

- General Location of Elk River, Kansas
- General Plan - Elk River Watershed Joint District No. 47 (Partial)¹
- Soil Survey of Chautauqua County, Kansas (Partial)¹
- Geology, Mineral Resources and Groundwater Resources of Elk County, Kansas (Partial)¹
- Plan, Profile, and Cross Section Diagrams for the Longton Reach of the Elk River, Kansas²
- Monthly Streamflow Data for the Elk River Basin, Kansas

¹These items are not included in the information supplied in this appendix. The following bibliography lists the sources of these data.

²Only part of the diagrams included in the original data set are supplied and included in this appendix. The others may be reviewed at the office of the Instream Flow Group in Fort Collins, Colorado.

Elk River Bed Material
Elk River Sediment Yield
Frequency - Runoff Curve
Elk River Hydrographs and Elevation vs. Discharge Data
 Unit Discharge Hydrograph
 Unit Discharge Hydrograph Table
 Elevation - Discharge Plots
 Elevation - Discharge Table

USGS MAPS SHOWING THE ELK RIVER AREA INCLUDED IN THE ORIGINAL DATA SETS

1:250,000

Joplin, Missouri; Kansas
Wichita, Kansas

1:24,000

Longton NW, Kansas
Elk Falls, Kansas
Longton, Kansas
Oak Valley, Kansas

BIBLIOGRAPHY OF REPORTS USED IN DATA SET

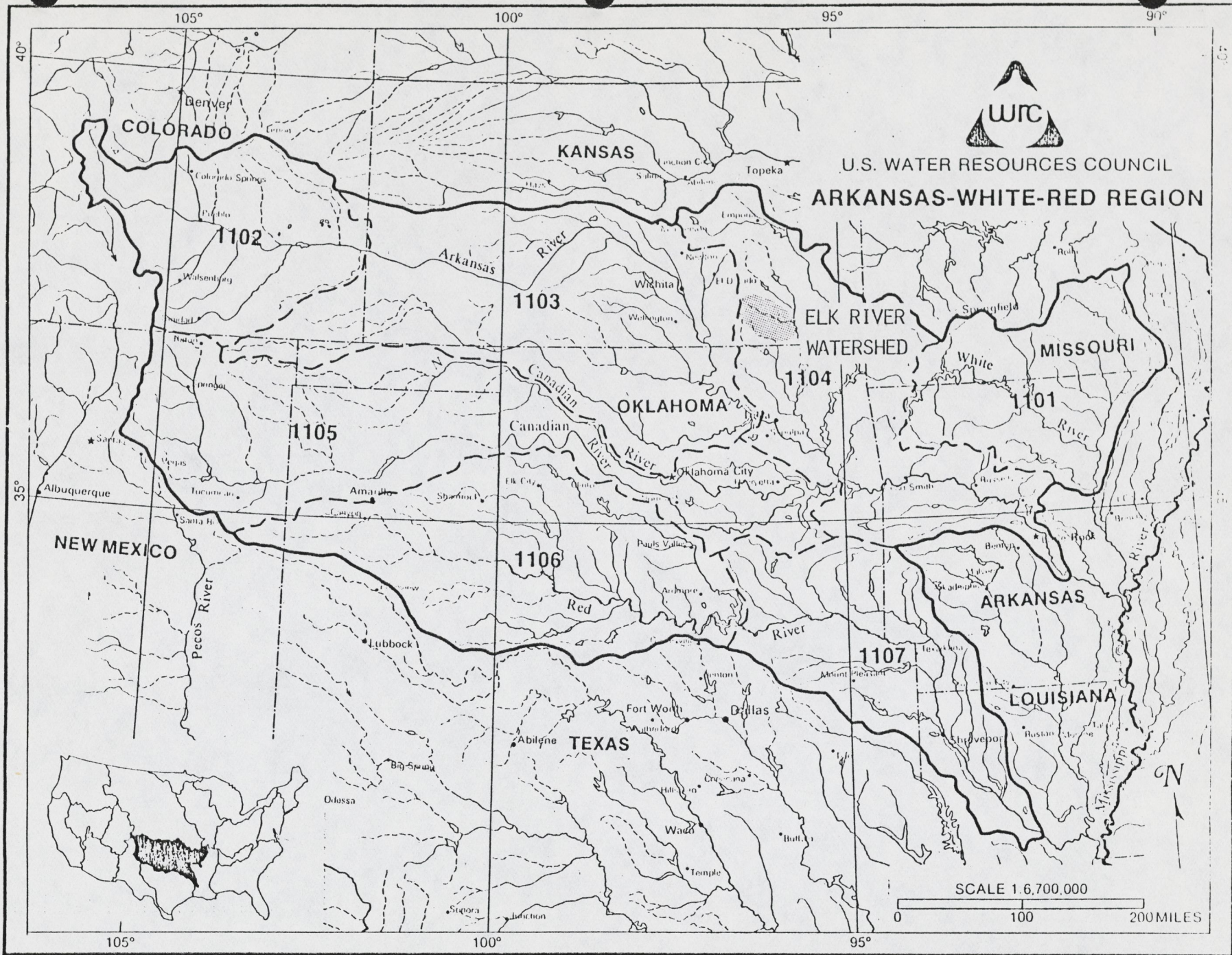
General Plan - Elk River Watershed Joint District No. 47, Kansas State Office, Soil Conservation Service, USDA, Salina, Kansas, January 1967.

Bell, E. L., and H. T. Rowland. Soil Survey of Chautauqua County, Kansas, Soil Conservation Service, USDA, Washington, D.C., October 1976.

Verville, G. J. R., Kulstad, N. Plummer, W. H. Schoewa, E. D. Goebel, and C. K. Bayhe. Geology, Mineral Resources, and Ground-Water Resources of Elk County, Kansas, State Geological Survey of Kansas (Volume 14), Lawrence, Kansas, July 1958.

PLAN, PROFILE, AND CROSS SECTION DIAGRAMS FOR THE LONGTON REACH OF THE ELK RIVER, KANSAS

The following figures illustrate the reach of the Elk River for which an estimate of the future form of the channel and substrate is needed. The reach is between Sections 3-3 and 11-15. Only a few of the diagrams are included in this appendix.



General location of Elk River, Kansas.

321

C-3

DIAGRAMS INCLUDED IN ORIGINAL DATA SETS

Plan-Profile Reach Numbers 3 and 11

Plan-Profile Reach Numbers 11 and 13

Cross Section 3-3
3-4
3-5
11-1 and 11-2
11-3 and 11-4
11-5 and 11-7
11-8 and 11-9
11-14 and 11-15

MONTHLY STREAMFLOW DATA FOR THE ELK RIVER BASIN, KANSAS

Drainage Areas

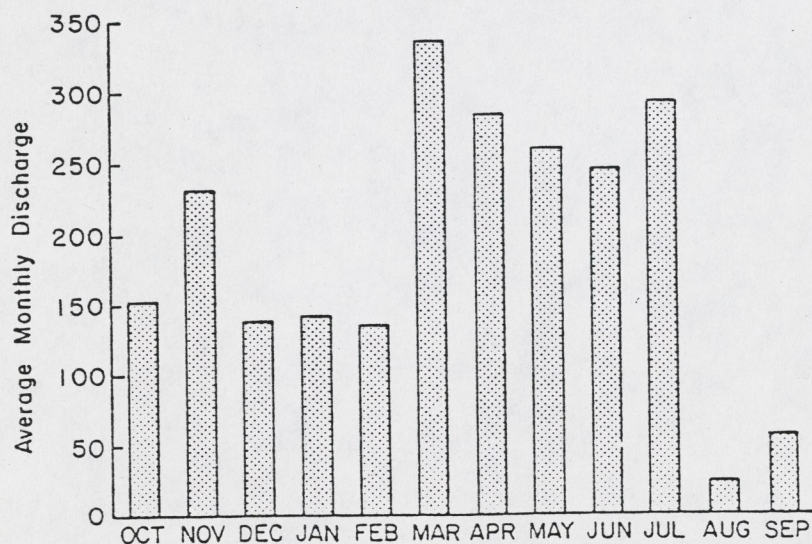
Study Reach:

Head of Reach: 285.7 square miles
Bottom of Reach: 405.3 square miles

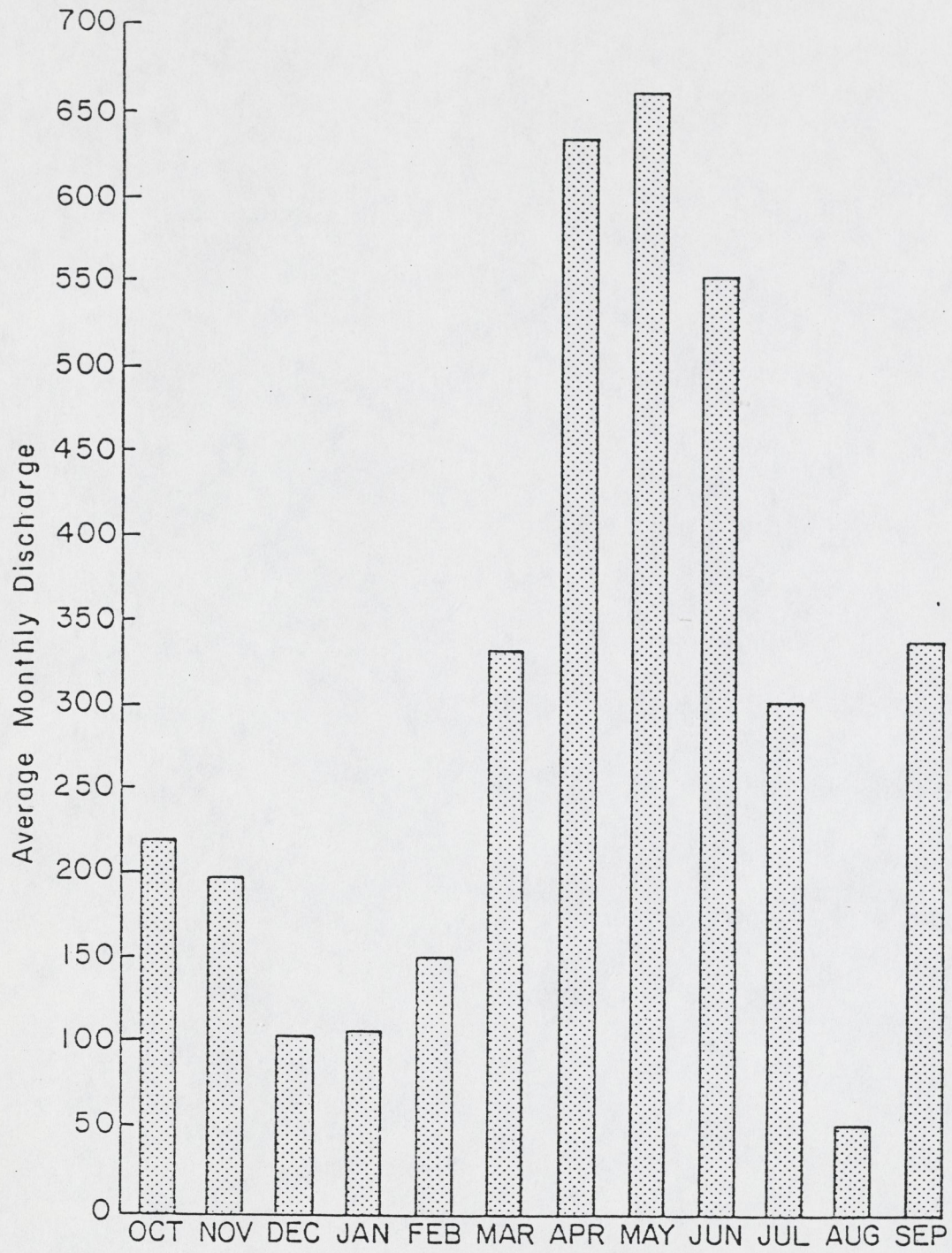
Gaging Station:

07-1698 Elk River at Elk Falls: 220 square miles
07-1700 Elk River near Elk City: 575 square miles

Included in the original data set were data from the U.S. Geological Survey report "Water Resources Data for Kansas." The data included were from water year 1978 and were on pages 15 and 297. Most of the data are from records of the U.S. Geological Survey.



Elk River at Elk Falls, Kansas (1968-1976)



Elk River near Elk City, Kansas (1939-1969)

ELK RIVER AT ELK FALLS, KANSAS
 ARKANSAS RIVER BASIN
 UNITS OF DISCHARGE ARE CUSECS

YEAR	OCT	NOV	DEC	JAN	FEB	MAR	APR	MAY	JUNE	JULY	AUG	SEPT	ANNUAL
1968	234.00	104.00	45.40	31.50	29.20	22.80	131.00	311.00	75.90	25.60	32.60	14.20	88.62
1969	197.00	390.00	143.00	84.00	158.00	332.00	444.00	614.00	758.00	34.20	2.29	169.00	219.23
1970	288.00	62.30	74.90	30.40	19.30	58.00	992.00	62.80	201.00	5.59	.25	13.90	149.87
1971	31.10	6.52	4.35	117.00	126.00	67.10	23.90	60.80	74.30	65.80	1.24	.33	47.86
1972	12.70	5.32	173.00	39.30	22.60	14.20	24.50	52.50	2.65	389.00	10.90	27.60	65.39
1973	19.40	394.00	229.00	394.00	265.00	1247.00	436.00	108.00	16.00	3.70	.77	128.00	219.13
1974	207.00	199.00	364.00	261.00	171.00	802.00	154.00	543.00	464.00	6.35	153.00	143.00	219.37
1975	389.00	930.00	215.00	309.00	426.00	477.00	140.00	354.00	508.00	20.00	12.70	11.70	314.31
1976	1.26	2.32	5.30	2.54	1.80	7.18	212.00	147.00	19.90	2080.00	5.73	2.76	210.48
AVERAGE	153.27	233.05	139.33	140.97	135.43	336.36	284.16	250.34	235.53	292.25	24.39	56.72	190.43
Q/QANN	6.831	10.052	6.210	6.283	5.501	14.992	12.256	11.158	10.159	13.025	1.087	2.446	100.000
COV VAR	.925	1.306	.859	1.016	1.039	1.299	1.078	.863	1.164	2.332	2.021	1.211	.549
SKEW	.327	1.733	.616	.880	1.163	1.384	1.776	.786	1.040	2.831	2.773	.899	-.209
MAXIMUM	389.00	930.00	364.00	394.00	426.00	1247.00	992.00	614.00	758.00	2080.00	153.00	169.00	314.31
MINIMUM	1.26	2.32	4.35	2.54	1.80	7.18	23.90	52.50	2.65	3.70	.25	.33	47.86

Monthly streamflow data.

STATISTICAL PARAMETERS FOR STATION 07169800 ELK RIVER AT ELK FALLS, KANSAS
 ARKANSAS RIVER BASIN

	MONTH	ARTH. AVERAGE	MEAN	LOG PARAMETERS		SKEW
				VARIANCE	STD. DEV.	
1	OCT	153.27	1.7793	.7025	.8331	-1.0480
2	NOV	233.05	1.7973	.8888	.9427	-.4516
3	DEC	139.33	1.8379	.5017	.7083	-1.0740
4	JAN	140.97	1.8306	.4686	.6845	-1.0222
5	FEB	135.43	1.7717	.5584	.7473	-.9952
6	MAR	336.36	2.0174	.6591	.8119	-.0464
7	APR	284.16	2.2031	.2982	.5461	-.4234
8	MAY	250.34	2.2297	.1782	.4221	.0952
9	JUNE	235.53	1.9147	.6731	.8204	-.5918
10	JULY	292.25	1.5653	.8158	.9032	.9755
11	AUG	24.39	.7039	.7483	.8651	.2077
12	SEPT	56.72	1.2505	.7817	.8842	-.7961
13	ANNUAL	190.43	2.1973	.0960	.3098	-.7273

SAMPLE SIZE 9 YEARS

LOG NORMAL DISTRIBUTION STATION 07169800 ELK RIVER AT ELK FALLS, KANSAS
 ARKANSAS RIVER BASIN

YEARS FLOW IS NOT EXCEEDED

MONTH	9 IN 10	2 IN 3	1 IN 2	1 IN 5	1 IN 10	1 IN 20	Q2 - Q10
OCT	714.17	137.68	60.16	11.85	5.07	2.52	55.09
NOV	1013.60	159.11	62.70	10.08	3.88	1.76	58.82
DEC	557.13	138.60	68.85	17.44	8.51	4.71	60.34
JAN	510.61	133.11	67.69	17.96	8.97	5.06	58.72
FEB	536.63	123.67	59.11	13.88	6.51	3.49	52.60
MAR	1143.32	232.08	104.08	21.57	9.47	4.81	94.60
APR	800.31	273.79	159.64	55.37	31.84	20.17	127.79
MAY	590.00	257.52	169.72	74.87	48.82	34.31	120.90
JUNE	925.87	184.80	82.18	16.75	7.29	3.67	74.88
JULY	528.72	89.70	36.76	6.38	2.56	1.20	34.20
AUG	65.00	11.88	5.06	.95	.39	.19	4.66
SEPT	242.08	42.64	17.80	3.21	1.31	.63	16.49
ANNUAL	393.05	213.90	157.52	86.40	63.13	48.73	94.39

Q2 - Q10 IS THE 1 IN 2 YEAR FLOW MINUS THE 1 IN 10 YEAR FLOW

LOG-PEARSON TYPE III DISTRIBUTION FOR STATION 07169800
 ELK RIVER AT ELK FALLS, KANSAS
 ARKANSAS RIVER BASIN

DATA SKEW K FACTORS

MONTH	SKEW	RETURN PERIOD					
		10 YRS	6.67 YRS	2 YRS	1.25 YRS	1.11 YRS	1.05 YRS
OCT	-1.05	1.0961	.5234	.1878	-.7382	-1.3405	-1.9023
NOV	-.45	1.2083	.4813	.0907	-.8041	-1.3254	-1.7851
DEC	-1.07	1.1015	.5221	.1839	-.7416	-1.3407	-1.8982
JAN	-1.02	1.0907	.5248	.1917	-.7349	-1.3402	-1.9064
FEB	-1.00	1.1270	.5152	.1648	-.7574	-1.3400	-1.8776
MAR	-.05	1.2636	.4451	.0256	-.8328	-1.2968	-1.6875
APR	-.42	1.2037	.4836	.0953	-.8019	-1.3268	-1.7916
MAY	.10	1.2706	.4393	.0162	-.8363	-1.2915	-1.6717
JUNE	-.59	1.1986	.4863	.1004	-.7992	-1.3284	-1.7988
JULY	.92	1.3367	.3298	-.1359	-.8555	-1.1613	-1.3744
AUG	.21	1.3003	.4080	-.0318	-.8497	-1.2589	-1.5883
SEPT	-.80	1.1653	.5015	.1326	-.7796	-1.3361	-1.8397
ANNUAL	-.73	1.1522	.5062	.1436	-.7720	-1.3382	-1.8528

FREQUENCY OF FLOWS

YEARS FLOW IS NOT EXCEEDED

MONTH	9 IN 10	2 IN 3	1 IN 2	1 IN 5	1 IN 10	1 IN 20	Q2 - Q10
OCT	498.85	165.21	86.44	14.47	4.53	1.53	81.91
NOV	863.65	178.23	76.35	10.94	3.53	1.30	72.82
DEC	415.09	161.32	92.93	20.54	7.73	3.11	85.20
JAN	377.68	154.81	91.57	21.26	8.19	3.35	83.38
FEB	411.00	143.43	78.49	16.06	5.89	2.34	72.59
MAR	1104.59	239.17	109.17	21.94	9.22	4.44	99.96
APR	725.30	293.26	179.95	58.24	30.10	16.78	149.86
MAY	583.49	260.10	172.41	75.29	48.37	33.43	124.04
JUNE	790.91	205.92	99.34	18.16	6.68	2.75	92.66
JULY	592.46	72.98	27.71	6.20	3.28	2.09	24.42
AUG	67.41	11.40	4.75	.93	.41	.21	4.33
SEPT	190.87	49.42	23.32	3.64	1.17	.42	22.15
ANNUAL	358.29	226.01	174.51	90.82	60.65	42.01	113.86

Q2 - Q10 IS THE 1 IN 2 YEAR FLOW MINUS THE 1 IN 10 YEAR FLOW

ELK RIVER NEAR ELK CITY, KANSAS
 ARKANSAS RIVER BASIN
 UNITS OF DISCHARGE ARE CUSECS

YEAR	OCT	NOV	DEC	JAN	FEB	MAR	APR	MAY	JUNE	JULY	AUG	SEPT	ANNUAL
1939	2.00	9.40	5.70	8.10	9.80	35.50	260.00	350.00	195.00	6.00	1.80	.10	73.63
1940	0.00	0.00	0.00	0.00	.40	.70	55.30	101.00	74.30	1.30	29.70	27.80	24.14
1941	.90	118.00	55.00	296.00	253.00	69.20	1181.00	158.00	892.00	10.10	7.10	584.00	298.10
1942	1424.00	645.00	208.00	83.70	380.00	181.00	1293.00	268.00	1284.00	48.10	154.00	1564.00	623.37
1943	408.00	106.00	496.00	206.00	212.00	181.00	85.40	3692.00	1373.00	68.30	12.00	3.90	574.95
1944	14.10	2.70	13.10	11.40	43.40	1008.00	3400.00	785.00	186.00	23.50	45.90	226.00	478.02
1945	538.00	20.10	778.00	74.70	73.10	1368.00	2556.00	220.00	501.00	564.00	11.00	2740.00	785.05
1946	802.00	40.30	29.70	516.00	340.00	498.00	248.00	58.70	23.10	6.41	27.60	27.50	218.65
1947	1.67	150.00	46.90	24.60	11.90	403.00	3370.00	1374.00	275.00	54.10	7.89	6.19	475.42
1948	.74	.53	1.94	2.19	2.74	156.00	216.00	112.00	1330.00	3096.00	105.00	9.38	422.84
1949	2.91	102.00	10.80	770.00	1629.00	493.00	673.00	918.00	1575.00	417.00	51.40	182.00	560.00
1950	30.30	20.00	22.00	55.60	29.50	39.60	19.60	74.80	682.00	644.00	495.00	110.00	186.12
1951	11.80	7.14	8.87	9.30	70.60	62.90	220.00	1237.00	2721.00	2461.00	51.80	415.00	607.33
1952	68.30	603.00	164.00	210.00	164.00	850.00	585.00	95.50	49.30	6.57	3.17	.05	232.93
1953	0.00	0.00	1.14	1.04	76.00	18.90	4.20	96.40	2.14	3.24	0.00	0.00	16.65
1954	.60	13.40	6.16	.27	4.76	2.69	264.00	608.00	21.20	.62	9.73	1.35	73.31
1955	72.20	.62	.07	2.80	13.90	3.45	4.50	625.00	149.00	18.39	.80	0.00	74.22
1956	91.90	0.00	0.00	0.00	0.00	0.00	0.00	4.36	25.60	35.90	.05	0.00	13.32
1957	0.00	0.00	0.00	0.00	.83	1.83	156.00	1562.00	2322.00	68.80	34.00	40.20	343.35
1958	16.20	88.60	20.00	49.00	88.90	1885.00	756.00	514.00	75.50	440.00	19.70	259.00	353.58
1959	13.90	26.50	12.90	20.40	30.20	62.70	292.00	452.00	40.30	767.00	79.40	25.30	153.41
1960	1754.00	78.40	52.30	172.00	393.00	774.00	437.00	405.00	107.00	53.40	67.90	7.85	360.43
1961	29.50	13.80	79.10	10.10	54.70	312.00	721.00	4773.00	193.00	322.00	21.60	2874.00	787.69
1962	737.00	2085.00	462.00	481.00	202.00	286.00	85.80	26.30	86.20	84.30	3.49	497.00	418.37
1963	155.00	52.80	56.20	296.00	48.70	291.00	43.30	82.30	58.30	2.65	4.24	.19	91.78
1964	0.00	0.00	0.00	0.00	.38	.46	11.10	54.30	133.00	1.55	43.40	25.50	22.42
1965	.22	922.00	189.00	95.60	72.70	192.00	1603.00	130.00	550.00	31.60	9.48	189.00	328.70
1966	6.29	1.81	26.40	16.00	78.60	99.90	72.10	75.90	16.10	13.90	27.90	.02	36.08
1967	0.00	0.00	0.00	0.00	0.00	.01	18.60	19.60	518.00	372.00	23.10	469.00	117.79
1968	383.00	281.00	120.00	82.60	68.30	146.00	361.00	646.00	177.00	109.00	334.00	36.10	230.04
1969	292.00	802.00	402.00	209.00	358.00	813.00	737.00	1055.00	1628.00	60.90	7.13	248.00	547.14
AVERAGE	221.18	199.68	105.40	119.46	152.11	330.16	636.42	663.65	556.87	315.54	54.46	340.92	307.78
Q/QANN	6.099	5.329	2.906	3.294	3.823	9.104	16.983	18.300	14.861	8.701	1.502	9.098	100.000
COV VAR	1.937	2.152	1.752	1.543	1.974	1.370	1.455	1.589	1.311	2.206	1.912	2.133	.760
SKEW	2.501	3.283	2.377	2.130	4.238	1.972	2.114	2.902	1.591	3.277	3.367	2.897	.450
MAXIMUM	1754.00	2085.00	778.00	770.00	1629.00	1885.00	3400.00	4773.00	2721.00	3096.00	495.00	2874.00	787.69
MINIMUM	0.00	0.00	0.00	0.00	0.00	0.00	0.00	4.36	2.14	.62	0.00	0.00	13.32

327

C-9

STATISTICAL PARAMETERS FOR STATION 07170000 ELK RIVER NEAR ELK CITY, KANSAS
 ARKANSAS RIVER BASIN

	MONTH	ARTH. AVERAGE	LOG PARAMETERS			SKEW
			MEAN	VARIANCE	STD. DEV.	
1	OCT	221.18	.7633	3.9216	1.9803	-.8753
2	NOV	199.68	.7421	4.2304	2.0568	-.9864
3	DEC	105.40	.7743	3.5297	1.8787	-1.2300
4	JAN	119.46	.8588	3.6232	1.9035	-1.2581
5	FEB	152.11	1.3333	2.1019	1.4498	-1.8474
6	MAR	330.16	1.6873	2.1792	1.4762	-1.7081
7	APR	636.42	2.1550	1.5293	1.2367	-2.5448
8	MAY	663.65	2.3898	.4654	.6822	-.3205
9	JUNE	556.87	2.2826	.5596	.7481	-.4246
10	JULY	315.54	1.6504	.9519	.9757	.0208
11	AUG	54.46	1.0957	1.1884	1.0901	-2.0410
12	SEPT	340.92	1.0676	3.6020	1.8979	-.9720
13	ANNUAL	307.78	2.2802	.2703	.5199	-.8993

SAMPLE SIZE 31 YEARS

LOG NORMAL DISTRIBUTION STATION 07170000 ELK RIVER NEAR ELK CITY, KANSAS
 ARKANSAS RIVER BASIN

YEARS FLOW IS NOT EXCEEDED

MONTH	9 IN 10	2 IN 3	1 IN 2	1 IN 5	1 IN 10	1 IN 20	Q2 - Q10
OCT	2004.55	41.00	5.80	.12	.02	.00	5.78
NOV	2392.80	42.12	5.52	.10	.01	.00	5.51
DEC	1523.61	38.05	5.95	.16	.02	.00	5.92
JAN	1990.93	47.36	7.22	.18	.03	.01	7.20
FEB	1555.64	90.21	21.54	1.30	.30	.09	21.24
MAR	3800.41	209.23	48.68	2.78	.62	.18	48.05
APR	5500.27	484.75	142.89	12.99	3.71	1.32	139.17
MAY	1838.20	481.37	245.36	65.37	32.75	18.52	212.61
JUNE	1744.31	401.35	191.69	44.95	21.07	11.27	170.63
JULY	796.51	117.20	44.71	6.74	2.51	1.11	42.20
AUG	311.38	36.59	12.47	1.51	.50	.20	11.97
SEPT	3167.47	76.17	11.68	.29	.04	.01	11.64
ANNUAL	884.43	318.58	190.63	69.58	41.09	26.61	149.54

Q2 - Q10 IS THE 1 IN 2 YEAR FLOW MINUS THE 1 IN 10 YEAR FLOW

LOG-PEARSON TYPE III DISTRIBUTION FOR STATION 07170000
 ELK RIVER NEAR ELK CITY, KANSAS
 ARKANSAS RIVER BASIN

DATA SKEW K FACTORS

MONTH	SKEW	RETURN PERIOD					
		10 YRS	6.67 YRS	2 YRS	1.25 YRS	1.11 YRS	1.05 YRS
OCT	-.88	1.1423	.5097	.1520	-.7663	-1.3392	-1.8527
NOV	-.99	1.1252	.5157	.1662	-.7562	-1.3401	-1.8773
DEC	-1.23	1.0479	.5333	.2205	-.7092	-1.3376	-1.9391
JAN	-1.29	1.0544	.5321	.2163	-.7131	-1.3382	-1.9309
FEB	-1.85	.9068	.5463	.3008	-.6175	-1.3058	-1.9927
MAR	-1.71	.9220	.5459	.2930	-.6283	-1.3106	-1.9884
APR	-2.54	.7343	.5332	.3724	-.4880	-1.2303	-2.0129
MAY	-.32	1.2191	.4754	.0795	-.8096	-1.3218	-1.7591
JUNE	-.42	1.2039	.4835	.0951	-.8020	-1.3268	-1.7919
JULY	.02	1.2795	.4316	.0035	-.8408	-1.2841	-1.6508
AUG	-2.04	.8542	.5457	.3255	-.5814	-1.2881	-2.0640
SEPT	-.97	1.1221	.5165	.1685	-.7544	-1.3403	-1.8618
ANNUAL	-.90	1.1469	.5081	.1481	-.7689	-1.3390	-1.8581

329

FREQUENCY OF FLOWS

YEARS FLOW IS NOT EXCEEDED

MONTH	9 IN 10	2 IN 3	1 IN 2	1 IN 5	1 IN 10	1 IN 20	Q2 - Q10
OCT	1060.16	59.25	11.59	.18	.01	.00	11.58
NOV	1138.42	63.49	12.13	.15	.01	.00	12.12
DEC	553.43	59.73	15.44	.28	.02	.00	15.42
JAN	734.09	74.41	18.64	.32	.02	.00	18.62
FEB	444.65	133.46	58.81	2.74	.28	.03	58.53
MAR	1117.96	311.32	131.79	5.75	.57	.06	131.23
APR	1156.30	652.23	412.61	35.61	4.30	.46	408.31
MAY	1665.20	517.78	278.01	68.78	30.77	15.24	247.24
JUNE	1524.83	440.88	225.80	48.16	19.50	8.76	206.30
JULY	792.06	117.88	45.06	6.76	2.50	1.10	42.57
AUG	106.41	49.05	28.22	2.90	.49	.08	27.73
SEPT	1574.91	111.66	24.40	.43	.03	.00	24.37
ANNUAL	752.34	350.22	227.61	75.94	38.38	20.62	189.23

Q2 - Q10 IS THE 1 IN 2 YEAR FLOW MINUS THE 1 IN 10 YEAR FLOW

C-11

ELK RIVER BED MATERIAL

The following information on the bed material was prepared by the staff of the Kansas Office of the Soil Conservation Service. (Additional diagrams on sphericity and shape factor were included in the original data set. Also, the plotted size distribution curves were included.)

Cross Section 3-1: Inspected 4/2/80

The section was investigated with an 8 feet probe to determine the thickness of the material in the bottom. Five feet of water with 2-3 inches of ice was penetrated before the probe encountered the bottom. The probe was pushed another 3 feet without encountering any resistance. It is assumed the bottom consisted of alluvium because both banks consist of alluvium.

Cross Section 3-3: Inspected 5/2/80

This section had 6.5 feet of water overlying the bottom. The probe was pushed another 12 inches before it become too hard to push. The material was not rock (Ls, SS, or Sh). It is assumed the bottom consists of alluvium, because both banks consisted of alluvium.

Cross Section 3-4: Inspected 5/2/80

Rock was encountered 2.5 feet below the channel bottom. The rock is assumed to be sandstone since the Ireland Sandstone member of the Lawrence Shale formation is exposed in the bluff north of the channel. Rock is exposed not in either bank of the channel. The material is CL, ML type alluvium.

The armor layer is estimated at 6 inches except for those areas where the 28 inches x 15 inches x 6 inches boulders are stacked one on another. The armor layer consists of sandstone fragments with the following gradation.

<u>% Retained By Size</u>	<u>Particle Dimensions</u>	
10	28 x 15 x 6	Sandstone
25	6 x 6 x 2	Sandstone
35	8 x 3 x 2	Sandstone
30	2 x 2.5 x 0.75	Sandstone

An alternate bar is located on the left (north) side of the channel. This bar contained 3 feet boulders along with materials found in the armor layer.

A sample of the bedload was not taken because of the depth of water (3 feet).

Cross Section 3-5: Inspected 5/2/80

Rock was encountered 6 feet below the bottom of the channel. The rock is assumed to be sandstones. The depth of water made it impossible to see the rock.

The Amazonia limestone member of the Lawrence Shale formation was exposed in the right bank. Alluvium was observed in the left bank.

The armor layer is estimated to range from 4-6 inches across the channel bottom. The armor layer consists of sandstone and limestone fragments with the following gradation:

<u>% Retained By Size</u>	<u>Particle Dimensions (inches)</u>	
1	28 x 17 x 14	Located by right bank limestone
2	24 x 15 x 2.5	Located by right bank limestone
10	16 x 10.5 x 2	Sandstone

Cross Section 3-5: Inspected 2/5/80

<u>% Retained By Size</u>	<u>Particle Dimensions (inches)</u>	
2	13 x 7 x 2	Sandstone
15	8.5 x 4 x 2	Limestone
30	5.5 x 3.5 x 2	Limestone
40	3.5 x 3.5 x 1	Limestone

A sample of the underlying material was taken and will be sent to the SML at Lincoln, Nebraska, for sieve analysis.

Cross Section 11-1: Inspected 5/2/80

The probe was pushed 18 inches below the bottom of the channel. The material pushed like shale. It is assumed to be shale of the Lawrence Shale formation. Alluvium was observed in both banks.

The armor layer is estimated at 4-6 inches. This layer consists of sandstone, limestone, and shale fragments with the following gradation:

<u>% Retained By Size</u>	<u>Particle Dimensions (inches)</u>	
15	43 x 38 x 1.75	Gray silty clay shale
5	40 x 26 x 11	Sandstone
10	28 x 20 x 4	Limestone
10	23 x 14 x 2	Limestone
5	13 x 11 x 1.5	Gray silty clay shale

<u>% Retained By Size</u>	<u>Particle Dimensions (inches)</u>	
20	9.5 x 6.5 x 1.5	Sandstone
20	6 x 4.5 x 1.5	Sandstone
10	4.5 x 2.5 x 3	Sandstone
5	4 x 3 x 1	Limestone

A sample of the underlying material was taken and will be sent to the SML at Lincoln, Nebraska, for sieve analysis.

Cross Section 11-7: Inspected 5/2/80

It was not possible to determine what is in the bottom at this section because the water depth is too deep (6-8 feet). Probing was done near the shore with the probe penetrating to 8 feet without encountering resistance. Alluvium was observed in both abutments. No rock was observed 1000' upstream or downstream from the bridge. Shale is assumed to underlie the bottom.

Cross Section 11-8: Inspected 6/2/80

Shale was encountered 4 feet below the bottom of the channel. The shale is the Lawrence Shale. The shale is a gray silty clay shale. Alluvium is in both banks of the channel.

The armor layer is estimated to range from 4-8 inches. This layer consists of limestone and sandstones fragments with the following gradation:

<u>% Retained By Size</u>	<u>Particle Dimensions</u>	
1	24 x 14 x 4	Sandstone
5	13 x 13 x 2	Sandstone
35	9 x 6 x 1	Limestone
5	7 1/2 x 12 x 4	Sandstone
40	7 x 5 x 1.75	Sandstone
14	4 x 3 x 2.5	Sandstone

A sample of the underlying material was taken and will be sent to the SML at Lincoln, Nebraska for sieve analysis.

Cross Section 11-9: Inspected 6/2/80

Under 6 inches of ice was 3.5 feet of water at this section. Probing indicates 6 to 8 feet of gravelly material on the bottom with softer material underneath to the depth probed (12-16 inches). Because of the water depth, samples were not taken. Alluvium was observed in both banks.

Between Cross Sections 11-11 and 11-12 at Station 2180:
Inspected 6/2/80

Shale was encountered 12 inches below the bottom of the channel. The shale is the Lawrence Shale formation. The shale is fairly soft for 6 inches and then a very hard layer is encountered. This layer is assumed to be sandstone. The shale is a gray silty clay shale and is exposed 2 feet above the channel bottom in the right bank. Alluvium is observed in both banks. The right bank is vertical and eroding.

The armor layer is 4 inches thick and consists of limestone and sandstone fragments with the following gradation:

<u>% Retained By Size</u>	<u>Particle Dimensions (inches)</u>	
2	8 x 5.5 x 3	Sandstone
3	7 x 5 x 1	Sandstone
40	5 x 3 x .5	Shaley Limestone
4	4.5 x 3 x 2.5	Limestone
25	3.5 x 2.5 x .5	Sandstone
26	3 x 2 x .25	Sandstone

A sample of the underlying material was taken and will be sent to the SML of Lincoln, Nebraska for sieve analysis. A soil sample was taken from the right bank.

Cross Section 11-12: Inspected 6/2/80

Shale is exposed in the bottom of the channel. A few 2 feet sandstone fragments are scattered over the bottom. No armor or bedload exists at this location. Depth of weathering in the shale is 8". The shale is the Lawrence Shale formation. Alluvium is exposed in both banks. The shale is exposed 12 inches above the waterline in the left abutment. The shale is a gray silty clay shale. The left bank is vertical and eroding.

Between Cross Sections 11-14 and 11-15 at Station 2080: Inspected 6/2/80

Shale was encountered 3.5 feet below the bottom of the channel. This shale is the Lawrence Shale formation. It is a gray silty clay shale.

The armor layer is 4-6 inches. The layer consists of limestone and sandstone fragments with the following gradation:

<u>% Retained By Size</u>	<u>Particle Dimensions (inches)</u>	
1	11.5 x 6 x 3	Limestone
5	8 x 6 x 1	Limestone
25	5 x 3 x 2	Limestone

<u>% Retained By Size</u>	<u>Particle Dimesnions (inches)</u>	
30	4 x 3 x .75	Limestone
19	2.5 x 2.25 x .25	Limestone
20	2.5 x 2 x .75	Limestone

Alluvium is exposed in both abutments. A sample of the underlying was taken and will be sent to the SML in Lincoln, Nebraska for sieve analysis.

Cross Section 11-15: Inspected 6/2/80

Rock was encountered 12 inches below the bottom of the channel. This rock is interpreted to be limestone.

The armor thickness is 6 inches with 6 inches of bedload underlying the armor. The armor layer consists of limestone and sandstone fragments with the following gradation:

<u>% Retained By Size</u>	<u>Particle Dimensions (inches)</u>	
5	22 x 17 x 4	Limestone
5	16 x 11 x 3	Shaley Sandstone
15	11.5 x 7 x 4	Limestone
15	9 x 7 x 3	Limestone
10	8 x 5 x 1.5	Limestone
10	3.5 x 2.5 x 2	Sandstone
5	3 x 2.5 x 1	Sandstone
25	3 x 2.25 x .25	Limestone
10	1.5 x 1.75 x .5	Limestone

Alluvium exposed in both banks of the channel.

Cross Section 13-2: Inspected 2/6/80

Limestone is exposed in the bottom of the channel. The limestone is clear of any armor or bedload. This limestone is the Amazonia limestone member of the Lawrence Shale formation.

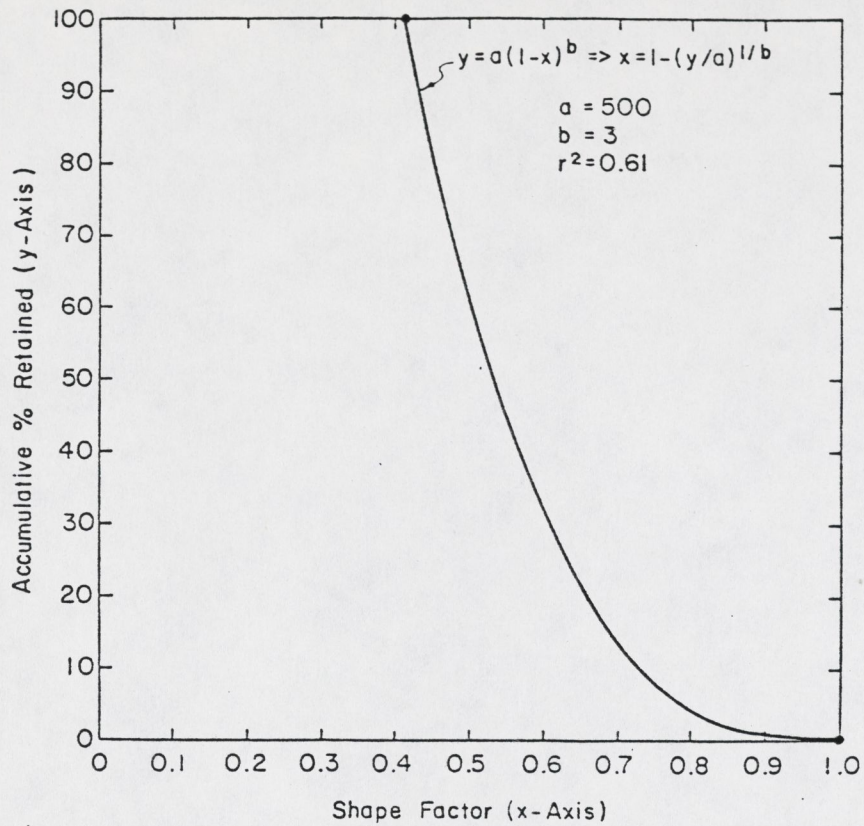
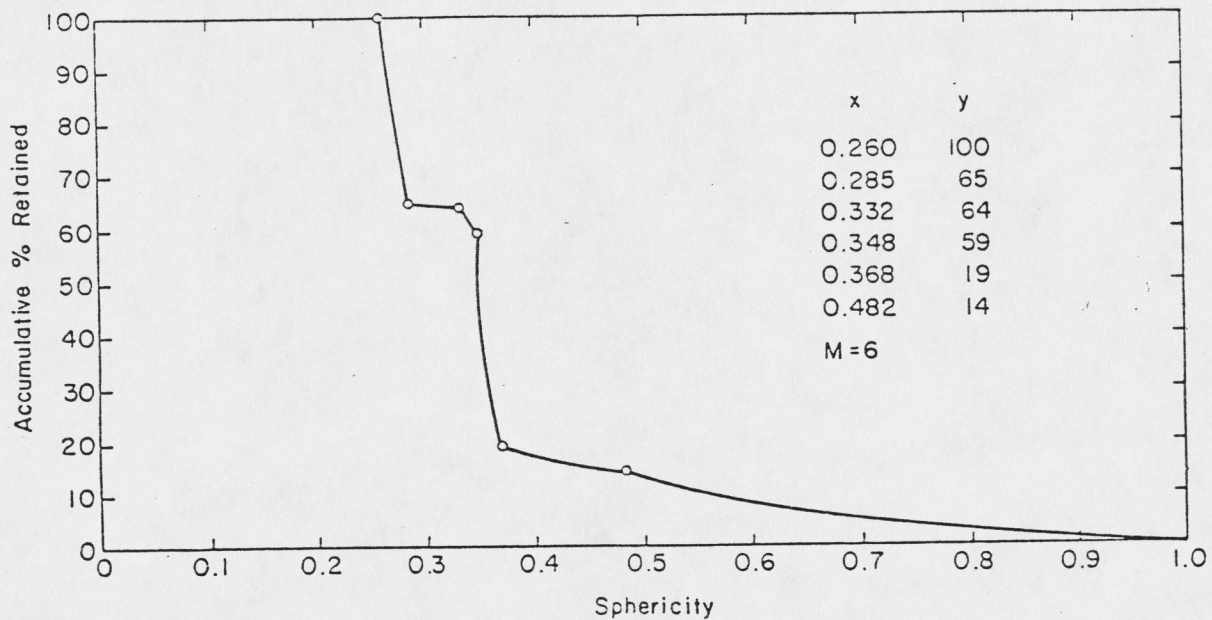


Figure C- . Channel bed material armor, shape factor-composite curve fit. $S_f = [(a \cdot b \cdot c)^{13}] / a$ where a is the longest axis of rock particle and b and c are the remaining axis of rock particle (Elk River, Kansas).



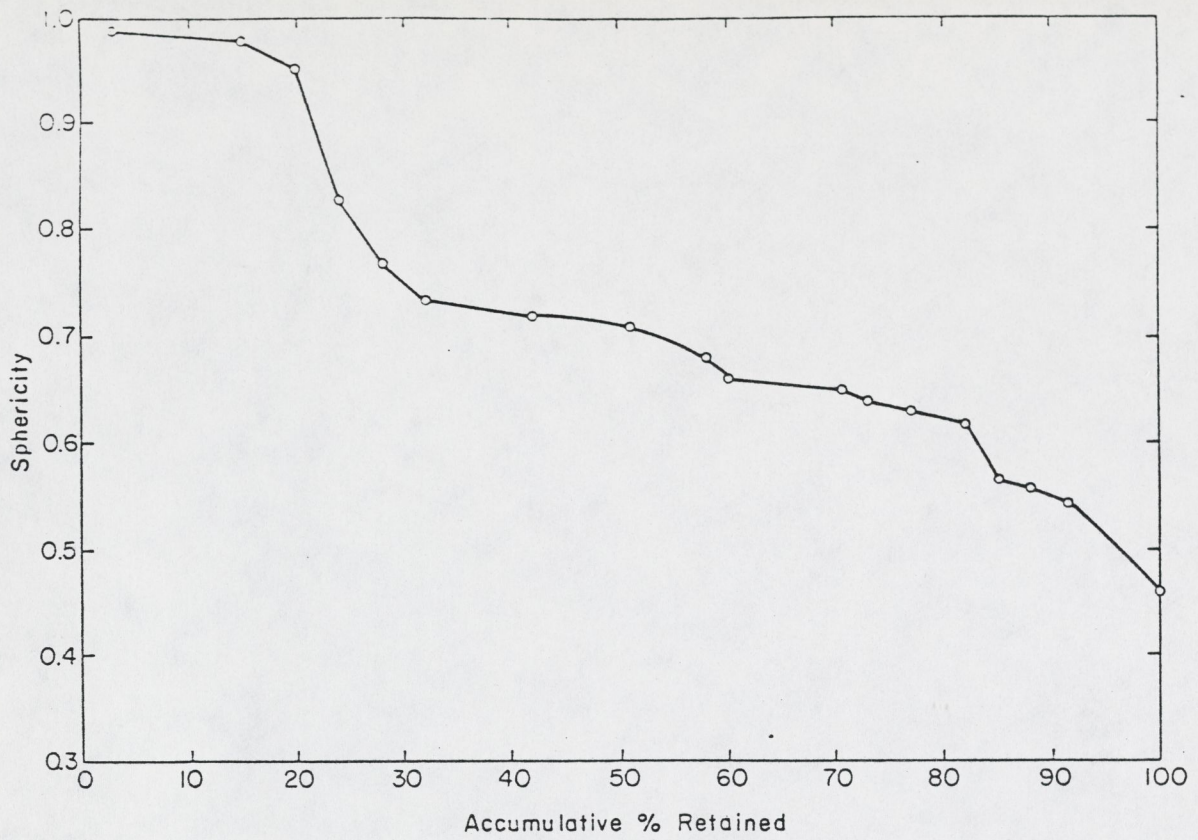
MECHANICAL ANALYSIS - PERCENT FINER

LOCATION	6"	3"	2"	1 1/2"	1"	3/4"	1/2"	3/8"	No. 4	No. 10	No. 40	No. 60	No. 200
Stream, x-sect 3-5, Bedload	8	9	13	14	17	24	34	45	53	64	71	81	100
Stream, x-sect 11-1 Bedload	7	7	8	9	11	19	32	48	57	70	78	87	100
Stream, x-sect 11-8, Bedload	4	5	6	7	9	17	28	39	45	56	65	84	100
Stream, x-sect 11-11, Bedload	8	9	12	13	17	28	43	59	68	80	89	98	100
Stream, x-sect 11-15, Bedload	7	8	10	12	16	24	36	53	64	82	90	98	100

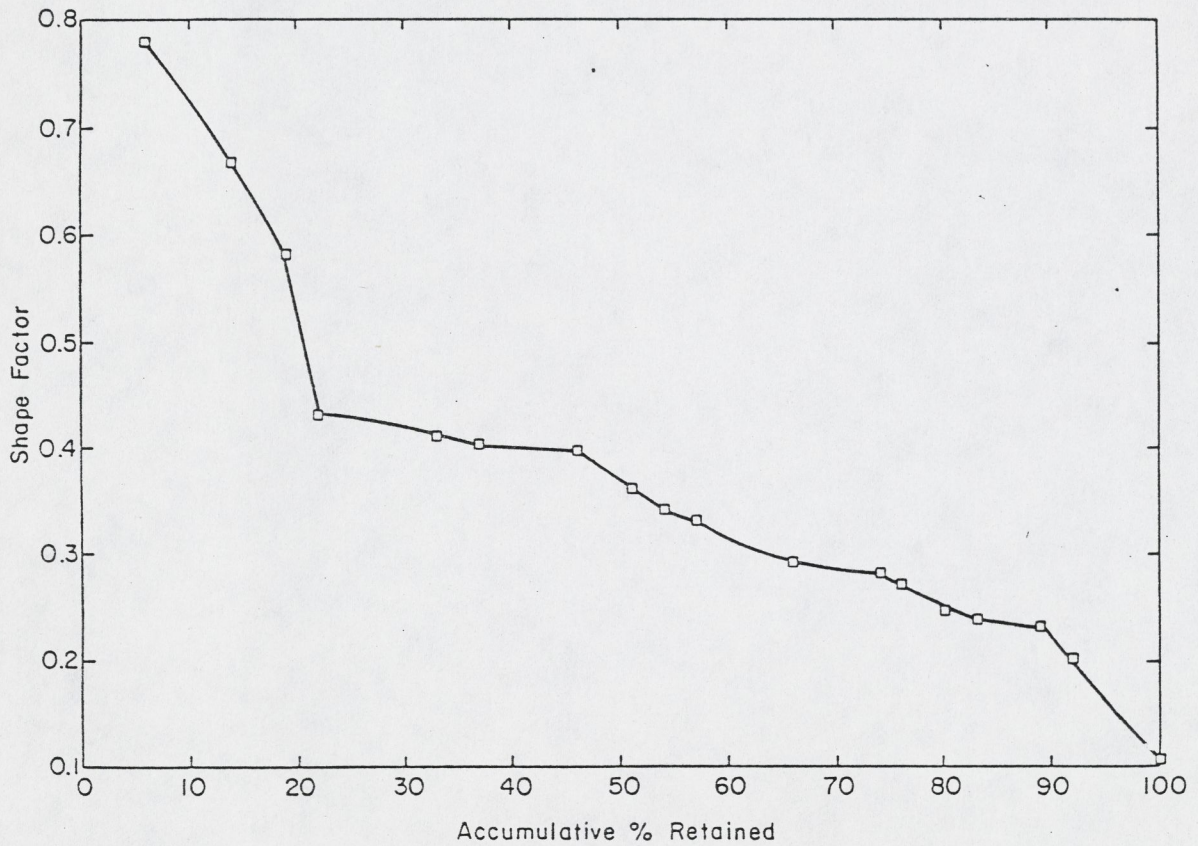
336

C-17

Soil mechanics laboratory data (lower Elk River stream channel).



Sphericity compared with percent retained (Elk River Watershed, Kansas; cross section 11-8, 3/31/80 C.T.)



Shape factor compared with percent retained (Elk River Watershed, Kansas; cross section 11-8, 3/31/80 C.T.)

ELK RIVER SEDIMENT YIELD

The following information on sediment yield was developed by the staff of the Kansas Office of the Soil Conservation Service. Only one suspended sediment measurement has been made and is included.

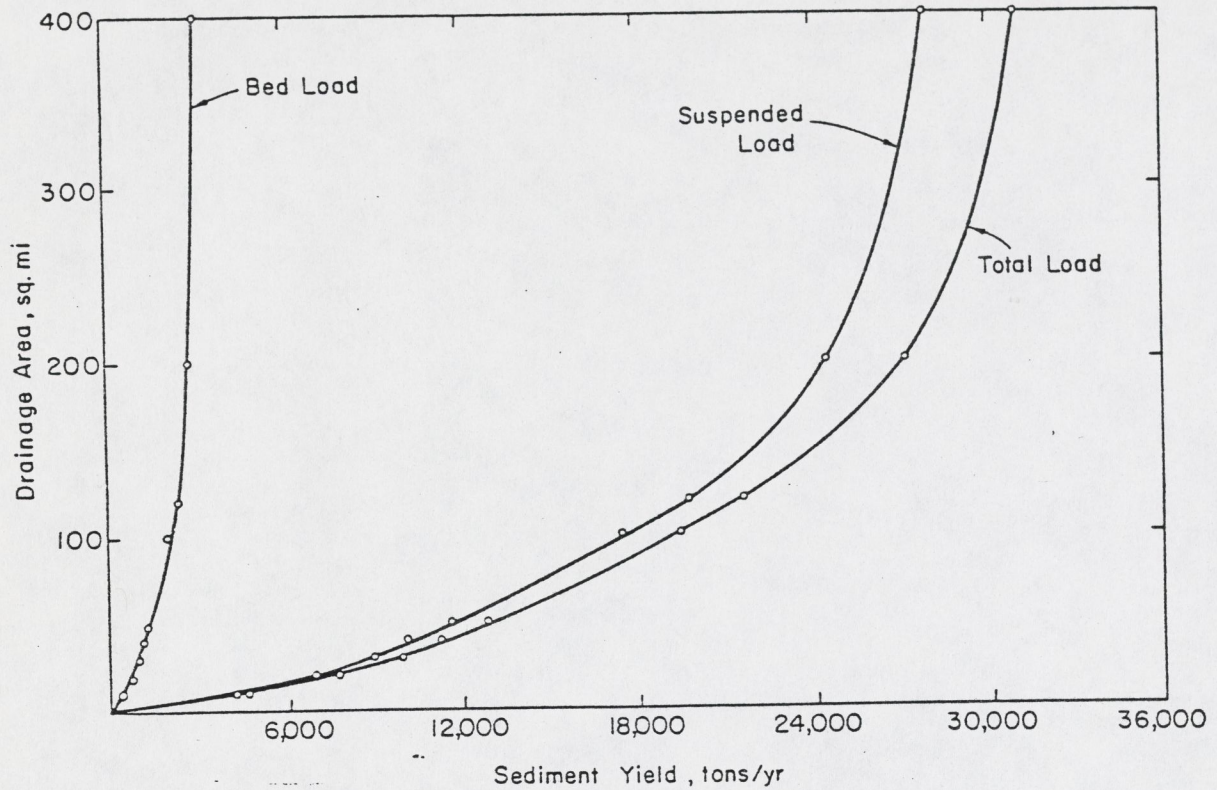


Figure C- . Drainage area compared with sediment yield without project (Elk River, Kansas 1980).

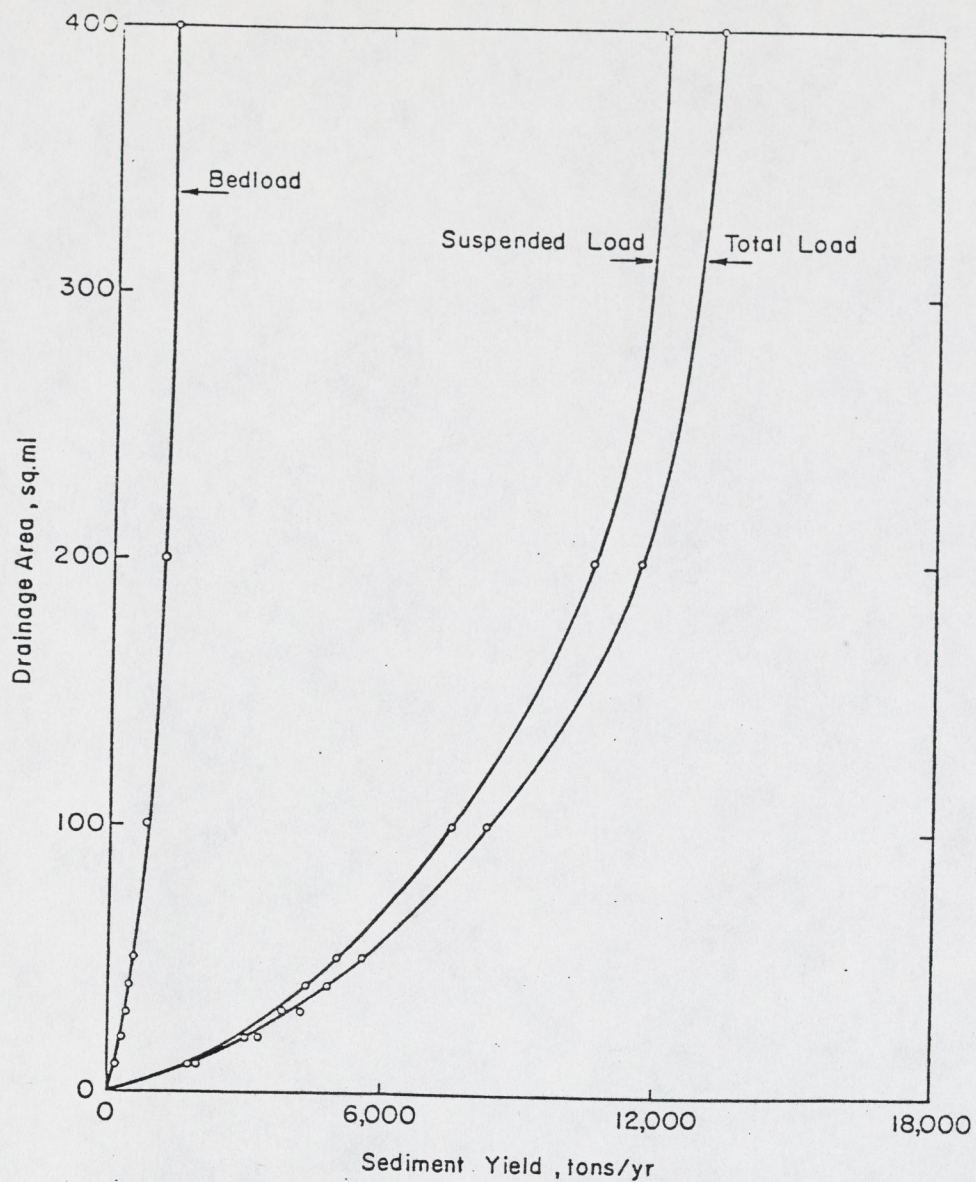


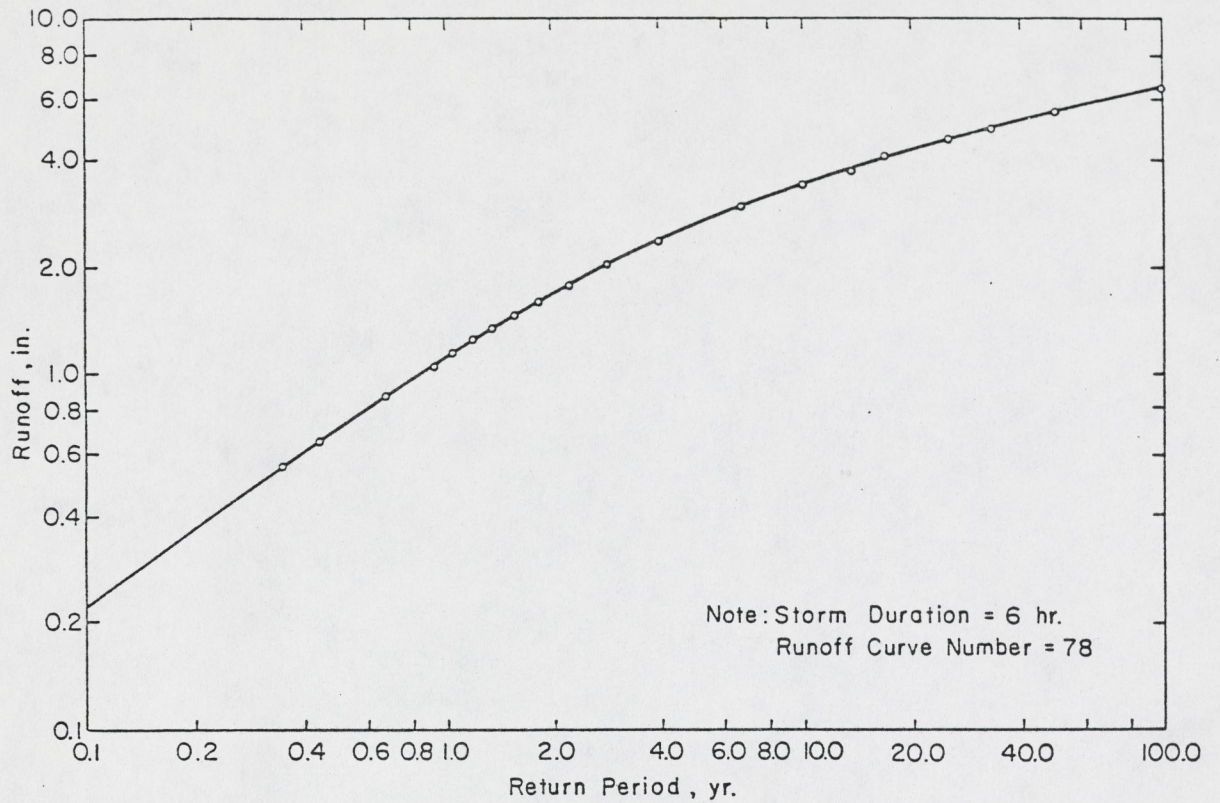
Figure C- . Drainage area compared with sediment yield with project (Elk River, Kansas 1980).

Suspended Sediment Measurement

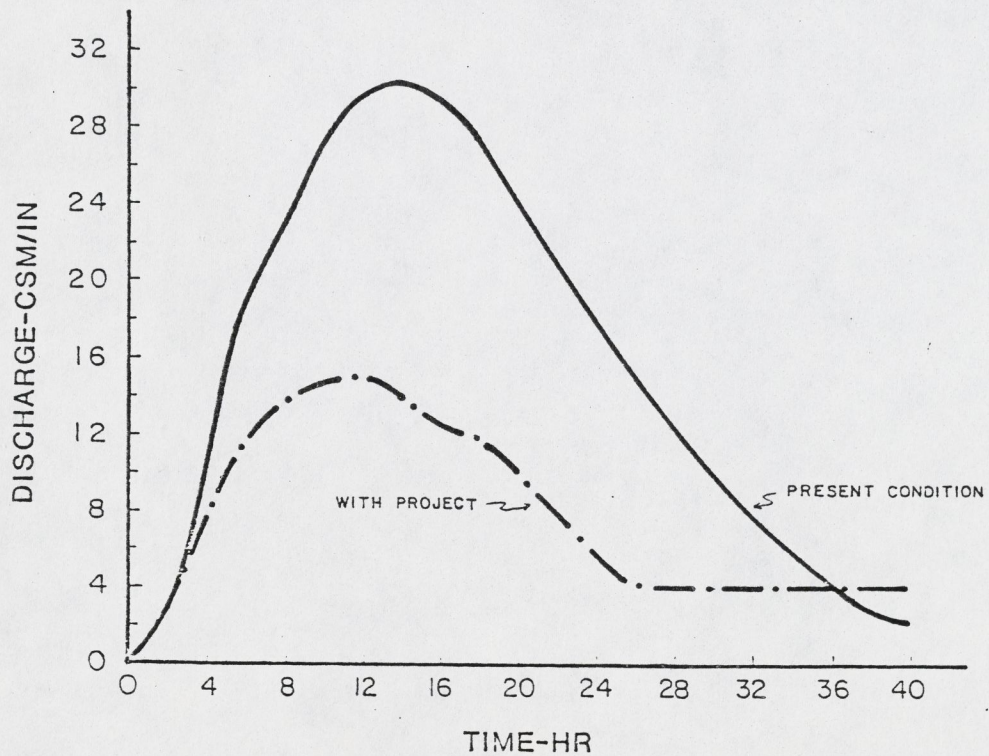
07169800-Elk River at Elk Falls, Kansas

Date	Time	Discharge (cfs)	Specific conductance (micromhas)	Sediment suspended (MG/L)	Sediment discharge (T/Day)
11-15-77 (1978 water year)	1430	154	460	61	25

ELK RIVER HYDROGRAPHS AND ELEVATION VS. DISCHARGE DATA



Frequency-runoff curve (annual series) (Elk River, Kansas).



Unit discharge hydrograph (Elk River, Kansas).

Table C. Unit Discharge Hydrograph (Elk River, Kansas)

[See preceding page]

^a discharge in cfs per sq. mi. per inch of runoff applicable for reaches 3 and 11.

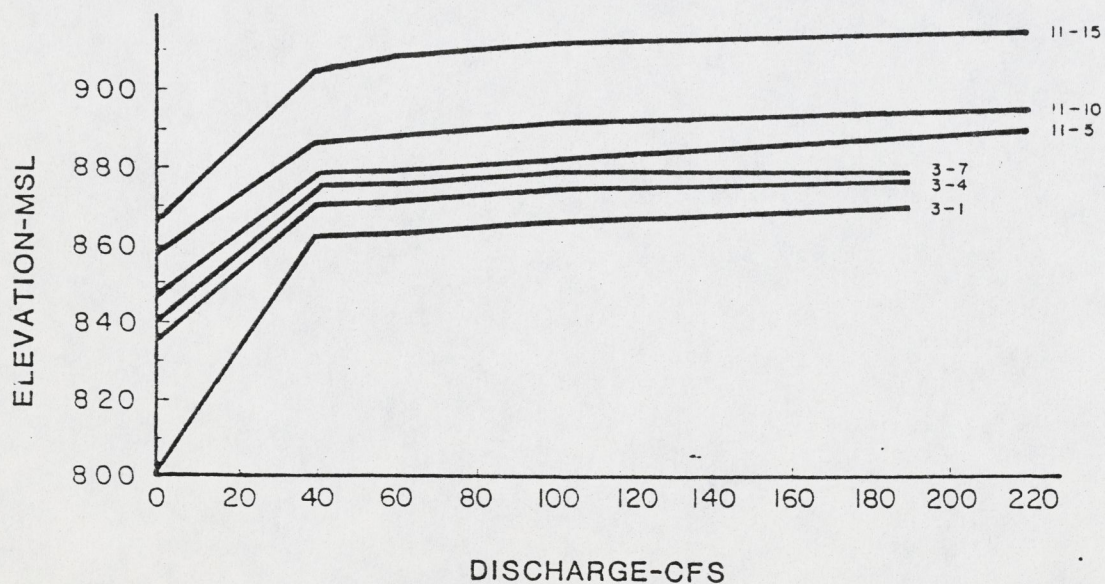
^b drainage area controlled equal 239.83 sq. mi. with average uncontrolled release rate from structures equal to 20 csm.

Unit Discharge Hydrograph Table (Elk River, Kansas)

hr. Time	Discharge-csm/in.*		Reach	Drainage Area Sq. Mi.
	Present	with Project**		
0	0	0		
2	3.69	2.69		
4	11.32	8.21	11 U.S.	284.5
6	18.51	11.92	11 D.S.	297.1
8	23.02	13.93	3 U.S.	341.0
10	26.66	14.61	3 D.S.	405.3
12	29.55	14.74		
14	30.29	13.93	U.S. - Upstream	
16	29.41	12.22	D.S. - Downstream	
18	27.21	10.91		
20	24.23	9.36		
22	21.01	7.88		
24	17.85	5.45		
26	14.87	4.00		
28	12.12			
30	9.72			
32	7.55			
34	5.56			
36	3.86			
38	2.70			
40	2.00			

*discharge in cfs per sq. mi. per inch of runoff applicable for reaches 3 and 11.

**drainage area controlled equal 239.83 sq. mi. with average uncontrolled release rate from structures equal to 20 csm.



Elevation-discharge plots (Elk River, Kansas).

Table C-. Elevation-discharge table (Elk River, Kansas)^a

^aSpatially-varied steady flow assumption was used.

^bContributing drainage area changes at low flows due to confluence.

Elevation-Discharge Table (Elk River, Kansas)^a

x-sec	DA Reach	Q=0		Bankfull		ft. ² A	fps \bar{v}	← Out-of-Bank Flow →									
		Elev-msl	E-msl	Q				40 csm		60 csm		100 csm		190 csm		220 csm	
				cfs	csm			Q-cfs	E-msl	Q-cfs	E-msl	Q-cfs	E-msl	Q-cfs	E-msl	Q-cfs	E-msl
3-1	405.3	832.2	855.5	6500	16	1862	3.49	15,708	863.6	23,562	865.8	40,530	865.6	77,007	870.5	—	—
3-2	392.2	832.2	855.0	3750	10	2086	1.80	15,688	866.5	23,532	867.4	39,220	868.9	74,518	871.8	—	—
3-3	390.5	833.2	864.5	9200	24	3586	2.57	15,620	867.9	23,430	869.0	39,050	870.9	74,195	873.6	—	—
3-4	388.0	835.9	866.0	9700	25	3123	3.11	15,520	869.5	23,280	871.1	38,800	873.1	73,720	875.9	—	—
3-5	387.6/341.3	837.2	867.5	7250	21	2180	3.33	13,652	871.9	20,478	873.4	38,760	875.5	73,644	878.1	—	—
3-7	341.0	840.0	869.5	7250	21	3232	2.24	13,640	874.0	20,460	875.3	34,100	877.3	64,790	879.9	—	—
<u>Reach 11</u>																	
11-1	297.1	842.5	873.0	8750	29	4630	1.89	11,884	874.9	17,826	876.5	29,710	879.6	—	—	65,362	882.4
11-3	297.1	842.5	871.0	6300	21	3078	2.05	11,884	875.1	17,826	876.8	29,710	880.6	—	—	65,362	887.9
11-4	296.6	844.8	871.0	5500	19	2261	2.43	11,864	876.6	17,796	878.7	29,667	881.7	—	—	65,252	888.6
11-5	296.3	845.5	872.5	6500	22	1802	3.61	11,852	877.7	17,778	879.8	29,630	882.4	—	—	65,166	888.9
11-7	295.2	849.2	877.5	8300	28	2667	3.11	11,808	880.0	17,712	881.9	29,520	884.3	—	—	64,944	889.8
11-8	294.0	856.0	881.5	9600	33	3924	2.45	11,760	883.3	17,640	885.5	29,400	887.5	—	—	64,680	891.6
11-9	291.8	857.2	881.0	6750	23	2091	3.23	11,672	885.8	17,508	888.1	29,180	890.5	—	—	64,196	891.8
11-10	291.2	858.5	880.0	5000	17	1999	2.50	11,648	887.2	17,472	889.6	29,120	892.4	—	—	64,064	895.9
11-11	290.8	860.6	886.0	7000	24	2730	2.56	11,632	889.3	17,448	890.8	29,080	893.6	—	—	63,976	897.2
11-12	288.4	861.0	869.0	8500	29	2074	4.10	11,536	891.2	17,304	893.1	28,840	895.5	—	—	63,448	899.0
11-14	286.3	863.2	891.0	8500	30	3546	2.40	11,452	893.2	17,178	895.6	28,630	897.8	—	—	62,986	901.2
11-15	285.7	865.6	895.0	10,750	38	3030	3.88	11,428	895.5	17,142	898.2	28,570	901.0	—	—	62,854	904.7

^aSpatially-varied steady flow assumption was used.

^bContributing drainage area changes at low flows due to confluence.