

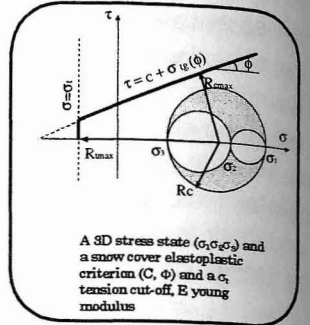
STATIC EQUILIBRIUM APPROACH : 3D STABILITY INDEXES F_{SHEAR} , $F_{TENSION}$

based on general (3D) stress state and static equilibrium at each point of the snow cover

It assumes a granular material with an elastoplastic behaviour (in our case : a Mohr Coulomb plasticity criterion with tension cut off)

$$F_s = \frac{R_{tmax}}{R_c} = \frac{\sin \Phi (2C \cot g\Phi + \sigma_1 + \sigma_3)}{\sigma_1 - \sigma_3}$$

$$F_t = \frac{R_{tmax}}{R_c} = \frac{\sigma_1 + \sigma_3 - 2\sigma_t}{\sigma_1 - \sigma_3}$$

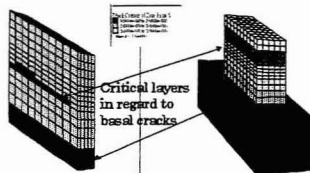


- Theoretical instability in shear [tension] is assumed when $F_s [F_t]$ is less than 1,
- it could be judicious using security factor a of 1.5 or 2
- The indexes are implemented in a finite differences code which estimates the stress inside the snow cover.
- Maps of indexes give a localisation of potential instable layers and areas

SOME APPLICATIONS

The "numerical Rutchblock" : knowing which layer leads to instability

Calculation of the F_s indexes to point out the shear critical layers by implementation in a stress strain code



E(MPa)	ϕ °	C(kPa)	σ_t (kPa)	e(cm)	density	
0.8	28	1.5	0.9	20	140	cche8
4	36	4	2.9	20	230	cche7
4	28	1.5	1.3	6	220	cche6
4	28	1.5	1.3	12	220	cche5
8	28	1.5	1.9	12	250	cche4
80	29	3	5.3	20	370	cche3
200	41	6	6	25	390	cche2
600	53	12	8.3	25	440	cche1

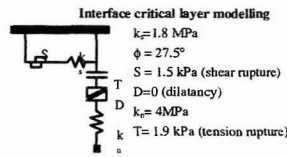
Mesh Model
half Rutchblock on a 35° rigid bedrock slope
no other loading than gravity

F_s is the smallest for layer 4 and 1.
Slab mechanism should be due to shear failure in layer 4

A 3D snow cover slope : effect of local defect of mechanical properties in the critical layer

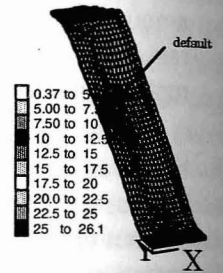
Snow cover meshed geometry:

- homogeneous top slab (50 cm) under gravity ($E=4$ Mpa, $C=5$ kPa, $f=38.7$, $\sigma_t=3.6$ kPa, $d=250$ kg/m³)
- The thin critical layer is assumed to be an interface without weight and depth



default size (m ²)	F_t
0	1.1
1*2	2.02
1*3	1.26
1*4	1.19
2*2	1.13
2*3	1.08

F_t as a function of the size of the defect (Interface Properties divided by 2)



- crest or comb (40 m wide, 100 m long)
- slope angle of 35° (majority of slabs triggering)
- boundary conditions :
4rigid bed rock
4only vertical displacements due to elastic behaviour allowed at the lateral boundary

The static equilibrium approach shows only the release model of the Griffith approach.

CRACK STABILITY WITH GRIFFITH APPROACH

Hypothesis : based on the evolution of an initial crack due to creep or skier triggering

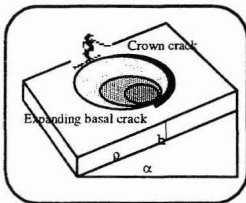
Two different release modes : scenarios may arrive according whether or not the basal crack fulfills Griffith's instability criterion before crown crack opening

Shear stress at the slab substrate interface :

$$\tau = \frac{1}{2} \rho \cdot g \cdot h \cdot \sin(2\alpha)$$

tensile stress in the slab at the upper tip of the basal crack:

$$\sigma = \rho \cdot g \cdot a_s \cdot \sin(\alpha)$$



Release Mode 1 : basal crack expands, and σ increases with the skier progression across the slab.

A tensile crack starts opening at the crown of the basal crack when the tensile stress σ reaches the tensile failure stress σ_t of the snow. Crack propagation is controlled by the energy balance between stresses relaxation and free surface creation (Jamieson and Johnson 91)

$$\text{so } a_s = \sigma_t / \rho \cdot g \cdot \sin(\alpha),$$

where a_s is the size of the basal crack in mode 1.

Mode 1 splitting slab fragments

Mode 2 breaking wave slab

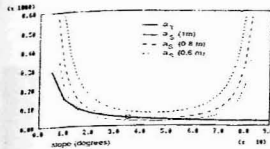
Release Mode 2 : basal crack meets alone Griffith's condition for unstable growth before the tensile stress σ at the crown reaches the critical value σ_t for tensile failure. the basal crack becomes unstable when the shear stress concentration factor $\tau(\pi a_s)^{1/2}$ exceeds the interface shear toughness K_{if} . The basal crack is already overcritical as the crown crack expand and the whole slope release. This suddenly open may be responsible of the "bang" heard in some avalanche.

$$\text{So } a_s = 1/\pi \cdot (2 K_{if} / \rho \cdot g \cdot h \cdot \sin(2\alpha))^2,$$

where a_s is the size of the basal crack in mode 2.

TRANSITION BETWEEN THE TWO TRIGGERING MODES

The competition between slab splitting and breaking wave modes is illustrated on this figure



For a given slope, the type of avalanche is determined by the smallest of the 2 crack size a_1 or a_2

The transition between the two modes is obtained for $a_1 = a_2$ and $da_1/d\alpha = da_2/d\alpha$ that leads to $\alpha = 35.6^\circ$ independent of the snow properties (thoughness, density ...)

EXPERIMENTAL STUDIES ON REDUCED MODEL WITH SAND

Experimental Installation

The installation is dimensioned by similitude equations where the symbol * concerns the sand properties and the reduced installation

$$\left\{ \begin{aligned} a_1 &= \frac{1}{\pi} \left(\frac{2Kc}{\rho g h \sin(2\alpha)} \right)^2 \\ a_2 &= \frac{1}{\pi} \left(\frac{2Kc^*}{\rho^* g h^* \sin(2\alpha)} \right)^2 \end{aligned} \right. \Rightarrow \frac{h^*}{h} = \frac{\rho Kc^*}{\rho^* Kc} \sqrt{\frac{a_1}{a_2^*}}$$

$$\left\{ \begin{aligned} \rho &= 400 \text{ kg/m}^3 \\ \rho^* &= 1400 \text{ kg/m}^3 \Rightarrow \frac{\rho}{\rho^*} = 3.5 \end{aligned} \right. \quad \left\{ \begin{aligned} a_1 &= \frac{\sigma_f \rho^*}{\sigma_f^* \rho} \\ a_2 &= \frac{\sigma_f^* \rho}{\sigma_f \rho^*} \end{aligned} \right.$$

For $h = 1\text{m}$, and $\alpha = 35^\circ$: We find: $a_1^* = 4,62\text{cm}$
 $h^* = 2,723\text{mm}$

• Around 30 experimental tests were performed in order to find the conditions of occurrence of the 2 modes of release

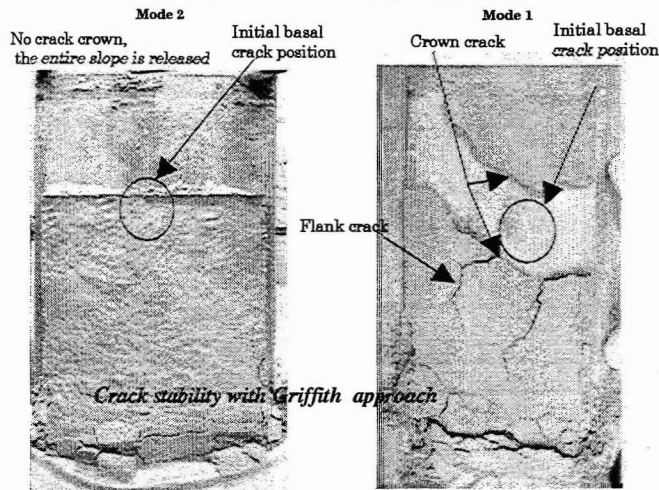
• The initial basal crack was made of a disk of polyethylene leave of 20 cm diameter.

• The slab layer is made of hand-compacted wet sand in order to get a non zero cohesion of the grain. (It is then inclined to find the slope which induced a release in mode 1 or in mode 2.

• The critical layer is layer of one mm of dry sand raining-like deposited.

As a result, there was not clearly a value of the slope angle which leads to one or other mode. Both were equally observed.

This may be due to the difficulty of realising repeatable samples (density, homogeneity...)



The measured α (between 40° to 60°) is always greater than the theoretical one

Dimensions $l=64\text{ cm}$, $L=80\text{ cm}$, $h = 5\text{ cm}$, properties: $\rho_{\text{sand}}=1400\text{ kg/m}^3$, $\sigma_f=11\text{N}$, $K_c=9.8\text{ Pa m}^{1/2}$.

CONCLUSION

- The static equilibrium approach can be used to predict which layers are potentially instable. It may predict the instability for the crown crack and the flank one in the slab, around the shape of the basal crack. But it doesn't predict the expansion of this last one. It may be reproduced by the mode 1 experimental tests.
- The Griffith rupture mechanic approach may predict two different release modes: the splitting slab fragment one and the breaking wave slab one.
- The experiments performed seem to show the two types of release. But it doesn't present clear parameters of transition due to the difficulty of doing those tests with a good reproducibility.
- However it is important to continue this work to get more repeatable results, especially experimental ones, and find amelioration of the installation

References

- Schillingier, L. Daudon, D., Flavigny, E., 1998. 3D modelling of snow slabs stability. In "25 years of snow Avalanche Research". Hestners E. (editor). NGI publications, Voss, Norway, 234-237.
- McClung, D., Schaerer, P., 1993. The Avalanche Handbook. The Mountaineers, Seattle.
- Mellor, M., 1974. A review of basic snow mechanics. *IAHS Pub.114*, 251-291.
- Roch, A., 1966. Les variations de résistance de la neige. *IAHS Pub.n°69*, 86-99.
- Sommerfeld, R.A., 1973. A correction factor for the Roch's stability index of slab avalanche release. *J. Glaciol...* 17(75), 145-147.
- Gubler H. Measurements and modelling to improve our understanding of avalanche formation. swiss federal institute for snow and avalanche research.
- Gibson and Ashby 1988, the mechanic of foam: basic results, cellular solids - structures and properties, pergamon press
- Daudon D, Flavigny E. 2000A way of studying snow-slab stability by numerical modeling. Innsbruck, Austria, may 2000.

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